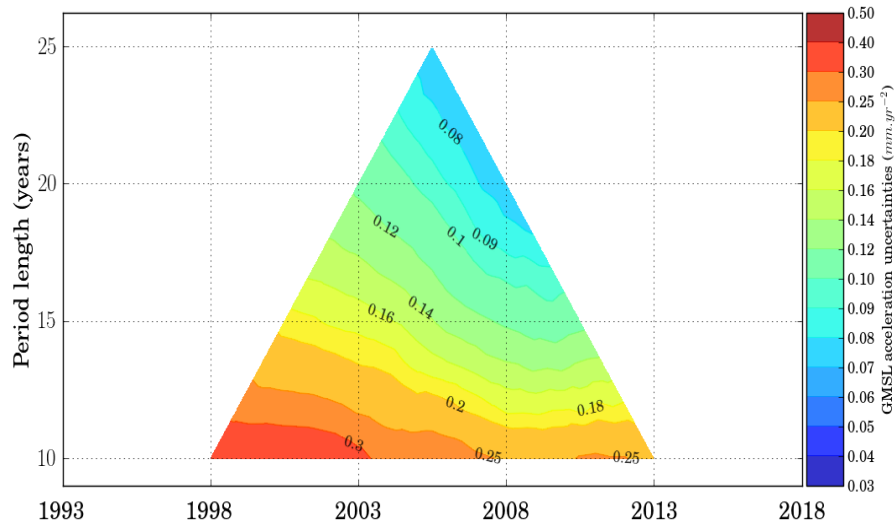


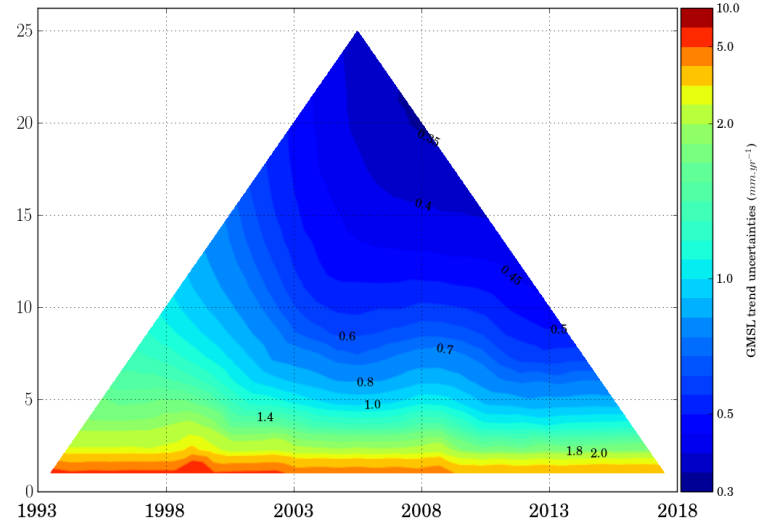
Uncertainty in Global mean sea level from Satellite Altimetry

Approach: estimate directly the GMSL error variance-covariance matrix from an error budget of the altimetry system. Then use the matrix to evaluate the uncertainty on some metrics like sea level trends and sea level acceleration (context: ESA climate change Initiative coastal sea Level project)



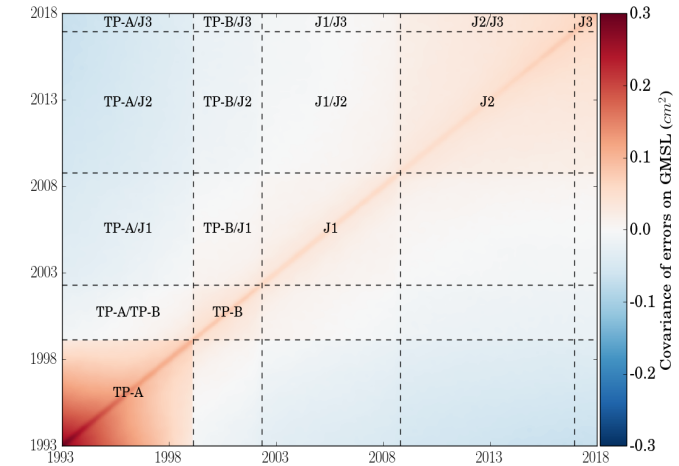
Center of the period over which accelerations are estimated

Uncertainty in sea level acceleration deduced from the var-covar matrix



Center of the period over which trends are estimated

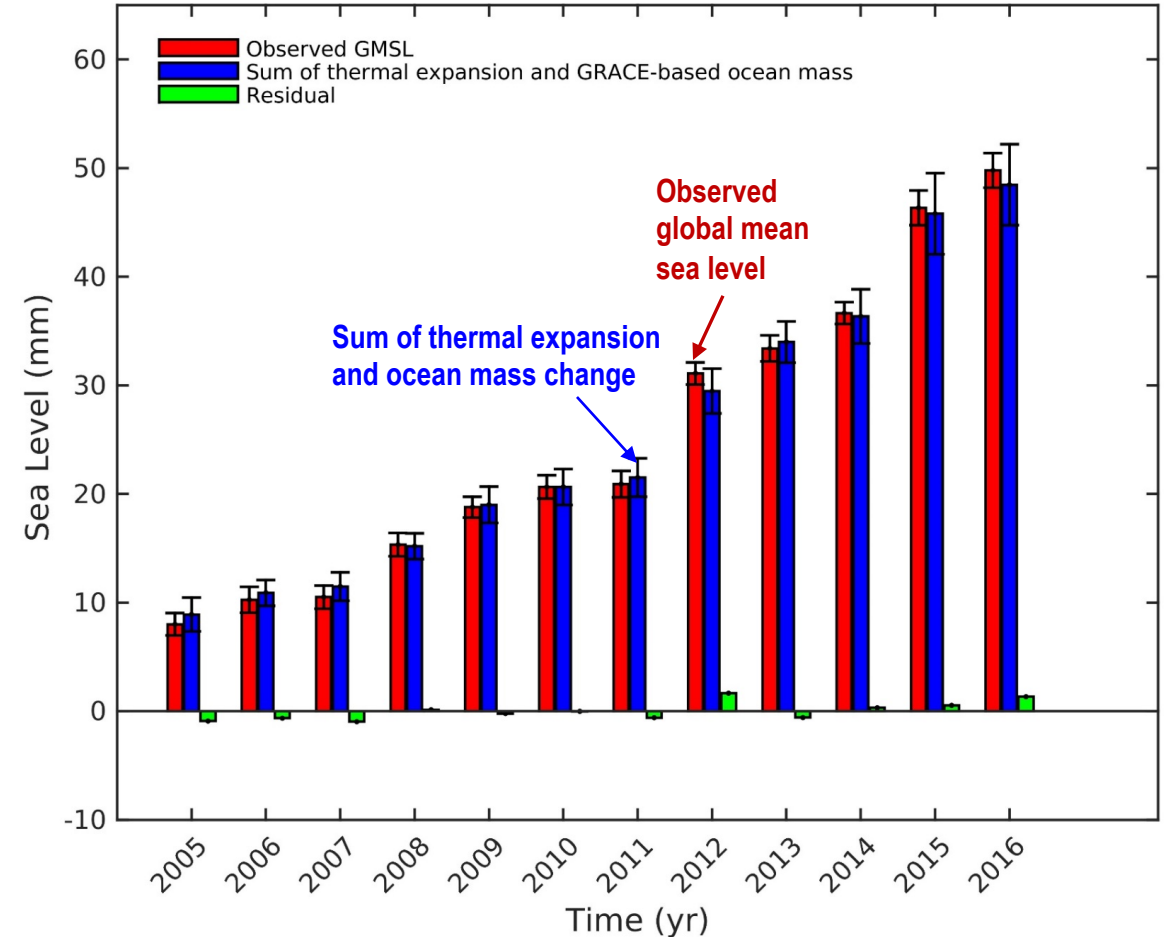
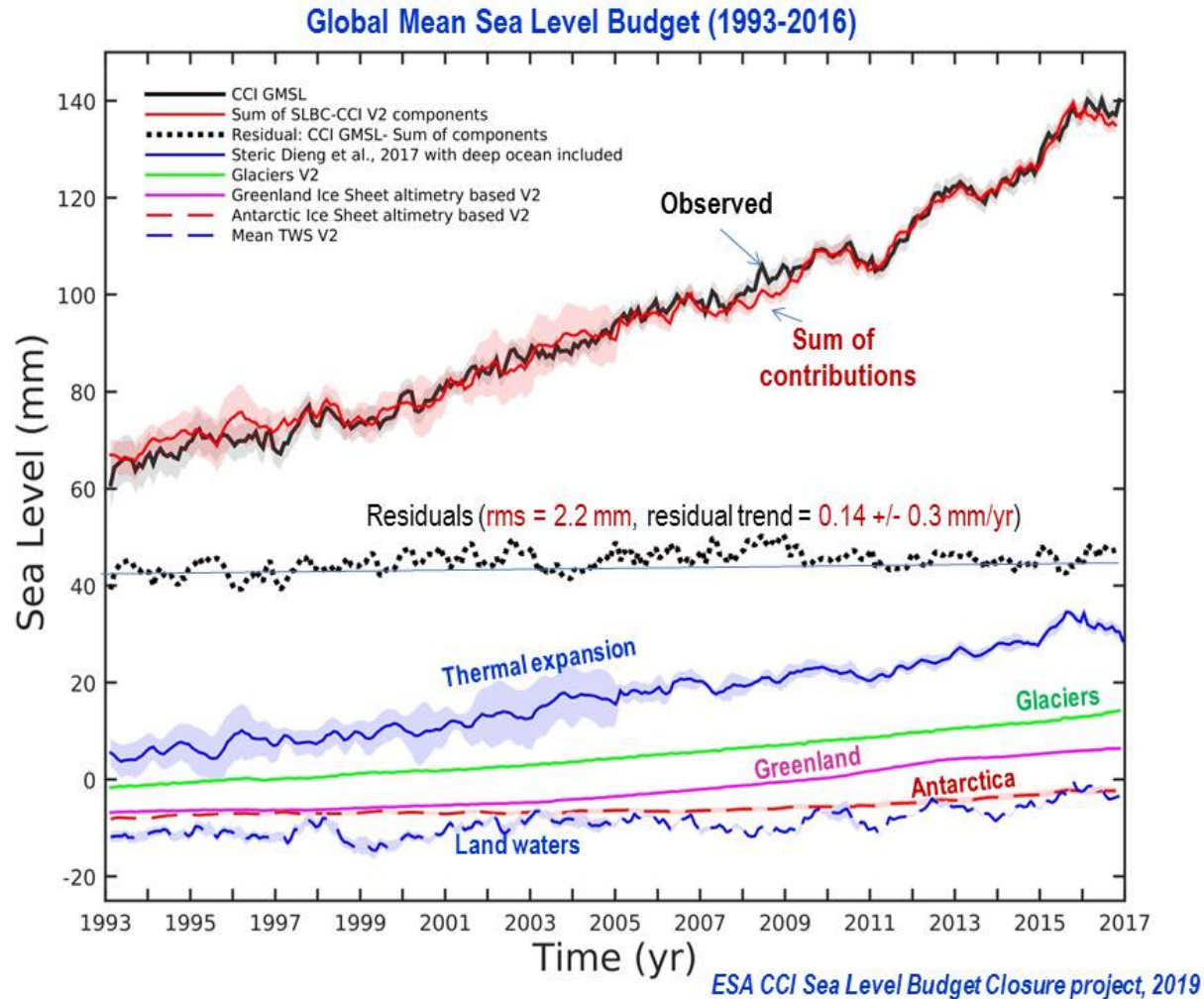
Uncertainty in sea level trends deduced from the var-covar matrix



Error Variance-covariance matrix over 1993-2018

Ablain et al., *ESSD*, 2019, Prandi et al, *Nature Scientific Data*, in revision

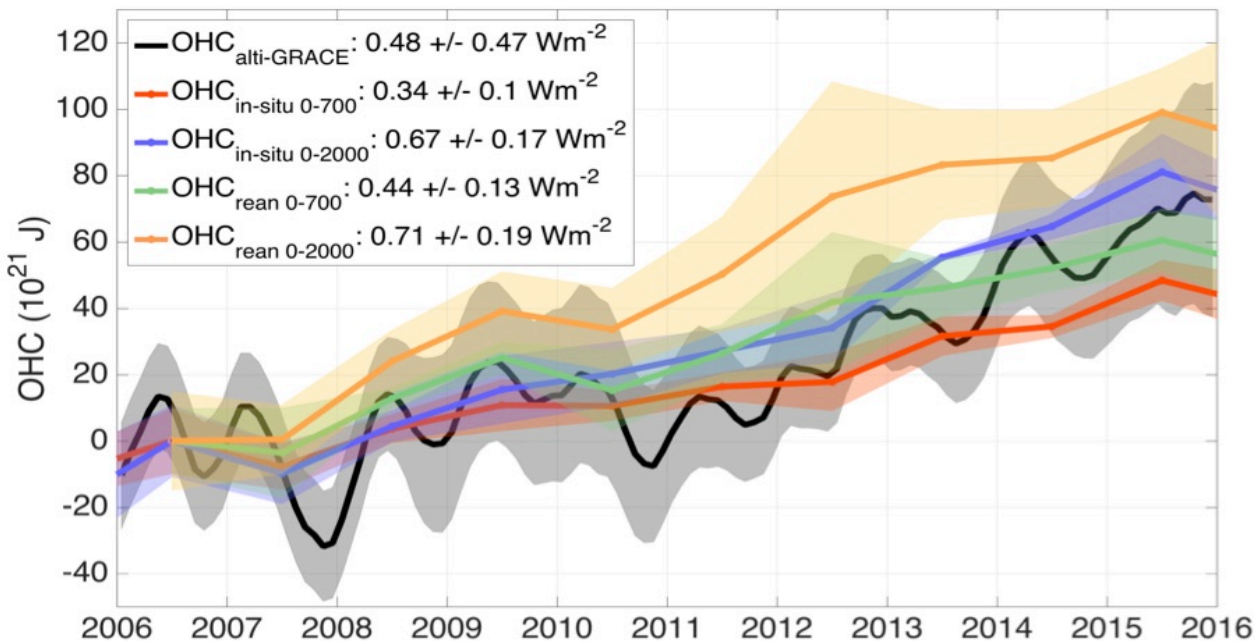
Closure of the Global Mean Sea Level Budget over the Altimetry Era



Dieng et al., *GRL*, 2017, Cazenave et al., *ASR*, 2018; The WCRP Global Mean Sea Level Team, *ESSD*, 2018; The ESA Sea Level Budget Closure project, Horwath et al., in preparation, 2020

Estimate of the Earth energy imbalance

Approach: estimate the global steric sea level from the difference between GMSL estimated by altimetry and ocean mass estimated by GRACE. Then estimate the global ocean heat content, which is a precise proxy of the Earth energy imbalance, by estimating the expansion efficiency of heat and multiplying it with the global steric sea level



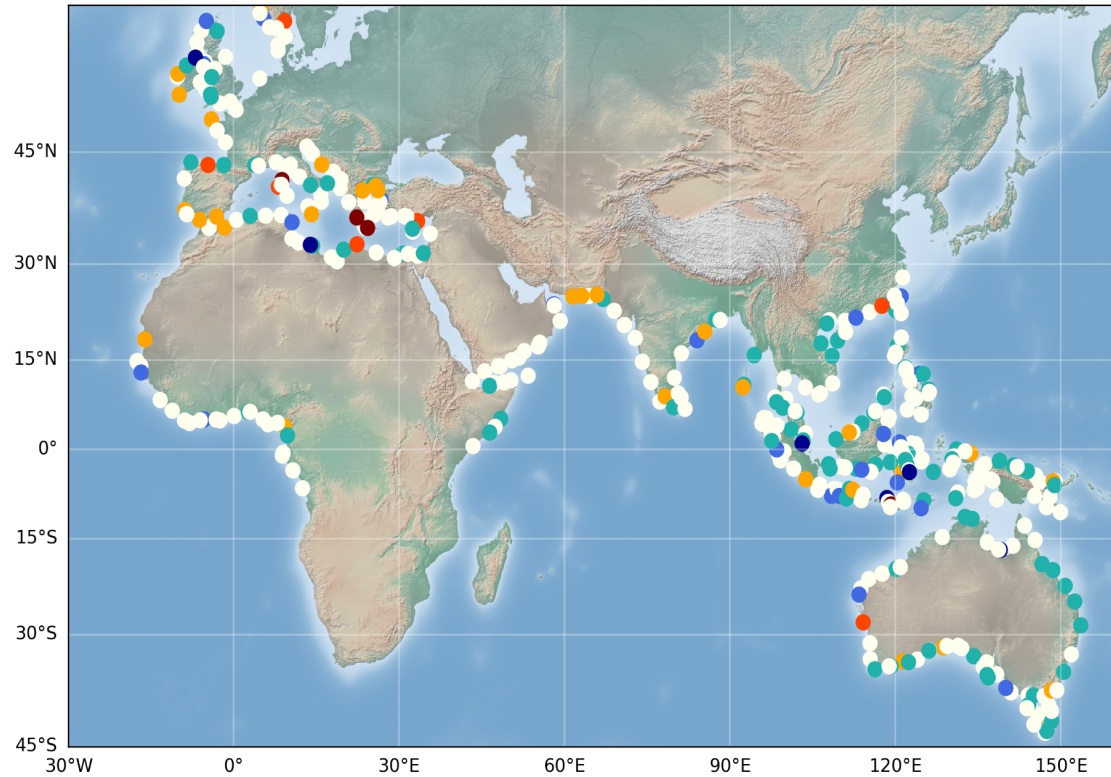
Global ocean heat content from alti-Grace, Argo data and ocean reanalysis

Ocean heat uptake	Time period	Spatial Coverage and/or depth range	mean in Wm ⁻²	Uncertainty in Wm ⁻² at the 5–95%CL	Correlation with CERES EBAF EEI	RMSE with CERES EBAF EEI Wm ⁻²
From <i>in situ</i> observations	2006–2015	0–2000m	0.61 (update of Johnson et al. 2018) ^a	±0.1 ^b	0.44	0.40
		deep ocean contribution 0–bottom	0.04 (update of Purkey & Johnson, 2010).	±0.04 (update of Purkey & Johnson, 2010);		
	1993–2017	0–2000m (no marginal seas, no ice covered areas)	0.65 ^c	±0.11 ^d		
		deep ocean contribution (below 2000m, no ice covered areas) 0–bottom	0.62 (update of Johnson et al. 2018) ^a 0.04 (update of Purkey & Johnson, 2010); 0.66 ^c	±0.22 ^e ±0.04 (update of Purkey & Johnson, 2010); ±0.22 ^d		
From surface net heat flux	2006–2015	Net ocean surface heat flux	10 to 15 ^f	±15 ^g		
From satellite altimetry and GRACE	2006–2015	0–bottom (no sea ice covered areas above 82°N)	0.53 ^h	±0.38 ⁱ	0.89	0.26
	2002–2016	0–bottom (no sea ice covered areas above 82°N)	0.57 ^j	±0.29 ^k		
From ocean reanalyses	2006–2015	0–2000m	0.7 (update von Shuckmann et al. 2008) ^l	±0.13 (update von Shuckmann et al. 2008) ^m	0.50	0.41
		deep ocean contribution (below 2000m, no ice covered areas) 0–bottom	0.04 (update of Purkey & Johnson, 2010). 0.74 ^c	±0.04 (update of Purkey & Johnson, 2010); ±0.14 ^d		
	1993–2008	0–bottom	0.71 (from Palmer et al. 2017)	±0.7 (spread across 15 ocean reanalyses from Palmer et al. 2017)		
From CMIP5 climate model simulations	2000–2010	0–bottom	0.73 (from Smith et al. 2015)	±0.21 (spread across 21 CMIP5 climate model simulations, from Smith et al. 2015)		

Uncertainty in Global ocean heat content and EEI from alti-Grace, Argo data and ocean reanalysis

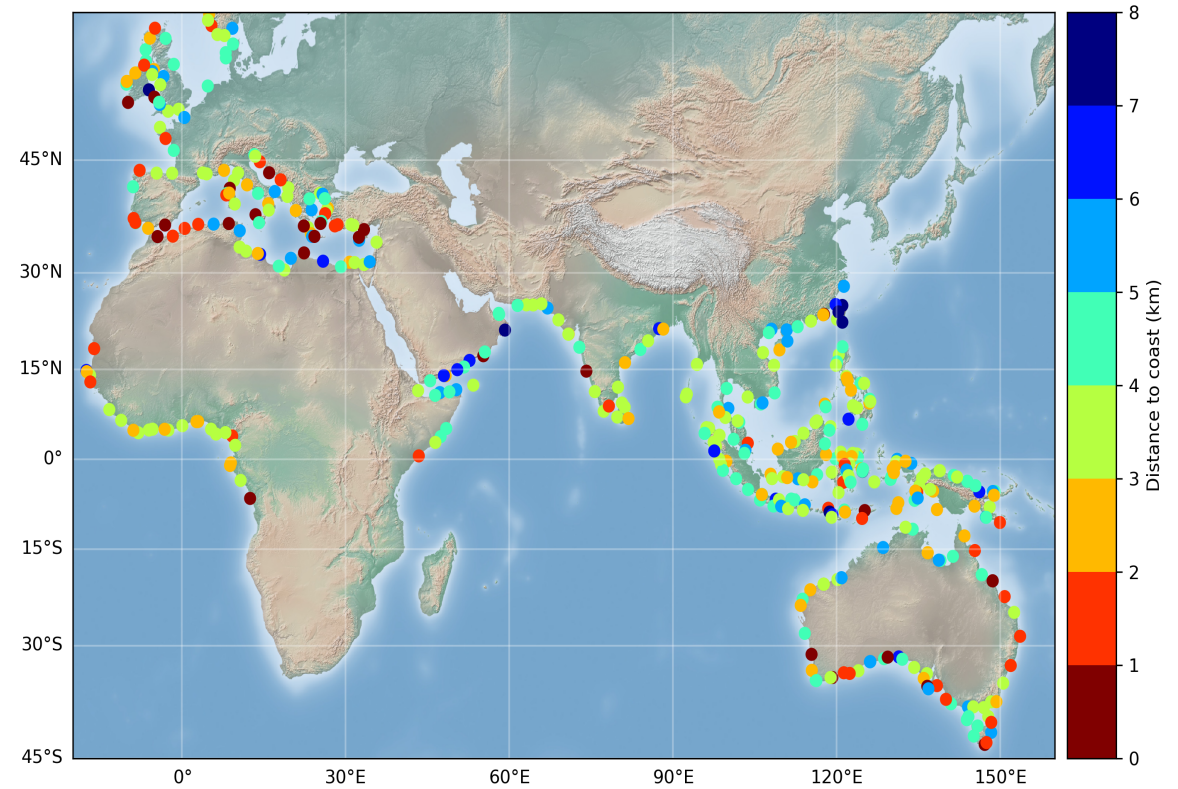
Coastal Sea Level Changes from Reprocessed Satellite Altimetry

Approach: retracking of Jason-1, 2, 3, Envisat and Saral/AltiKa altimetry missions to estimate sea level trends in the world coastal zones (context: ESA climate change Initiative coastal sea Level project)



Difference in sea level trends between open ocean and coast (mm/yr) (2002-2018)

- In 20% of the sites, the coastal trend significantly differs from open ocean trend
- Coastal processes (e.g., T/S changes, currents, waves, fresh water input from rivers, etc.) are under investigation to explain this observation



Closest distance to coast (km) where reliable sea level trends can be estimated

Marti et al., ASR, 2019; Gouzenes et al., Ocean Sciences, 2020; The CCI Coastal Sea Level Team, Nature Scientific Data, 2020