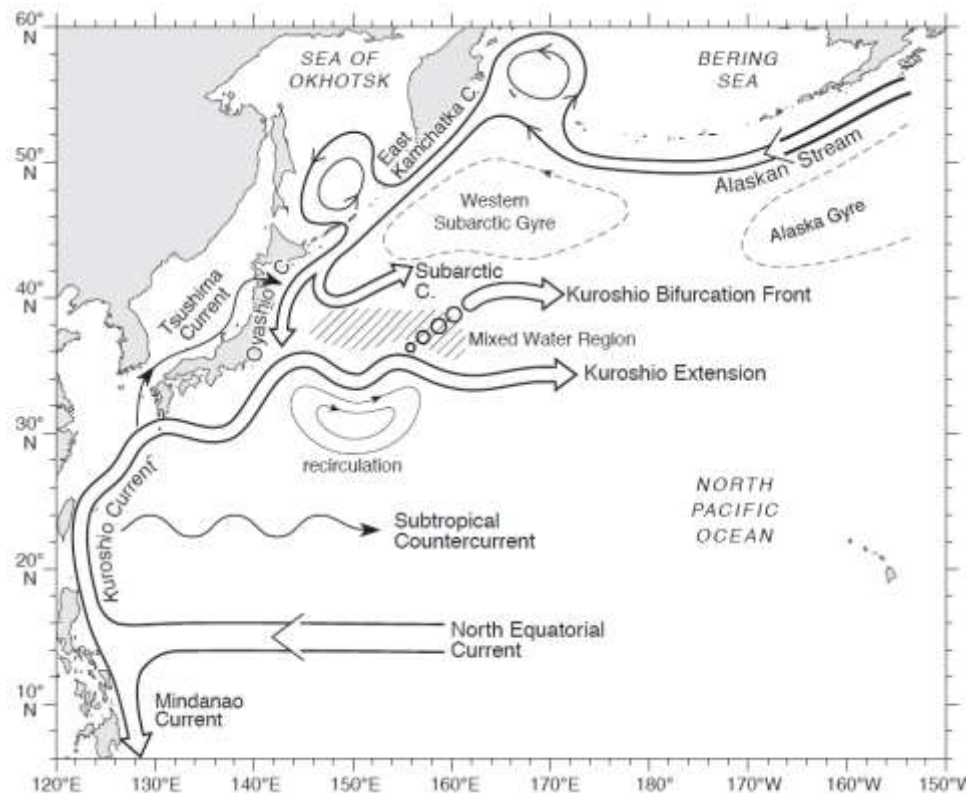


# Interplay of Intrinsic versus Forced Decadal Kuroshio Extension Variability: Observations & Predictability

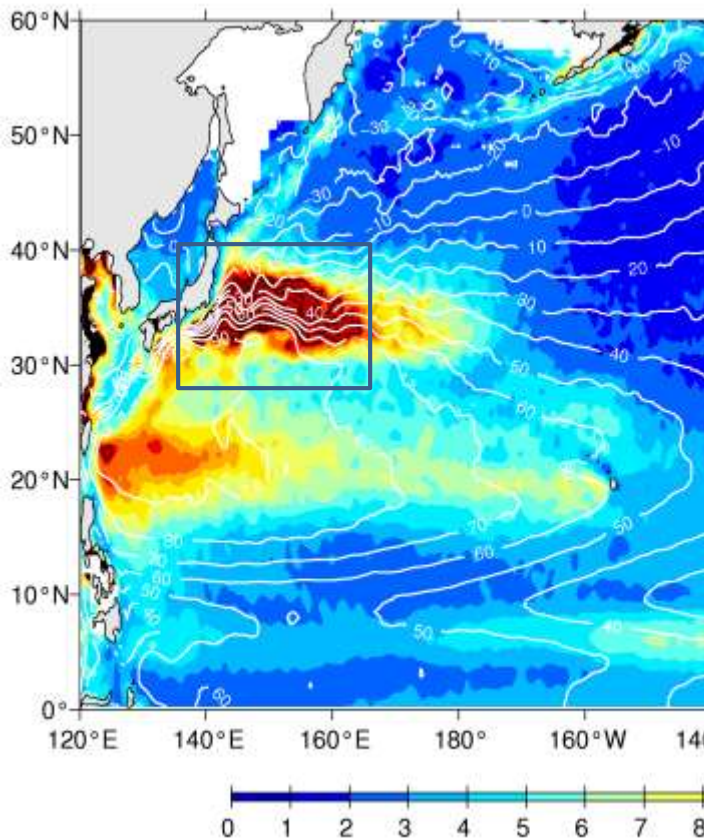
Bo Qiu, Shuiming Chen & Niklas Schneider

Department of Oceanography, University of Hawaii at Manoa

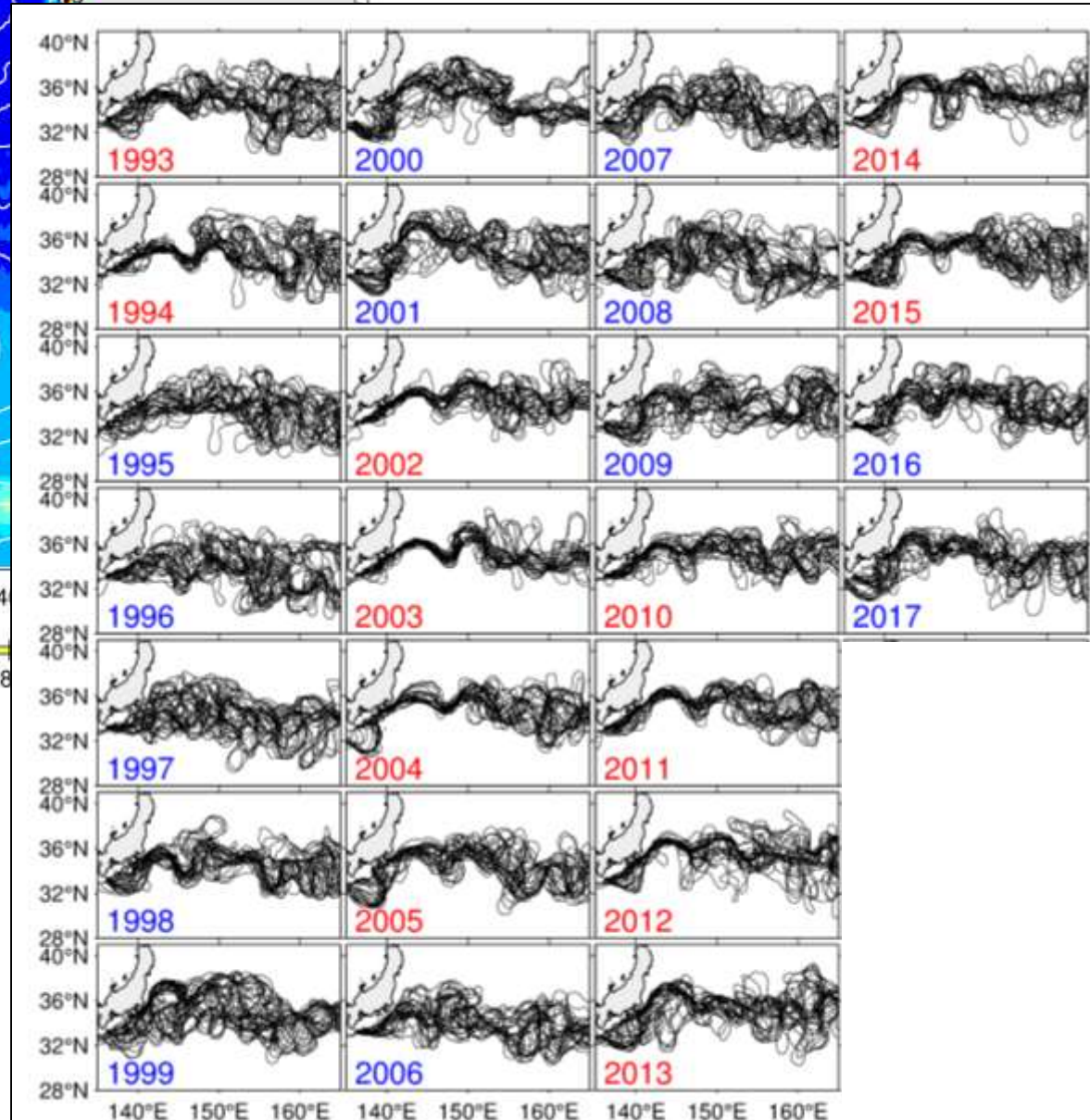


Schematic of NW Pacific Ocean Circulation  
(Qiu 2019, Encyc. Oce. Sci.)

# Annual maps of bi-weekly paths of the Kuroshio/KE jet

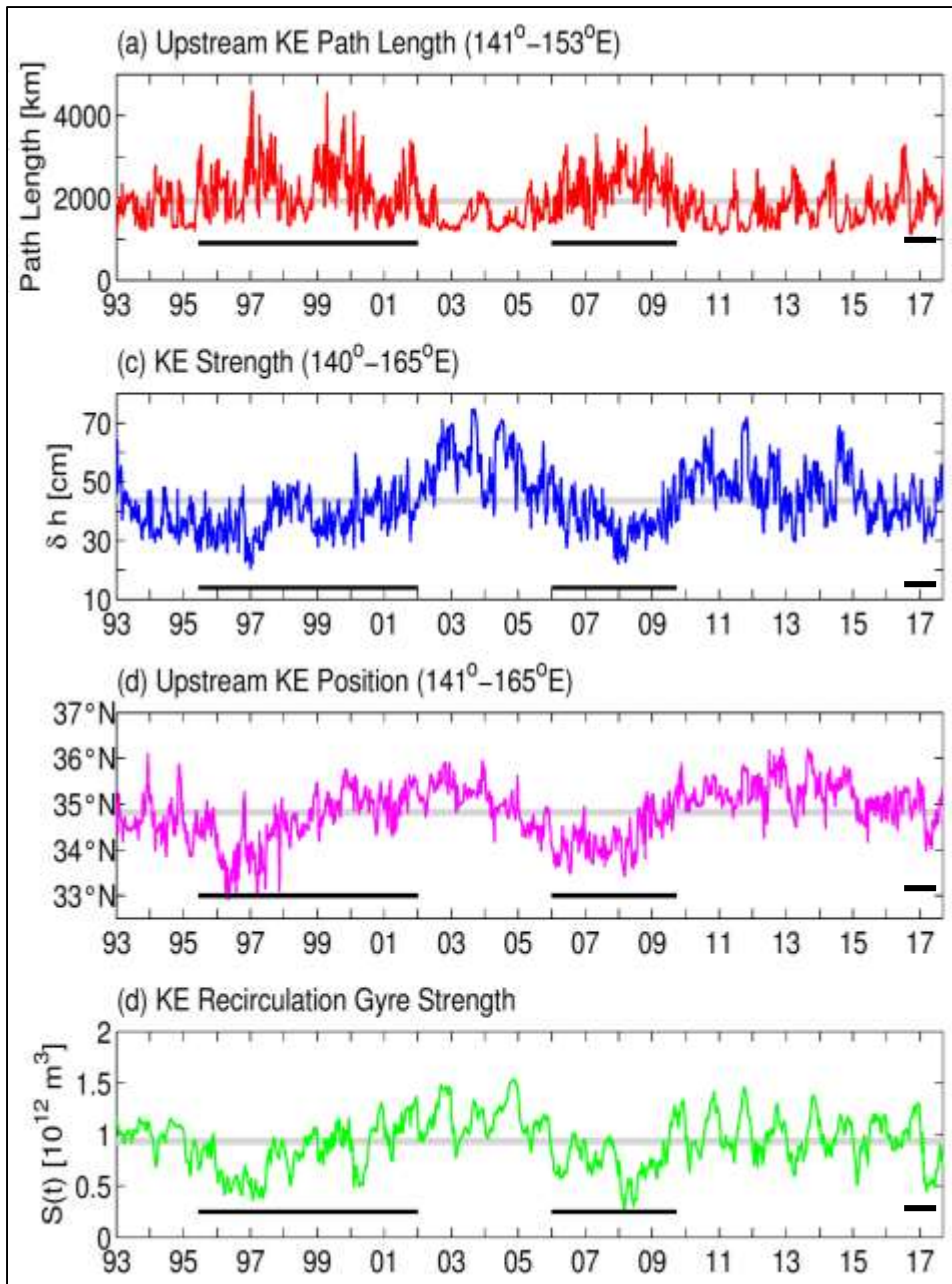
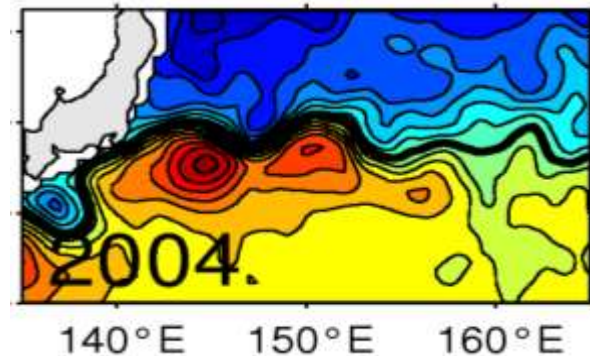
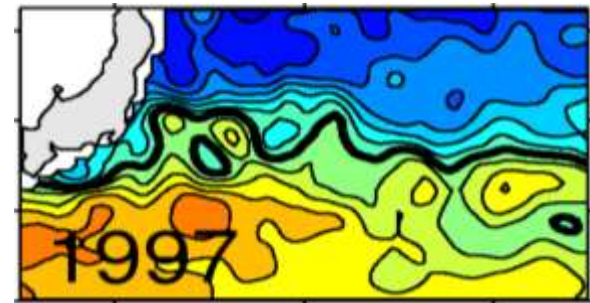


Alternation between **unstable**  
versus **stable** states



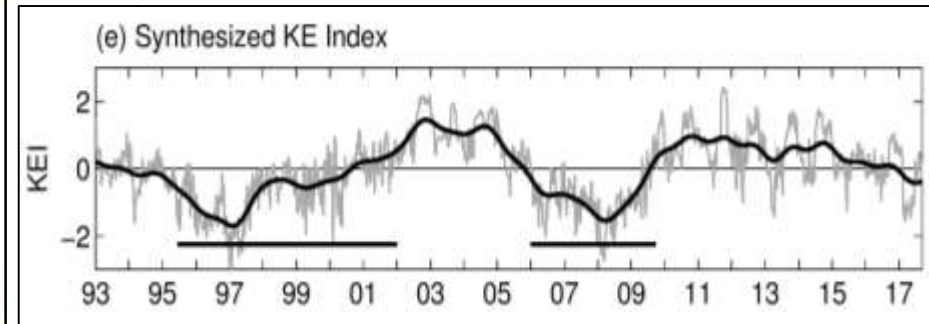
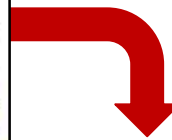
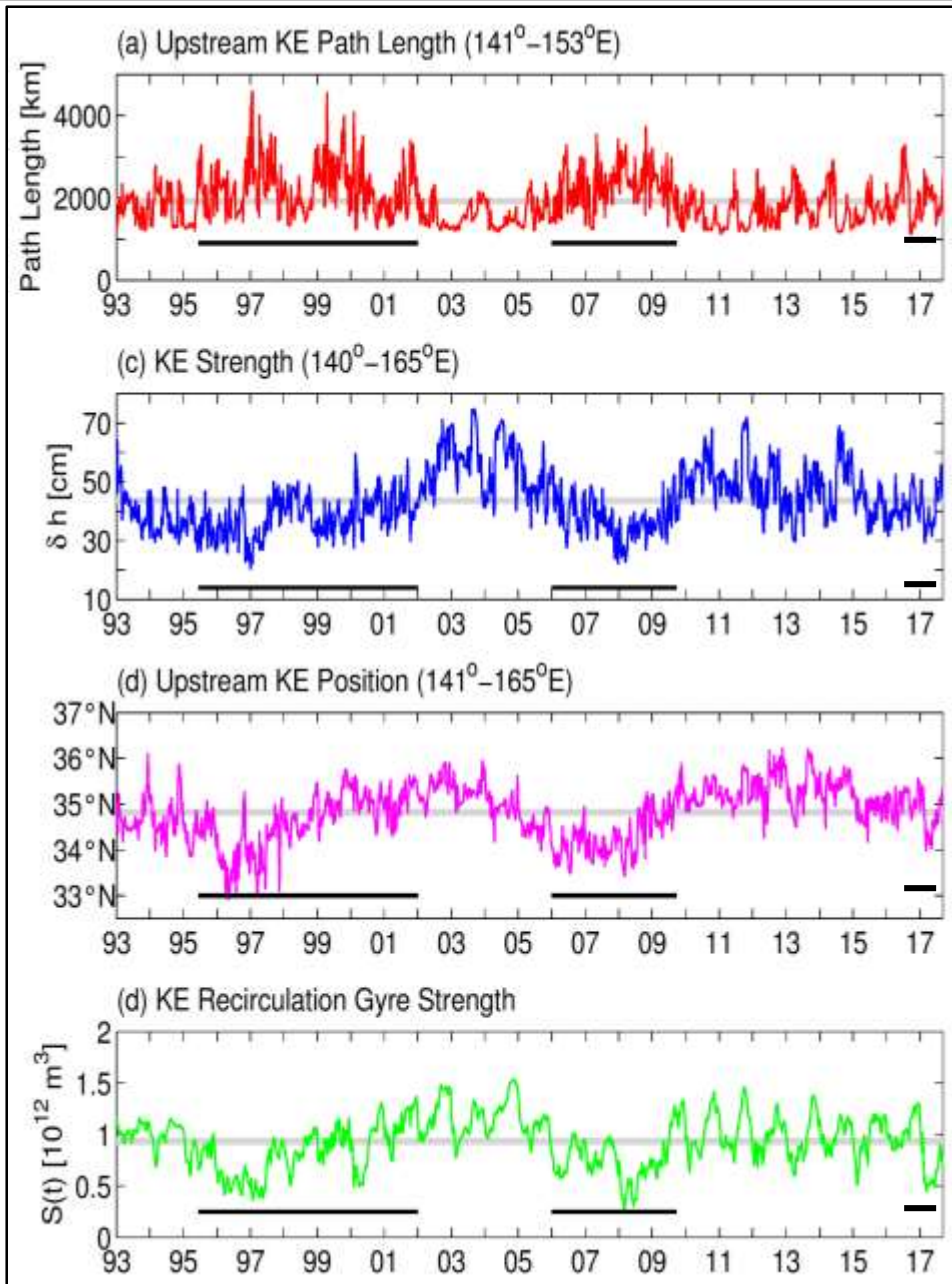
# Dynamically **unstable** vs. **stable** states of the KE system

Typical yearly SSH patterns in  
**unstable** vs. **stable** dynamic states



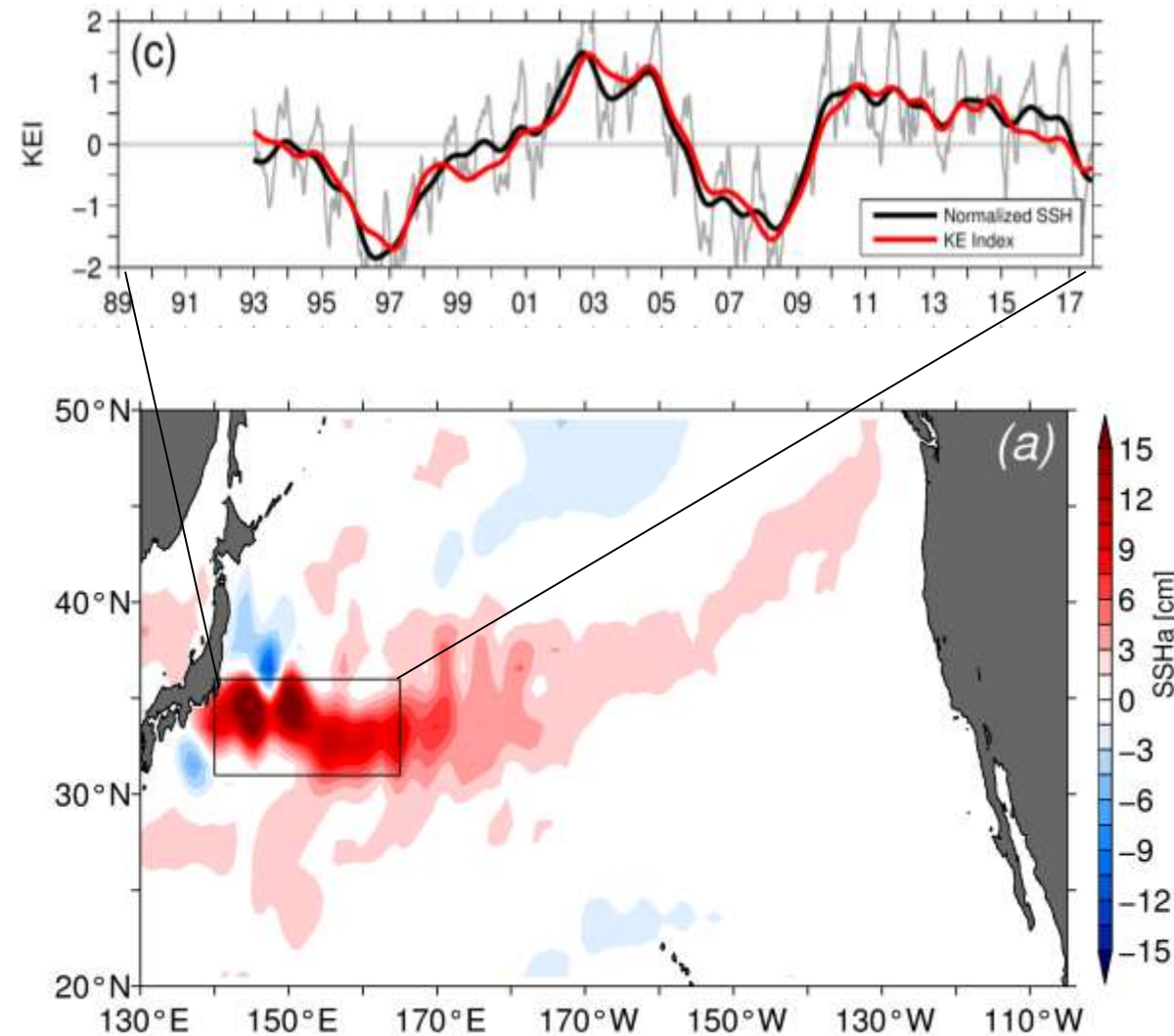


# Form a **comprehensive index** representing the KE variability



**KE index** : average of these 4 dynamic quantities (normalized)

# Regression between the **KE index** and basin-wide SSH anomaly field



**KE index:** represented well by SSH anomalies in the **southern RG box** (31-36°N, 140-165°E)

- KE index from dynamic properties
- box-averaged SSH time series

## Implications

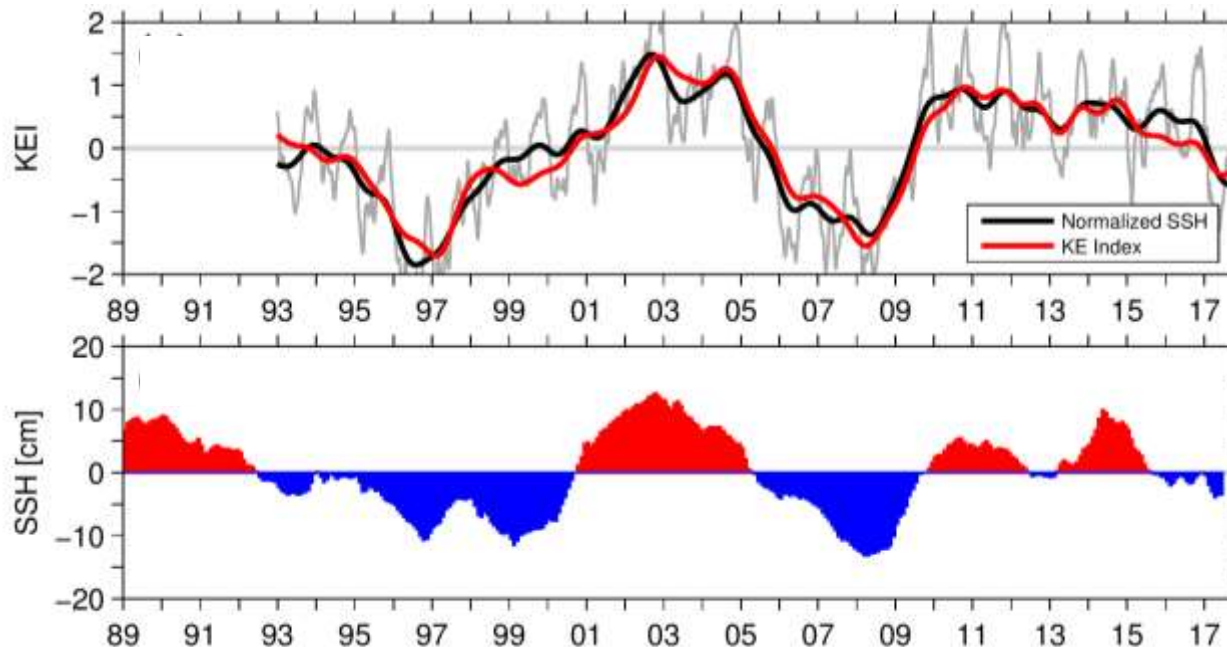
- Examining **KE index** becomes equivalent to examining SSH anomalies in this **key box**

# Quantifying the **KE index** using a wind-forced linear vorticity model

- The proxy KE index is governed by wind forcing along the 31°-36°N band across the North Pacific basin:

$$\frac{\partial h}{\partial t} - c_R \frac{\partial h}{\partial x} = -\frac{1}{\rho} \nabla \times \left( \frac{\tau_{wind}}{f} \right)$$

- Hindcast of KE index using **ECMWF interim Ekman pumping** data along Rossby wave characteristics:



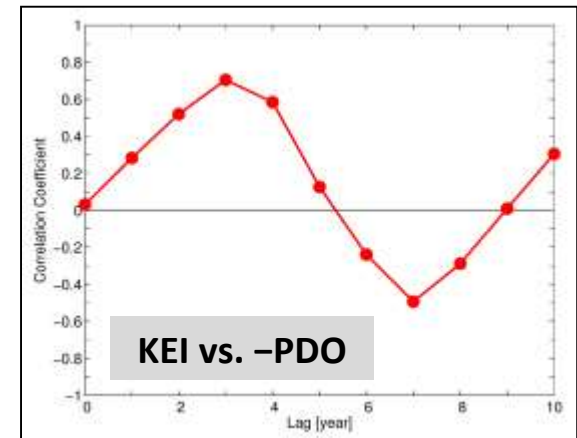
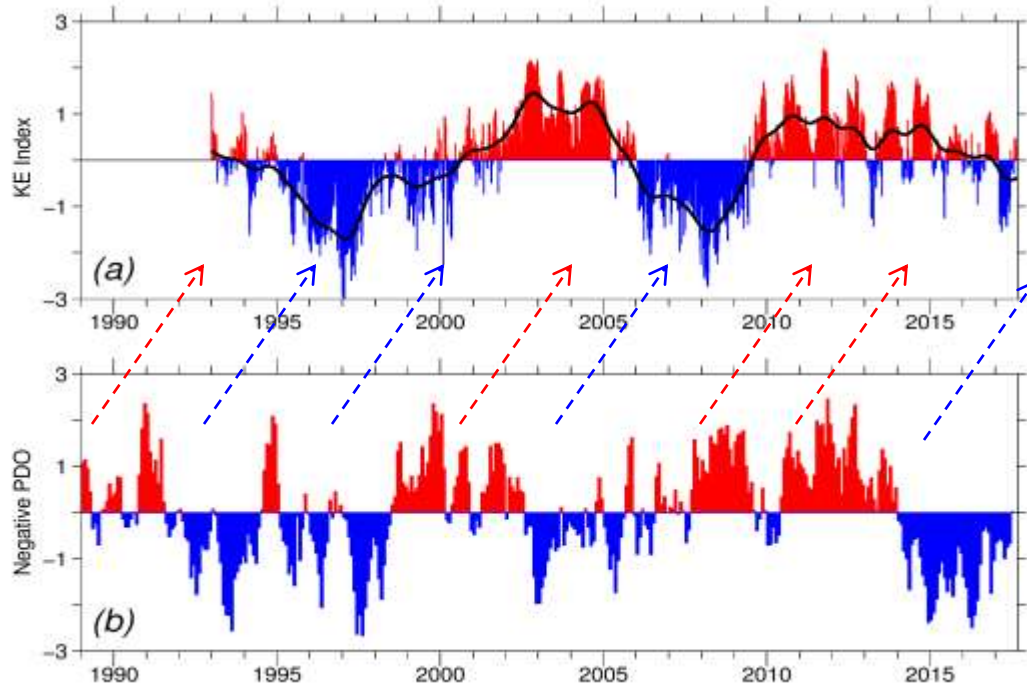
original time series:

$r = 0.65$

low-pass filtered TS:

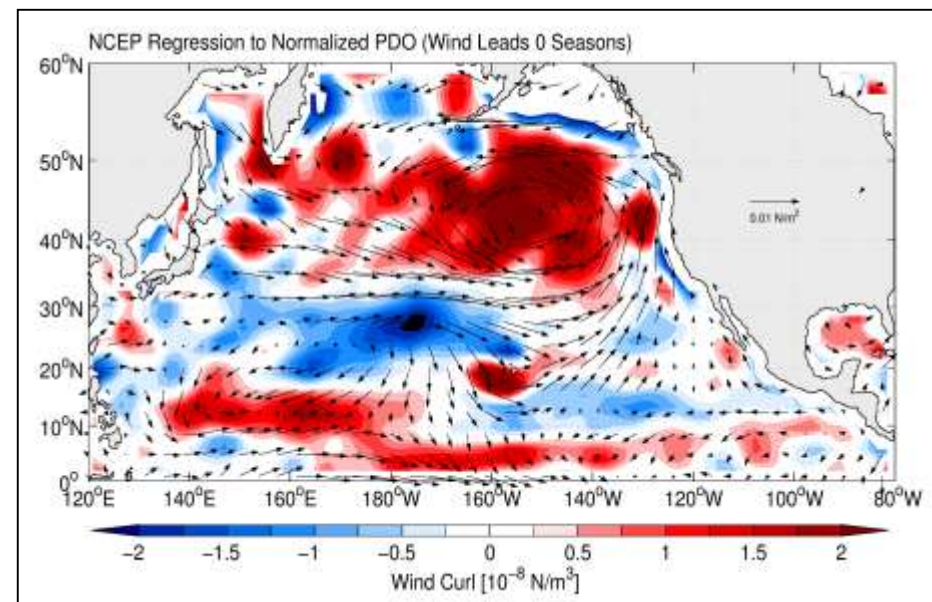
$r = 0.87$

# Decadal KE variability lags the **negative PDO** index by $\sim 3$ years ( $r = 0.74$ )



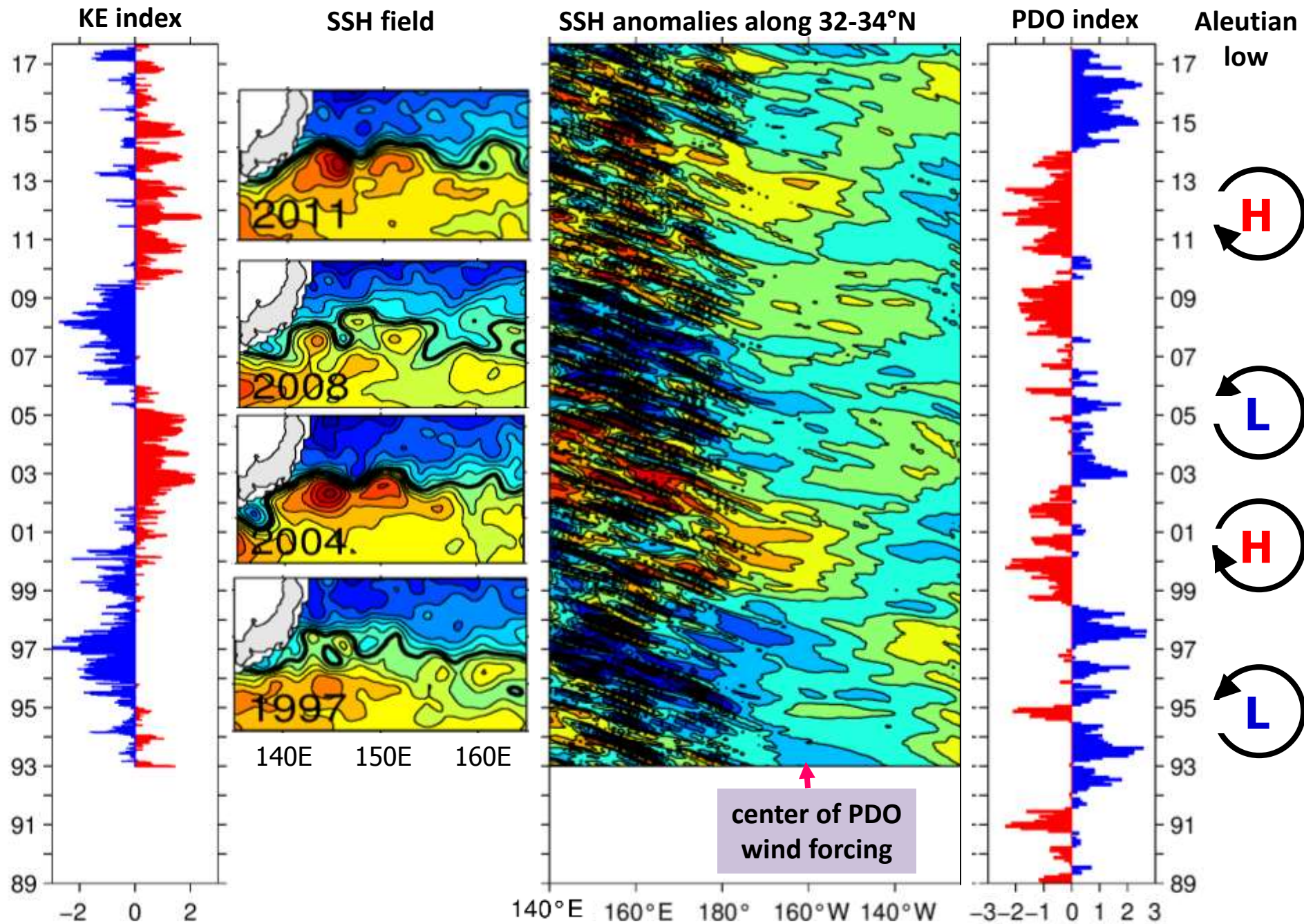
## PDO-regressed wind stress vector & curl (color)

- Center of PDO wind forcing is in **eastern** half of N Pacific basin



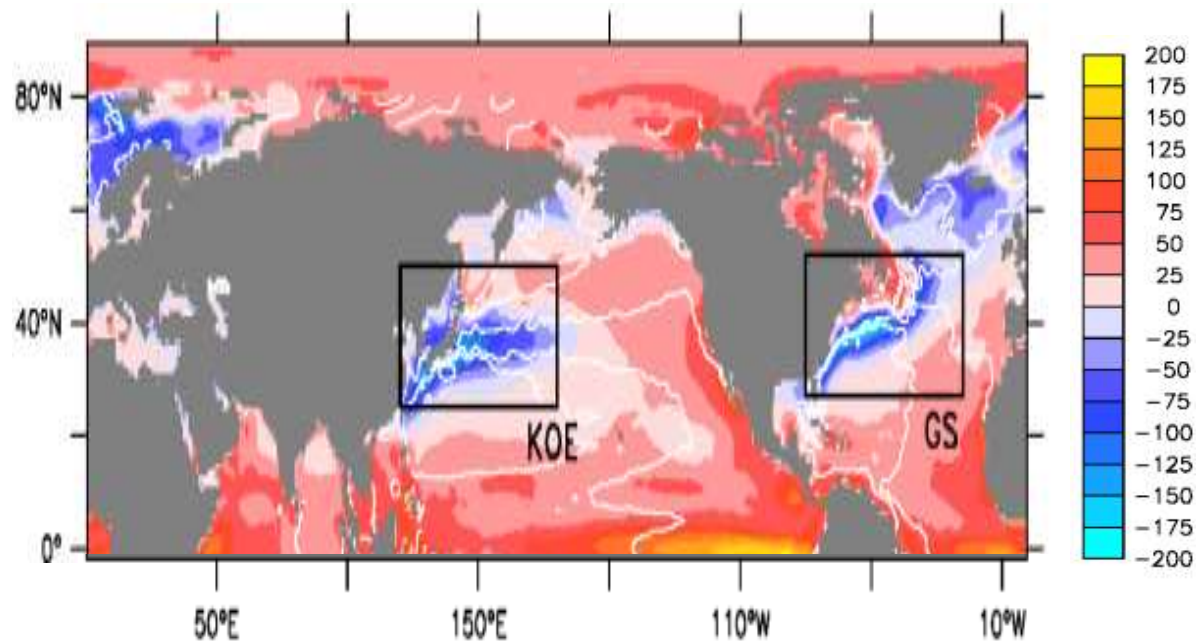


# Connections between PDO forcing, cross-basin SSH adjustment & KE index

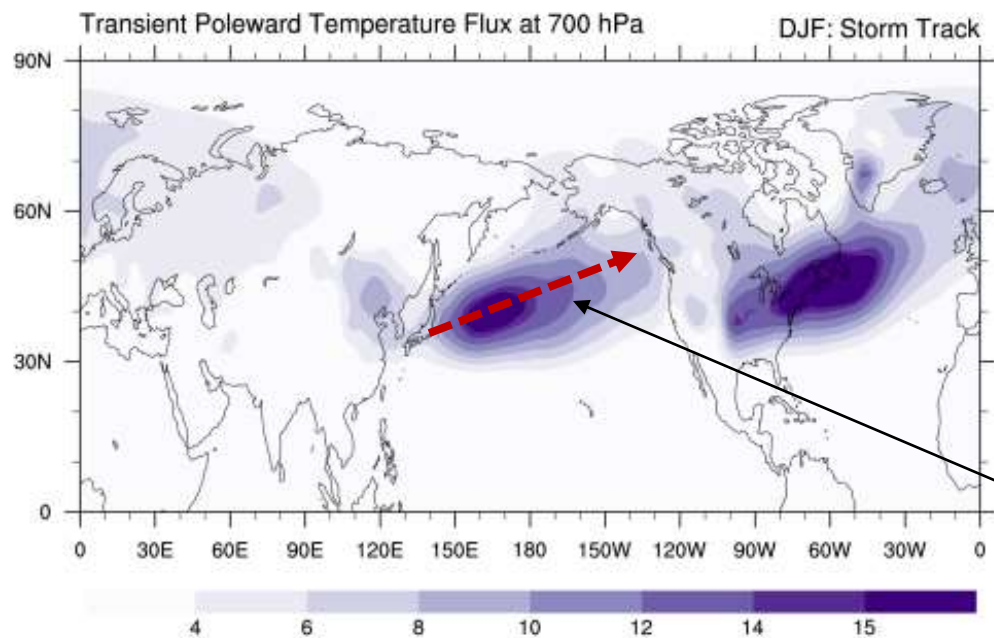




# Kuroshio/KE supply heat to overlying atmosphere & anchor stormtracks



Time-mean net surface heat  
flux from air to ocean  
(W/m<sup>2</sup> ; Cronin et al. 2009)

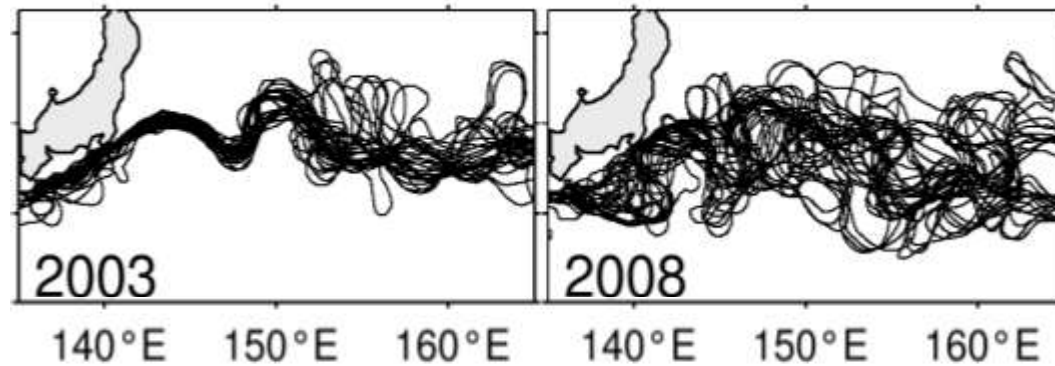


DJF 700mb rms  $v'T'$  fluxes  
indicative of storm-track variability  
(Nakamura et al. 2004)

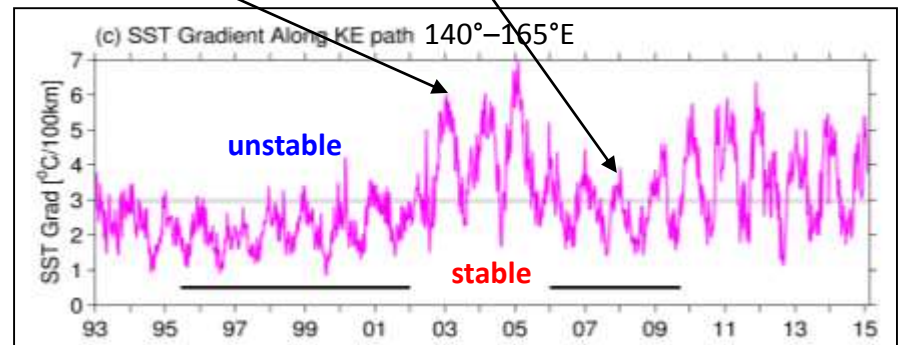
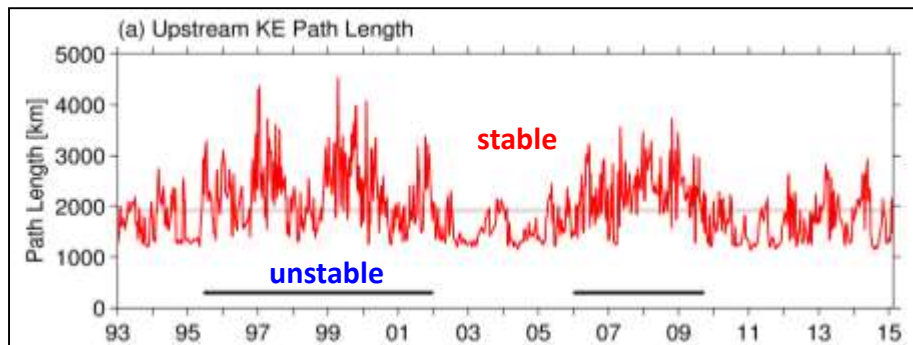
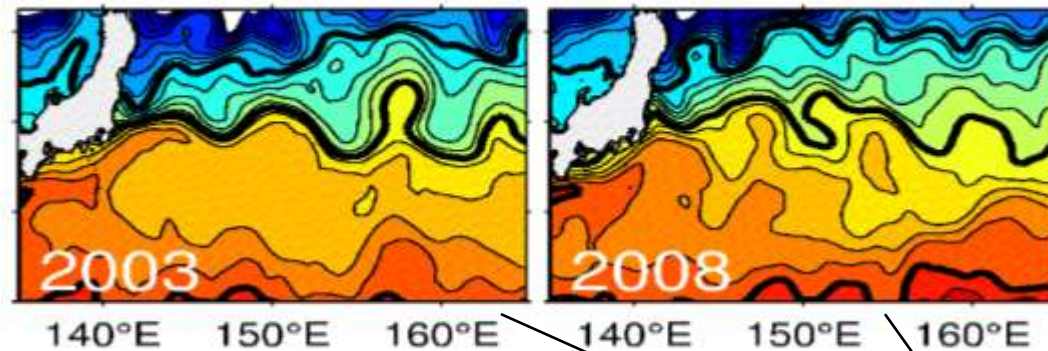
$\text{curl } \mathbf{x} \boldsymbol{\tau} = 0$  line/storm-tracks

# KE dynamic state affects regional SST & cross-front SST gradient

Bi-weekly  
KE paths

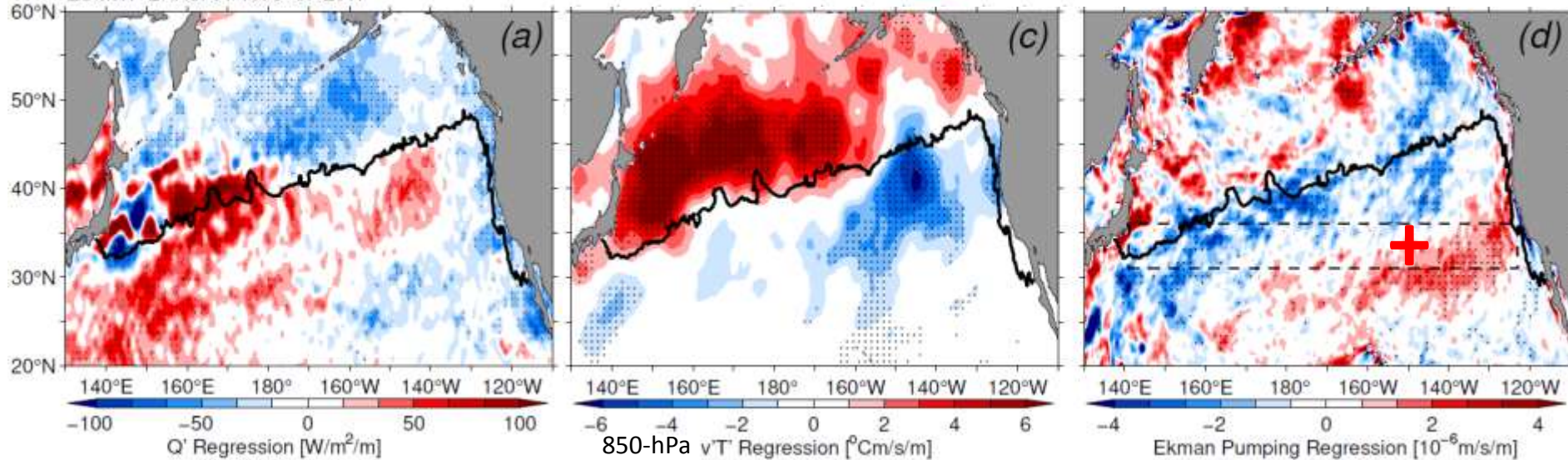


Feb-March  
SST maps

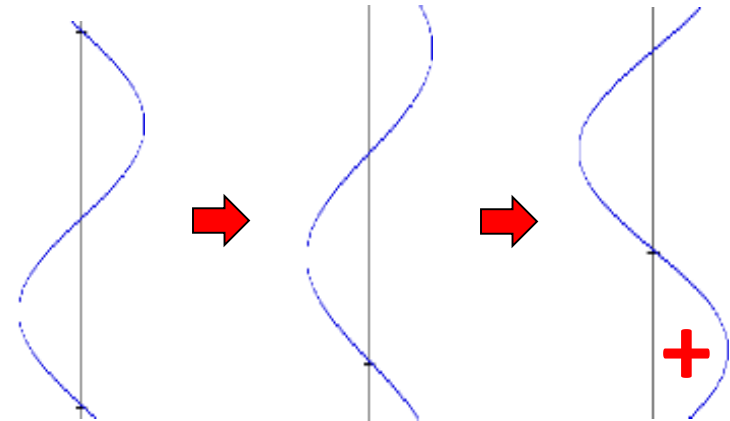


# Turbulent heat flux, storm-tracks & wind curl regressed to the KE index

ECMWF ERA5: 01/1979–07/2017



- A stable KE (warmer SST & stronger KE front) enhances upward turbulent heat flux in downstream KE
- This induces a poleward shift in storm-tracks as inferred from 850-hPa  $\langle v'T' \rangle$  fluxes
- Surface wind stress pattern moves similarly poleward → **positive curl anomaly in eastern basin along 31°–36°N band**



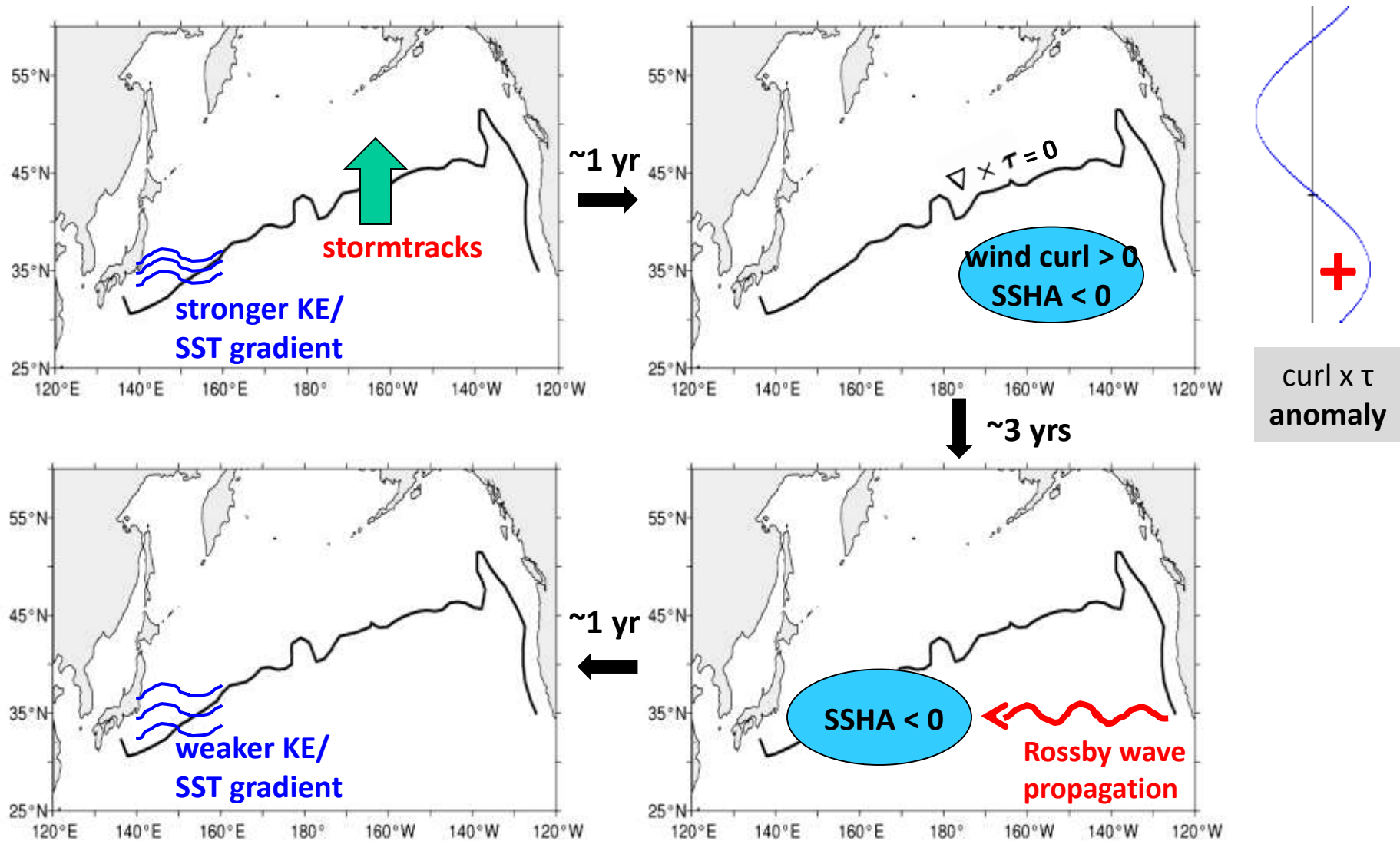
mean curl  $\times \tau$   
pattern

poleward-  
shifted curl  $\times \tau$

curl  $\times \tau$   
anomaly



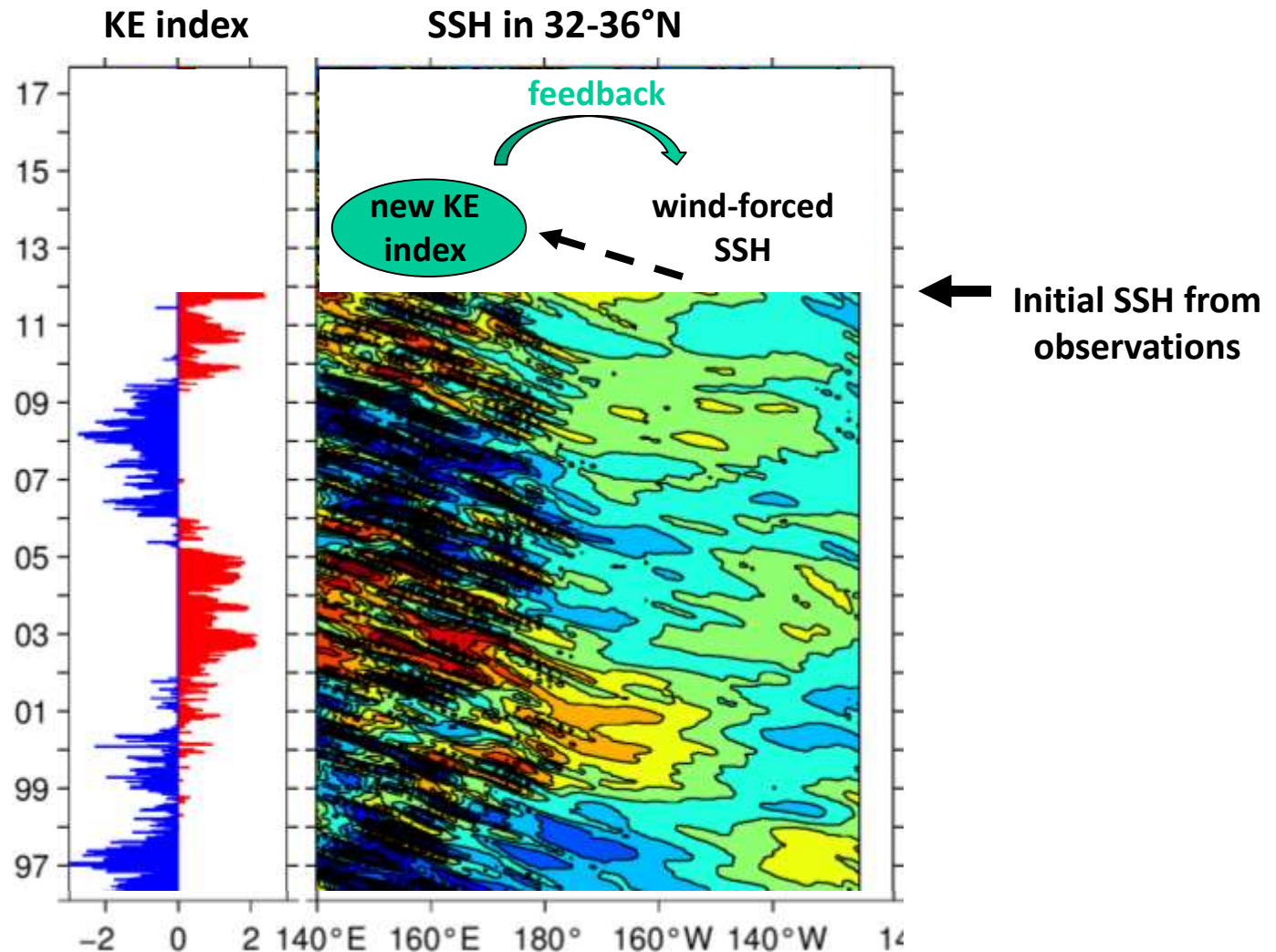
# KE interaction with stormtracks favors a **delayed negative feedback loop**



Half a cycle = ~ 5 yrs with Pacific basin crossing + adjustment times

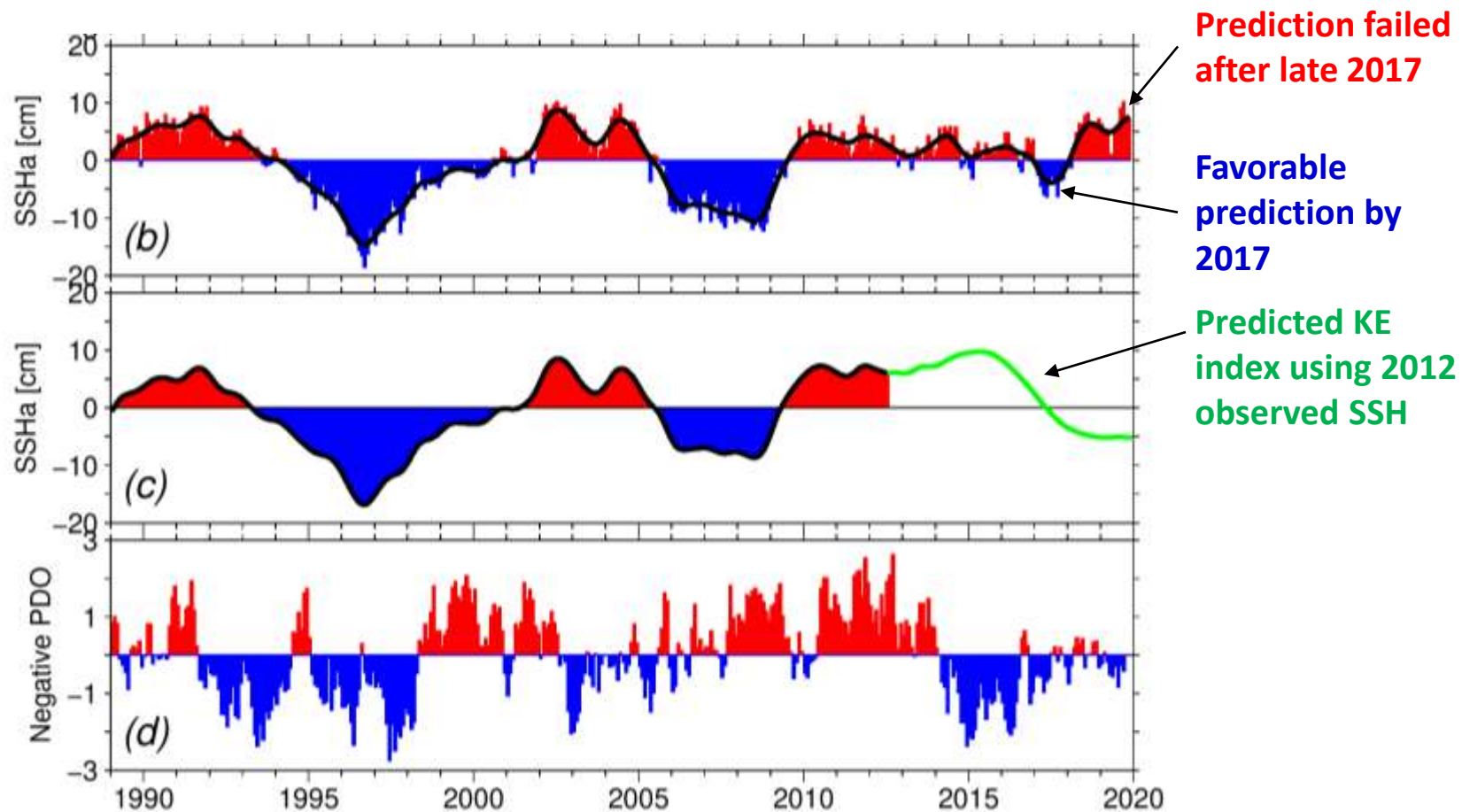
# Predicting **KE index** based on the negative feedback mechanism

- Prediction to **KE index** can be made by taking into account the SSH signals propagating from the east, plus the SSH signals generated by the KE-induced wind forcing across the Pacific basin:



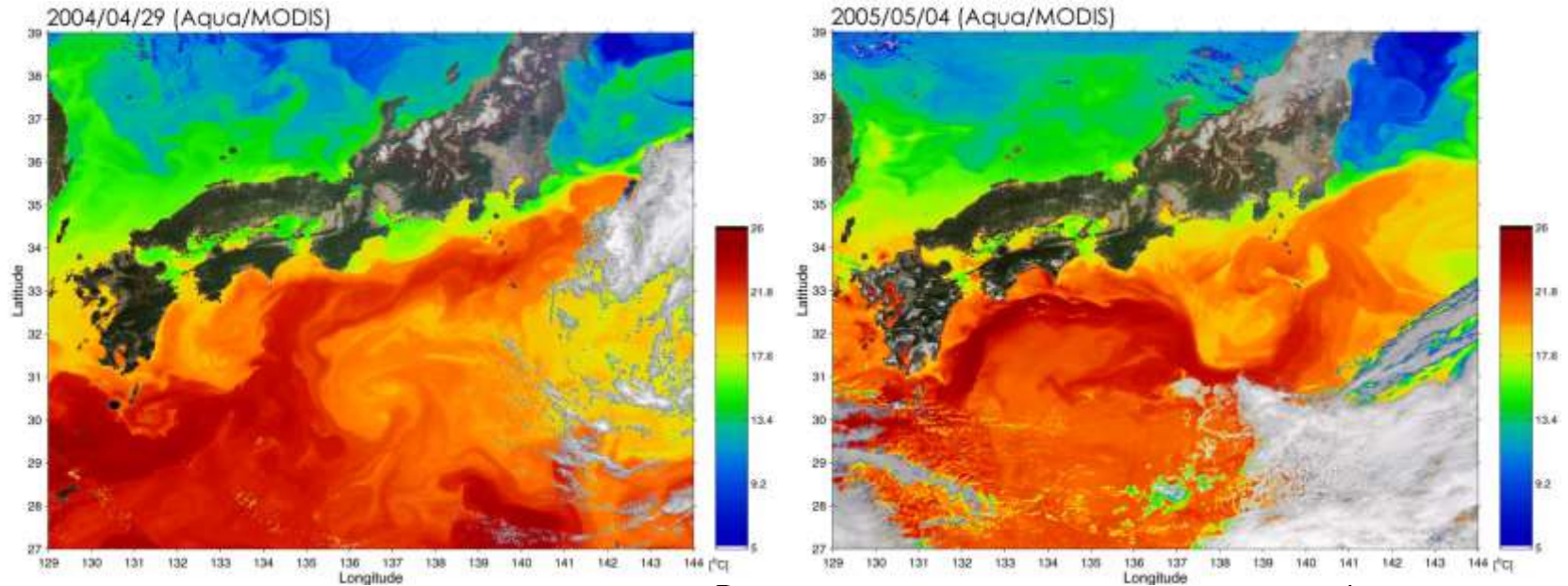
# Verification of the predicted KE index based on 2012 AVISO SSH data

AVISO  
data

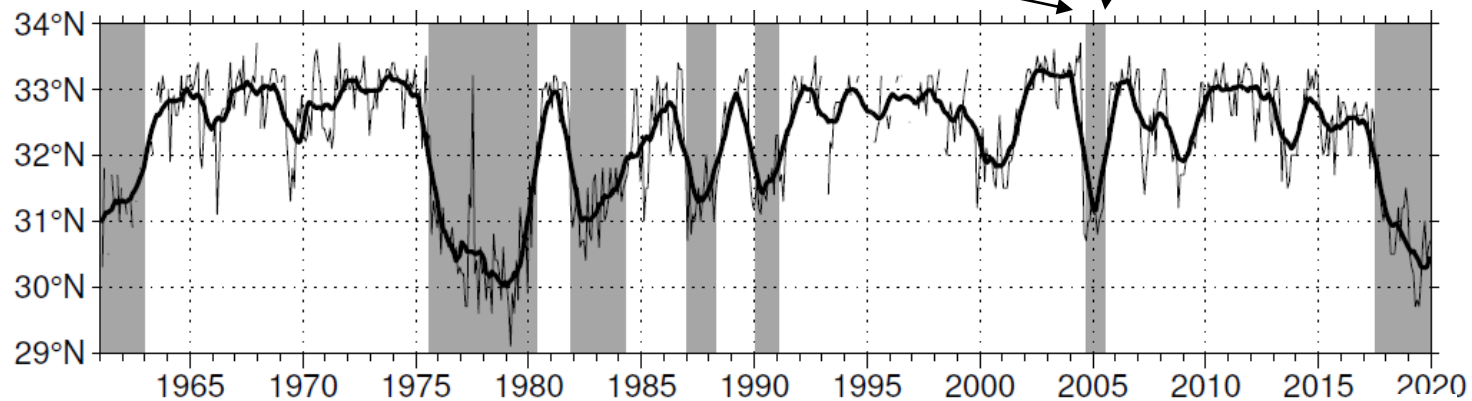




# Bimodal paths of the Kuroshio south of Japan



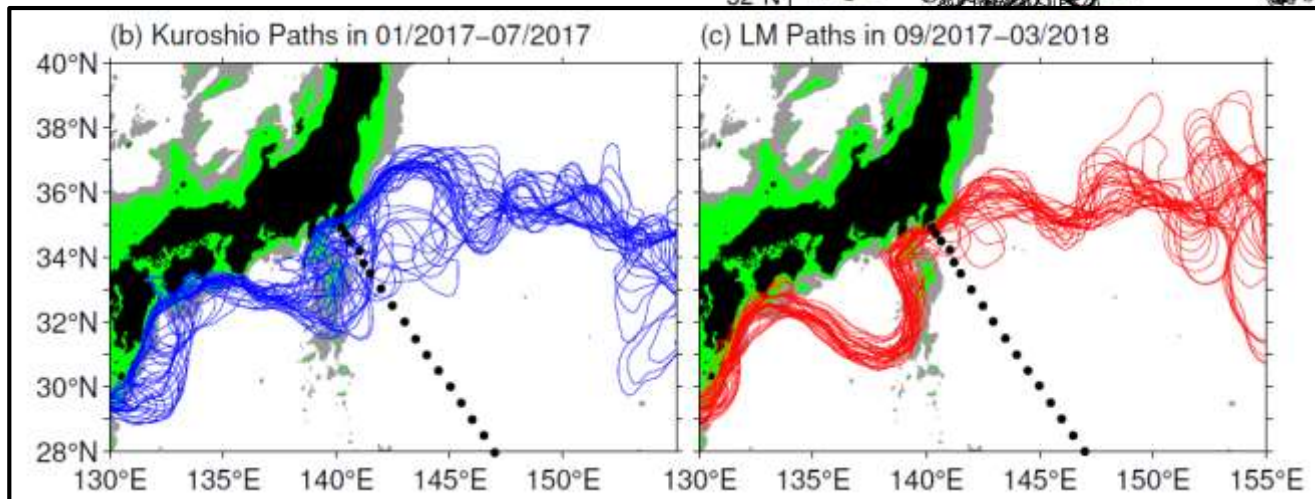
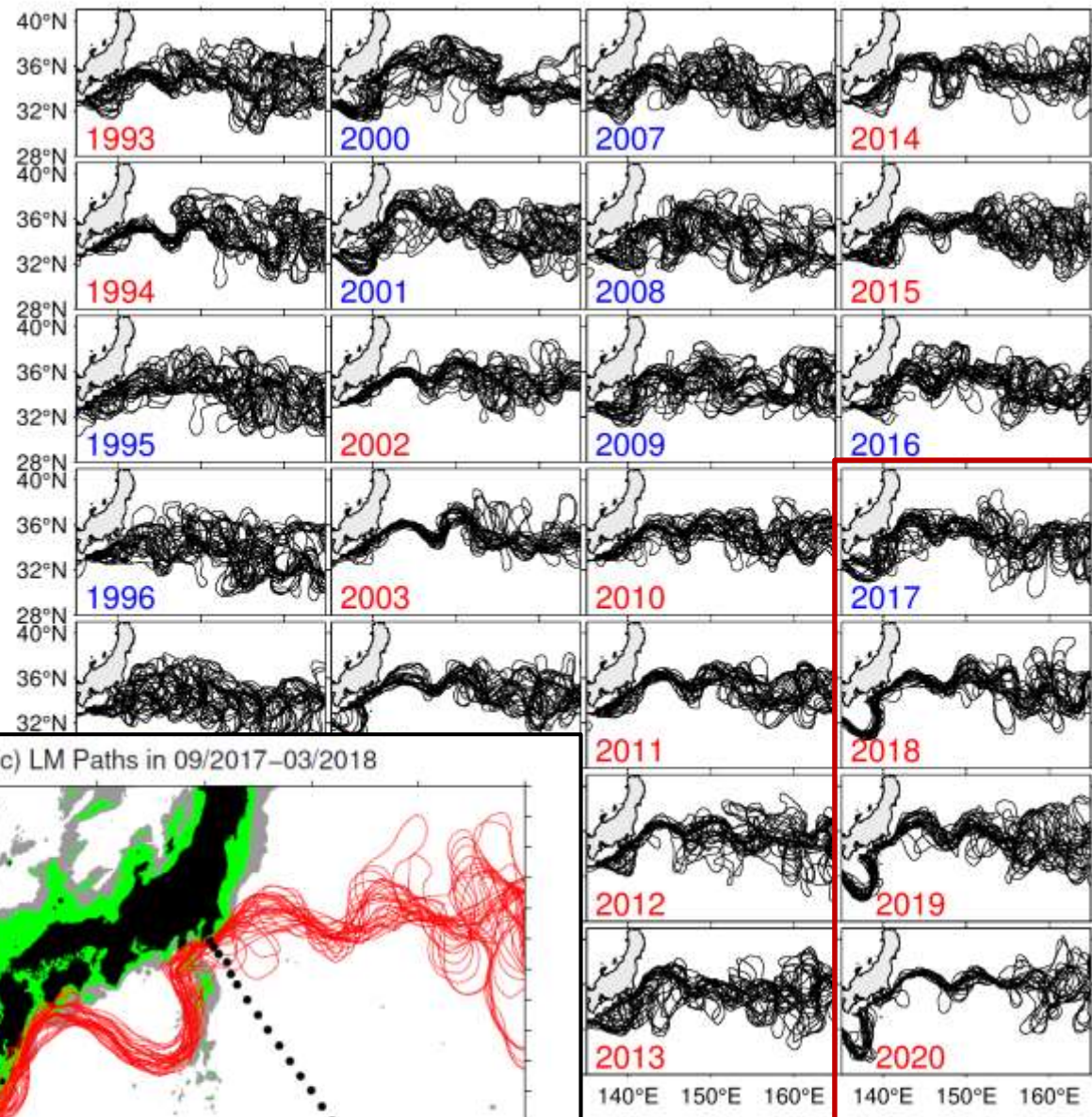
Southernmost Kuroshio latitude from Japan coast (JMA)



- Irregular path fluctuations on interannual & longer timescales

# Annual maps of bi-weekly paths of the Kuroshio/KE jet

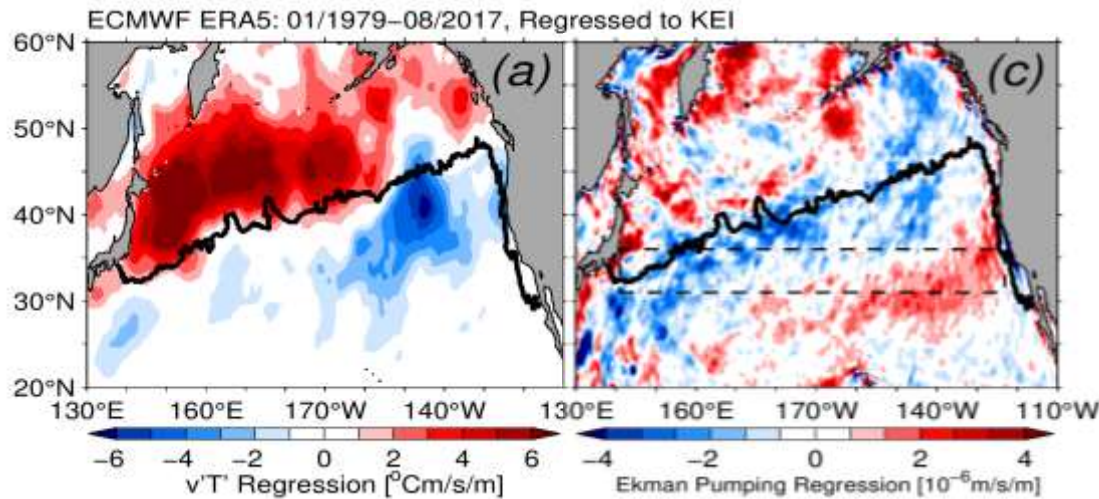
- Occurrence of the Kuroshio LM in August 2017 forced the KE path poleward, avoided the shoaling Izu Ridge & switched KE's dynamic state back to a **stable state**





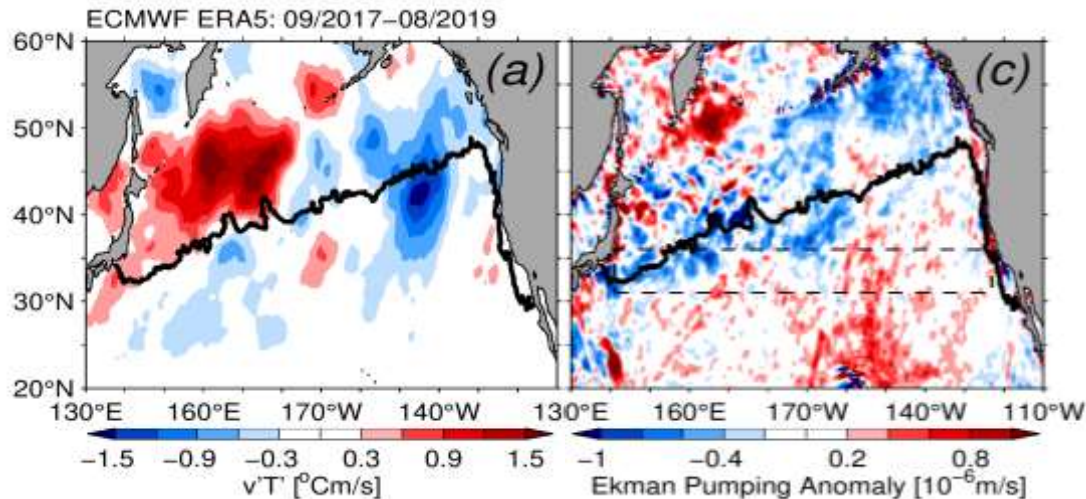
# Many questions remain to be addressed & answered

Do the **LM-induced KE change** after 2017 alter the basin-wide atmospheric circulation in a similar way as the **wind-forced KE change** ?



Storm-tracks & Ekman pumping variations regressed to the KE index of 01/1979–08/2017

wind-forced KE



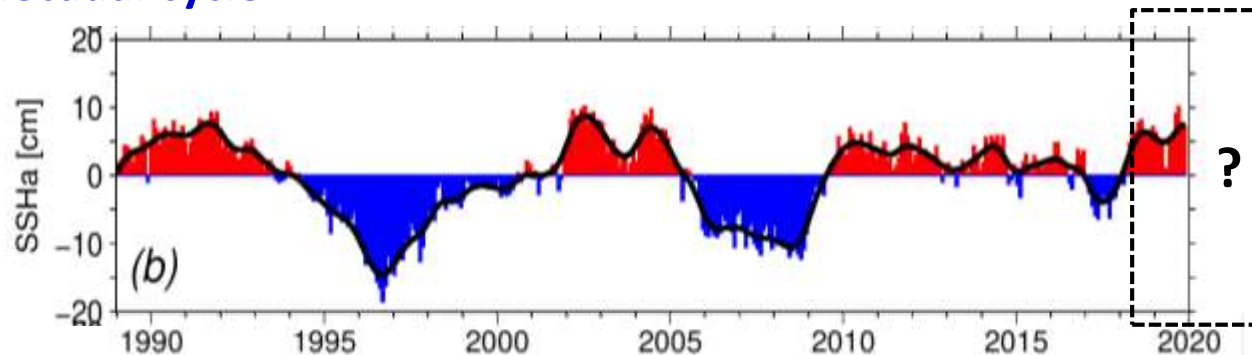
Storm-tracks & Ekman pumping anomalies in 09/2017–08/2019

LM-forced KE



# Summary

- The KE dynamic state & the PDO index after 1980s are dominated by decadal variations as a result of emergence of a coupled negative feedback loop
- This feedback loop is disrupted in 2017 due to the formation of Kuroshio large meander south of Japan
- Similar basin-wide atmospheric responses are generated by the KE variability, irrespective if it is LM-forced or wind-forced
- Implications: KE dynamic state has likely entered a newly-set decadal cycle



- It remains a challenge to simulate intrinsically-forced Kuroshio path changes off Japan → a hindrance for decadal KE prediction