Improved Representation of Eddies in Fine Resolution Forecasting Systems Using Multi-Scale Data Assimilation of Satellite Altimetry

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Impact of Altimetry Data in Real-Time Mesoscale Prediction during the SPURS Field Campaign

AVISO

- The assimilation of the altimetry data played a significant role in the prediction of mesoscale eddies (OSTST 2013)
- The model encouragingly showed an ability of predicting submesoscale features (OSTST 2014)
- Today's topic: how the assimilation altimetry data improves representation of eddies down to tens of km



2012 Fall, Model Predicted Mesoscale Eddies



2013 Spring, Model Predicted Sub-mesoscale features

Busecke et al. (2014)

MODEL

Observed Mesoscale and Sub-Mesoscale Features During SPURS

Salinity Processes in the Upper Ocean Regional Study (SPURS)



Focus on the maximum sea surface salinity (max-SSS) area



Different Spatial Scales - Salinity Finer Structures



ROMS model

- 3-km resolution
- 50 levels
- 6-hourly GFS atmospheric forcing



Mesoscale and Sub-Mesoscale Characteristics

Remove

- Two monthly mean
- Zonal Mean

April 6, 2013





Mesoscale SSH Fronts Responsible for Submesoscale Vertical Velocity



Implication to data Assimilation: Constraints on mesoscale eddies partially constrain submesoscale features ?

Improved Mesoscales and Mesoscales with Data Assimilation



Average from March 5 through April 5, 2013 (Scaled for plotting)

- The SSH analysis is close to AVISO for scale larger than 250 km.
- The model with data assimilation realistically reproduces submesoscale features observed from high resolution SSTs and SPURS measurements.
- The vertical velocities is largest at a scale of around 20 km (likely depending on 3-km resolution).

How Altimetry Data Impacts Small-Scale Analysis and Forecast



$$\delta \mathbf{x}^{a} = \mathcal{K}(\varsigma^{o} - H\varsigma^{f})$$

$$\delta \varsigma^{a}, \delta T^{a}, \delta S^{a}, \delta \vec{V}^{a} \qquad \begin{array}{c} \text{Gain} \\ \text{Matrix} \end{array} \quad \begin{array}{c} \text{Altimetry} \\ \text{Data} \end{array}$$

Direct but smoothed constraint

(Li et al., 2015, MWR; 2015, Ocean Dyn.)



Partial constraint

SPURS Reanalsysis and Eddy Flux Analysis (1) Horizontal Divergence



$$10^{-7}$$
 psu/s=3.1psu/year

$$\frac{\partial \bar{s}}{\partial t} + \bar{u} \cdot \nabla \bar{s} + \bar{w} \frac{\partial \bar{s}}{\partial z} = -\nabla_H \cdot \bar{u} \cdot \bar{s} - \frac{\partial \bar{w} \cdot \bar{s}}{\partial z} + \bar{s}$$

The eddy flux divergence fresh and stratify the maximize salinity are

SPURS Reanalsysis and Eddy Flux Analysis (2) **Vertical Divergence**



• **Comparable to other regions?**

Salinity Budget for the Maximum Salinity (22.5°N-25.5°N)



Summary

- 1. The assimilation of altimetry data constrains eddies larger than 200 km and also improves the representation of submesoscales down to tens of km.
- Near surface, salinity and buoyancy have considerably stronger submesoscale features, which are subject to submesoscale vertical velocities associated with mesoscale SSH fronts.
- 3. Reanalysis has been produced for SPURS with the assimilation of altimetry data.
- 4. The eddy flux divergence analysis from the reanalysis shows that both mesoscale and submesoscale flux divergences account for a significant part of the salinity budget in the maximum salinity area.



How Does the Altimetry Data Impacts Representation of Meso-Scale Eddies in Data Assimilation?

AVISO vs Along Track

In Data Analysis

- Eddies down to 100 km along track (?)
- Eddies down to 200 km in 2D maps, with reduced amplitudes approaching 200 km (e.g., Chelton et al., 2011)
- Difference between the OI maps from four and two altimeters as large as 10 cm (Pascual et al., 2006)



In Data Assimilation

• Frontogenesis deterministic conditioned on the accurate placement of the mesoscale (Jacobs, 2014).

Multi-Scale 3DVAR with Background Error Covariance of **Multi-Decorrelation Length Scales**

δ



$$x = x_{L} + x_{S}$$

$$B = B_{L} + B_{S}$$

$$i = \sum_{x} J(\delta x) = \frac{1}{2} \delta x^{T} (B_{L} + B_{S})^{-1} \delta x + \frac{1}{2} (H \delta x - \delta y)^{T} R^{-1} (H \delta x - \delta y)$$

$$i = \sum_{x} Low resolution obs$$

$$\min_{\delta L} J(\delta x_{L}) = \frac{1}{2} \delta x_{L}^{T} B_{L}^{-1} \delta x_{L} + \frac{1}{2} (H \delta x_{L} - \delta y)^{T} (H B_{S} H^{T} + R)^{-1} (H \delta x_{L} - \delta y)$$

$$\min_{\delta x_{S}} J(\delta x_{S}) = \frac{1}{2} \delta x_{S}^{T} B_{S}^{-1} \delta x_{S} + \frac{1}{2} (H \delta x_{S} - \delta y)^{T} (H B_{L} H^{T} + R)^{-1} (H \delta x_{S} - \delta y)$$

High resolution obs





- 1. Fresher-warmer water intrusion
- 2. Intensification of mesoscale systems
- 3. Rapid stratification
- 4. Secondary circulation
- 5. Baroclinic instability

Baroclinity: Density Compensation and Enhancement





The density anti-compensation establishes intense density gradients and thus baroclinicity south of 24°N



Stratification: Up-Front Winds



- Up front winds
 - Stratification
- Frontogenesis