Effect of swell and wind-waves on the altimeter-derived estimates Analyzing real and simulated data



T. Moreau, N. Tran, M. Raynal, S. Labroue, J. Aublanc, L. Amarouche, R. Husson (CLS), C. Tison, S. Le Gac, F. Boy (CNES), A. Mouche (Ifremer), L. Aouf (Meteo France)



- To correct effect of non-gaussian statistics of wave surface
 - Estimator with higher order statistical descriptors of the distribution [Rodriguez, 1988; Kerbaol et Chapron, 1999]
 - o Removing non-negligible contribution of skewness to SSB [Srokosz, 1986]



- To correct effect of non-gaussian statistics of wave surface
 - Estimator with higher order statistical descriptors of the distribution [Rodriguez, 1988; Kerbaol et Chapron, 1999]
 - o Removing non-negligible contribution of skewness to SSB [Srokosz, 1986]
- CLS/CNES simulation studies showing swell and wind sea effects with gaussian surface wave properties
 - R&D on SAR-mode sensitivity (and LRM) to long waves
 - o Analyze of impact of wind-sea on LRM in CFOSAT project









- To correct effect of non-gaussian statistics of wave surface
 - Estimator with higher order statistical descriptors of the distribution [Rodriguez, 1988; Kerbaol et Chapron, 1999]
 - o Removing non-negligible contribution of skewness to SSB [Srokosz, 1986]
- CLS/CNES simulation studies showing swell and wind sea effects with gaussian surface wave properties
 - o R&D on SAR-mode sensitivity (and LRM) to long waves
 - Analyze of impact of wind-sea on LRM in CFOSAT project
- First evidence of impact of swell on real SAR altimetry data by *Aouf et Phalippou* [2015]
 - Showing Cryosat-2 wave height estimates likely biased by ocean swell (most noticeably for longest waves and swell fields propagating in a direction parallel to the satellite track)



• Wave conditions are frequently represented by **Hs** and **Tm**

$$m_x = \iint f^x S(f,\varphi) \, df \, d\varphi$$

Significant wave height: $Hs = 4 \ sqrt(m0)$ Mean wave period:Tm = sqrt(m0/m2)

 m_x : x – order moment S: wave height spectrum f: wave frequency ϕ : wave direction



• Wave conditions are frequently represented by **Hs** and **Tm**

$$m_x = \iint f^x S(f,\varphi) \, df \, d\varphi$$

Significant wave height: $Hs = 4 \ sqrt(m0)$ Mean wave period:Tm = sqrt(m0/m2)

 m_x : x – order moment S: wave height spectrum f: wave frequency ϕ : wave direction

- But no theoretical model for bivariate probability density function (pdf) exists
- In altimetry, processing design (retracking) only considers wave height (pdf based on Hs), ignoring Tm



• Wave conditions are frequently represented by **Hs** and **Tm**

$$m_x = \iint f^x S(f,\varphi) \, df \, d\varphi$$

But no theoretical model for bivariate probability density function (pdf) exists

Hs), ignoring Tm

wave periods (i.e. swell)

In altimetry, processing design (retracking) only considers wave height (pdf based on

With increasing spatial resolution (from 10km to 300m in SAR-mode), **radar measurements**

shall become sensitive to long wavelength

Significant wave height: $Hs = 4 \ sqrt(m0)$ Mean wave period:Tm = sqrt(m0/m2)

- m_x: x order moment
 S: wave height spectrum
 f: wave frequency
 φ: wave direction
- Ocean swell Sea (fetch Wave spectra swell wind sea



• For a given Hs (or Tm respectively) sea surface may have different aspects which may result in different altimetry radar returns



fixed Hs, variable T

variable Hs, fixed T



 For a given Hs (or Tm respectively) sea surface may have different aspects which may result in different altimetry radar returns



 In addition, ocean surface is complex with mixing wave systems (crossing swell fields and wind sea). Two wave fields with the same Hs and Tm may also be different in detail.









• Doppler echo shapes are distorted, and retrieved parameters altered





 Doppler echo shapes are distorted, and retrieved parameters altered





- Doppler echo shapes are distorted, and retrieved parameters altered
- Depend on Hs, Tpeak, orientation of wave propagation, location of the focusing point on surface





- Doppler echo shapes are distorted, and retrieved parameters altered
- Depend on Hs, Tpeak, orientation of wave propagation, location of the focusing point on surface
- Brown assumption no more valid (surface statistics in small delay-Doppler cells non-Gaussian and non homogeneous with each other)





- Doppler echo shapes are distorted, and retrieved parameters altered
- Depend on Hs, Tpeak, orientation of wave propagation, location of the focusing point on surface
- Brown assumption no more valid (surface statistics in small delay-Doppler cells non-Gaussian and non homogeneous with each other)
- Physical ocean model (retracker) not adapted for these particular wave systems





- Doppler echo shapes are distorted, and retrieved parameters altered
- Depend on Hs, Tpeak, orientation of wave propagation, location of the focusing point on surface
- Brown assumption no more valid (surface statistics in small delay-Doppler cells non-Gaussian and non homogeneous with each other)
- Physical ocean model (retracker) not adapted for these particular wave systems
- Features also seen in LRM, for swell and particular wind-sea wave conditions, inducing potential impact on estimates





- Doppler echo shapes are distorted, and retrieved parameters altered
- Depend on Hs, Tpeak, orientation of wave propagation, location of the focusing point on surface
- Brown assumption no more valid (surface statistics in small delay-Doppler cells non-Gaussian and non homogeneous with each other)
- Physical ocean model (retracker) not adapted for these particular wave systems
- Features also seen in LRM, for swell and particular wind-sea wave conditions, inducing potential impact on estimates

➔ To identify these effects with real data and characterize impacts on surface altimeter parameters





DATASETS DESCRIPTION

- 8-months of altimeter data in open ocean from may to december 2015
 - o Jason-2 Low Resolution Mode (LRM) @ 1-Hz/20-Hz



Cryosat-2 in SAR mode SAR mode and LRM-like (PLRM) processed by CNES CPP @ 1-Hz/20-Hz

IFREMER WaveWatch3 products

- Global 0.5°x0.5° wave grid
- ECMWF surface wind forcing / 3-hours step (with no altimetric data assimilation)
- Fields used:
 - Hs
 - Tm (mean wave period)
 - Partitions: wind-sea; swell#1; swell#2; swell#3 Hs, Dir, Tp (peak period)







0

DATASETS DESCRIPTION

- 8-months of altimeter data in open ocean from may to december 2015
 - o Jason-2 Low Resolution Mode (LRM) @ 1-Hz/20-Hz

SAR mode and LRM-like (P)

Cryosat-2 in SAR mode Ο + Sentinel-3A in SAR mode



©UCL

IFREMER W

Global 0.5°x0.5° wave grid

processed by

@ 1-H

- ECMWF surface wind forcing / 3-hours step (with no altimetric data assimilation)
- Fields used:
 - Hs
 - Tm (mean wave period)
 - Partitions: wind-sea; swell#1; swell#2; swell#3 Hs, Dir, Tp (peak period)





J2 SWH vs WW3 Hs



 Good agreement between altimeter and WW3 Hs values



J2 SWH vs WW3 Hs



- Good agreement between altimeter and WW3 Hs values
- But their difference reveal features depending on wave period values

May impact wind/SSB computation



➔ Analyse over longer time period, separate wave systems and geographical regions to better understand these trends

➔ Analyse the consistency of WW3 as a reference (by comparing with SAR imagery estimates for example)













T02 = 6 s

T02 = 7 s

T02 = 8 s

R

- T02 = 9 s

5



SOSCLS

OSTST meeting – La Rochelle – 1-4 November 2016

6

T02 = 3 s

T02 = 5 s

T02 = 6 s

T02 = 7 s

T02 = 8 s

- T02 = 9 s

T02 = 3 s

T02 = 4

T02 = 5

T02 = 6 s

T02 = 7 s

- T02 = 8 s

- T02 = 9 s

5

4

- T02 = 4 s



OSTST meeting - La Rochelle - 1-4 November 2016

6

T02 = 3 s

T02 = 4

T02 = 5

T02 = 6

T02 = 7

T02 = 85

T02 = 3

T02 = 8

- T02 = 9 :

5

T02 = 9

CS2 SARM AND PLRM STD



SOSCLS

OSTST meeting - La Rochelle - 1-4 November 2016

CS2 SARM AND PLRM STD



SOSCLS

OSTST meeting – La Rochelle – 1-4 November 2016

CS2 SARM AND PLRM STD



OSTST meeting – La Rochelle – 1-4 November 2016

CS2 AND S3A SARM RANGE STD



- Dependence of 20-Hz range std on Tm confirmed with S3A data in SAR-mode
- Much less significant dependency of 20-Hz range std on Tm observed in S3A P-LRM data (in backslide)



CS2 SWH STD wrt DIRECTION



- **20-Hz SWH std sensitive to wave direction in SAR-mode** but no apparent dependency of SWH measurements in direction (in backslide)
- Higher dispersion (over 7km) for wave propagation parallel to satellite ground track displacement
- No angular effect on 20-Hz SWH std in P-LRM (cylindrical symmetry)



CS2 RANGE STD wrt DIRECTION



- **20-Hz range std sensitive to wave direction in SAR-mode** but no apparent dependency of range measurements in direction (in backslide)
- No angular effect on 20-Hz range std in P-LRM



S3A RANGE STD wrt DIRECTION

Range Std for 2 < swh < 3



 20-Hz range std sensitivity to Tm and wave direction in SAR-mode confirmed with S3A data



S3A RANGE STD wrt DIRECTION

Range Std for 2 < swh < 3





 20-Hz range std sensitivity to Tm and wave direction in SAR-mode confirmed with S3A data Occurrences of 20-Hz range std in SARM greater than in PRLM are observed for waves propagating parallel to orbit track over long segments



56 satellite positions

11111

COMPLEMENTARY SIMULATIONS

 Use an *end-to-end* radar altimeter simulator able to generate waveforms in SAR and conventional modes for any instrumental configuration and different scenario of waves system

→ To study the dispersion of 20-Hz SARM and LRM measurements and dependencies



 Shape of Doppler power echoes impacted by the wave period and Hs, but also the direction



56 satellite positions

COMPLEMENTARY SIMULATIONS

 Use an *end-to-end* radar altimeter simulator able to generate waveforms in SAR and conventional modes for any instrumental configuration and different scenario of waves system

→ To study the dispersion of 20-Hz SARM and LRM measurements and dependencies



 Simulations confirm the sensitivity of SARM measurements (in dispersion) to waves period and direction while no impact is observed on LRM (not shown here)



CONCLUSIONS & PERSPECTIVES

- Impact of swells and wind-waves on the altimeter-derived estimates has been clearly identified at small scale (in terms of dispersion over 7km)
- Triple colocation alti/WW3/S-1 to assess their consistencies is on-going





- Use of indirect SAR (S-1) measurements by propagating synthetic swell fields may be considered where static colocations with S-1 are not available
- To analyze sea state effects on Ka-band altimeter configurations (Altika)
- A more thorough study of this issue is planned to be done with Sentinel-3A data to better characterize those effects and biases, and to reveal any potential anomalies in SAR (and LRM) altimetry data and in the sea level content
- The results of this analysis soon in: "Impact of swell and wind-waves on SAR altimeter-derived estimates", Moreau et al. [in preparation]



THANK YOU !!

tmoreau@cls.fr



OSTST meeting – La Rochelle – 1-4 November 2016

WW3 MODEL: WAVE PERIOD AND WAVELENGTH



- Tpeak swell > Tpeak wind-sea
- Tm is lower than individual Tpeak
- Deep water: L ~ 1.6 Tm²



 Tm of 13.5s corresponds to SAR altimeter resolution ~ 300 m



WAVE SYSTEMS FROM WW3 MODEL





OSTST meeting – La Rochelle – 1-4 November 2016

WAVE SYSTEMS FROM WW3 MODEL





OSTST meeting – La Rochelle – 1-4 November 2016

WAVE SYSTEMS FROM WW3 MODEL





threshold = 0.01 m



CLS

OSTST meeting – La Rochelle – 1-4 November 2016

CS2 SWH SARM vs PLRM





OSTST meeting – La Rochelle – 1-4 November 2016

EFFECT ON RANGE



- No apparent dependency of the range difference on the wave mean period values
- But the dispersion of their difference depends on the wave period
- Correlation observed between range difference and SWH difference





S3A SARM AND P-LRM RANGE STD







CS2 SWH vs WW3 Hs wrt DIRECTION



 No apparent dependency of SWH (and range) measurements in direction



WAVEFORMS ANALYSIS







- Higher waveform amplitude dispersion in SAR-mode as Tm increases
- GOF of the SAR altimeter models worsens as Tm increases (nearly stable in PLRM)
- Larger SARM SWH and SLA std as Tm increases (almost unchanged in PLRM)



