

Performance and quality assessment of the new CMEMS global ocean monitoring and forecasting real-time system

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- Main ingredients of an analysis and forecasting system
- Main recent system updates and illustration of some of them
- Performance and quality of the new system
- Conclusions and Perspectives



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Ocean numerical model + observations + data assimilation (mathematical methods)

The aim of a data assimilation method is to force the ocean model to be as close as possible of the observations available in the past to obtain the best forecast in the future, taking into account observations and model errors.





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Main updates of the system

Previous system

Service

Weekly 7-day analyses and daily 7-day forecasts

Model (ocean and sea-ice)

MERCATOR

OCEAN

- NEMO 3.1, horizontal 1/12° et 50 vertical levels
- LIM2_EVP (sea-ice model)
- 3h ECMWF atmospheric forcing

Data assimilated

- SLA (CMEMS) + MDT CNES CLS09
- SST AV/HRP.
- T/S vertical profiles (CMEMS)

Assimilation

- SEEK (Kalman filter)
- 3D-Var large scales bias correction



R&D activities have been conducted these last years to improve this system

New system

(operational since Wednesday, October 19, 2016)

<u>Service</u>

Daily 10-day forecasts

Model

- Ice-sheets and glaciers melting added to the runoffs
- Large scale correction of precipitations
- Global steric effect added to SSH

Data assimilated & Assimilation

- MDT CNES-CLS13
- Assimilation SST (OSTIA CMEMS)
- Assimilation sea-ice concentration (OSI-SAF CMEMS)
- WOA 2013 "weak assimilation" below 2000m
- Adaptive tuning of SLA and SST observation errors

Available Period: 2007 → now



Without correction

With correction







Mean surface salinity innovation (2011)



The objective of this diagnostic is to improve the error specification by tuning an adaptive weight coefficient α acting on the error of each assimilated observation



Either for SLA or SST, we improve the performances of the system using this adaptive tuning





Adaptive tuning of observation errors (SST)

Section at latitude 3°N (seasonal cycle removed)

MERCATOR OCEAN

Model SST anomaly 2013 2012 2011 June 2010 Ы 2009 2008 June 2007 Feb Jan 140 160 220 240 260 280 140 160 180 200 220 240 260 280 180 200 °C °C -5-4-3-2-1012345 -1 0 1

Obs error SST anomaly

The SST anomalies in the equatorial Pacific clearly show the propagation westwards of TIW in the second half of the year. This is more pronounced during episodes of La Nina (mid-2007 and mid-2010).

The error anomalies estimated by "Desroziers method" show that the error increases when these TIWs are more marked.

When model SST and OSTIA SST are smoother, the error decreases.

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The SST error adapts both to the seasonal and interannual fluctuations



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Previous system

SLA residual



-0.861 meters Nov = 1.377 -0.075 -0.038 0.000 0.038 0.075 0.113 0.150 -0.150 -0.113

SLA RMS residual



Min =0.001 meters Mox =1.377 0.000 0.019 0.038 0.056 0.075 0.094 0.113 0.131 New system

SLA residual



Min =-1.412 meters Max = 1								
-0.150	-0.113	-0.075	-0.038	0.000	0.038	0.075	0.113	0.150

SLA RMS residual





New system





→ This allows having more accurate description of the water masses

Global mean SSH (2007-2015) Better mass balance and GMSL trend

OCEAN



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RMS (model – OSISAF)

Not yet assimilated in the system ...

Previous system

New system

U drift innovation in 2008-2013

Mean surface currents (OND 2008)

The Ekman spiral is improved (angle between the model or observed surface current and the wind

for different wind intensities)

The zonal and meridional slopes are improved

MERCATOR

1.0 1.0 PREVIOUS NEW Angle to wind for year 2008 OND at 0m m/s m/s Fit Slope=0.56 Fit Slope=0.64 (PREVIOUS) MODEL MODEL .⊆ **0.5** .⊆ **0.5** 80 PSY3V3R3 i PSY4V3R1 i 0.0 PSY3V3R3 in m/s PSY4V3R1 60 Obs (ARGO+AOML) MODEL MODEL 40 Ξ. m/s _-0.5 _____.5 20 Model \cap -t.0-t.0-1.0 -0.5 0.0 0.5 0.0 0.5 1.0 -1.0 -0.5 1.0 -20 U OBS drift 20080ND in m/s U OBS drift 20080ND in m/s 10 15 wind speed in m/s Angle to wind for year 2008 OND at 0m 1.0 1.0 PREVIOUS NEW (NEW) PSY3V3R3 in m/s 6 6 m/s Fit Slope=0.43 Fit Slope=0.51 V MODEL 80 MODEL ,⊆ 0.5 60 . PSY4V3R1 i PSY4V3R1 in PSY3V3R3 in m/s Obs (ARGO+AOML) 40 MODEL MODEL 20 Model m/s ~-0.5 0 -2a -1.0-1.0 10 15 5 0.0 0.5 1.0 -1.0 -0.5 0.0 0.5 1.0 -1.0 -0.5wind speed in m/s V OBS drift 2008OND in m/s V OBS drift 2008OND in m/s

Crash AF447 flight on the night of 1st June 2009 We would have done better ...

MERCATOR OCEAN

On the night of 1st June 2009, a Rio-Paris Air France flight (AF447) disappeared in the western tropical Atlantic Ocean. The first debris was located 5 days after the accident...

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- □ The new global system has a quite good statistical behavior with a accurate representation of the water masses, the surface fields and the mesoscale activity.
- □ Most of the components of the new system are improved: global mass balance, 3D T/S, SSH sea-ice, currents, ...
- □ 10 years of homogeneous products are now available \rightarrow "little reanalysis" at 1/12°.

Some other topics on which we have started to work and we want to continue in the future...

- Improvement of coastal runoff, atmospheric forcing, inclusion of the tide in the system,
- Use of the LIM3 sea-ice model,
- Assimilation of velocity observations,
- Assimilation of SSS satellite observations,
- From weekly to Daily Analysis,
- Production of a global 1/12° reanalysis covering the altimetric period,
- Take into account in the system the interaction between currents and waves,
- Extend the Data Assimilation to the ocean biogeochemical component.

Extra Slides

In order to obtain a consistent model equivalent for SLA, different space/time filters are applied.

These time filters act both on the sea level height and the barotropic height computed by the ocean model in order to remove high frequency barotropic signals.

First, the SSH is averaged over a one day window, and two running means of barotropic height spatially averaged (over a box of \sim 5°) are removed from it.

The model equivalent for SLA is :

$$SLA_{eq}^{f} = \overline{SSH}_{1day}$$
 - LargeScale (\overline{HBAR}_{1day} - \overline{HBAR}_{21days}) - MDT

where MDT is the Mean Dynamic Topography

The running mean over a one day window eliminates barotropic high frequencies (period < 1 day) and the one over a 21 day window eliminates barotropic signals with periods < 21 days, which correspond to the low pass filter (Lanczos) applied to along track altimetric data.

The objective was to collaccount:

- SLA innovations,
- HDYN innovations cor
- Most recent satellite-
- Latest CNES-CLS13 M

meters

vious systems, taking into

gy is used),

Some diagnostics with the previous system

mean sea level innovation in 2009

mean sea level residual in 2009

- ✓ Use of Dai and Trenberth 2009 rivers runoff database (instead of Dai and Trenberth 2002) : New data (mostly from recent years) from several sources, 925 collected rivers (73% of global total runoff), use streamflow simulated by Community Land Model version 3 (CLM3) to fill the gaps, all land areas except Antarctica and Greenland.
- ✓ Construction of a freshwater flux from ice-sheets melting in Greenland and Antarctica using Altiberg icebergs database project (Tournadre *et al.*, 2013).

EXA vfec = SLA Sat - DAC - LWE - TIDE where DAC = IB LF + MOG2D HF **EXA** similation: $SLA_{eq} = \overline{SSH}_{1day} - \underline{LargeScale}(\overline{HBAR}_{1day} - \overline{HBAR}_{21days}) - MDT - COMPARENT SHOULD - COMPARENT SHOULD - COMPARENT SHOULD - COMPARENT SHOULD - COMPARENT - COMPARE$

 $Model: \frac{\partial U}{\partial t} = \dots \qquad \text{If the model is forced by atmospheric pressure} + tide in the model}$ $Assimilati \text{ on }: \begin{cases} SLA_{eq} = (\overline{SSH} - SSp_{TDE}^{25h}) - (\overline{SSH} + SSH_{MOG 2D}] - MDT \\ SSH_{IB} = -\frac{1}{g\rho}(p_A - \langle P \rangle): & \text{Inverse Barometer} \\ p_A: \text{Atmospheric pressure (From ECMWF model)} \\ \langle P \rangle: \text{Spatial Average (over the ocean)} \end{cases}$

 \rightarrow Our future need: SLA_{vfec} + MOG2D_{HF} (+ IB_{LF} ?)