

**→ 25 YEARS OF PROGRESS
IN RADAR ALTIMETRY SYMPOSIUM**

OSTST MEETING

24–29 September 2018
Ponta Delgada, São Miguel Island
Azores Archipelago, Portugal

**The new time-variable gravity field model
RL04 from CNES/GRGS
based on GRACE+SLR
for the POD of altimetric satellites**

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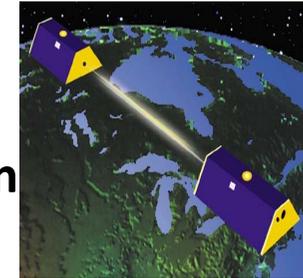
- ❖ Precise orbit determination is a key element in the overall accuracy of the altimetric measurements.
- ❖ Since 2002, thanks to the GRACE (and GOCE) missions, we have now a very good knowledge of the Earth gravity field and its time evolution.
- ❖ Based on 14 years of GRACE data (2002.5-2016.5), 3 years of GOCE data and 33 years of SLR data (1985-2018), the **EIGEN-GRGS.RL04.MEAN-FIELD** is the gravity model that is proposed for the GDR-F standards.
- ❖ It contains a time-variable gravity (TVG) part until degree and order 90, and a static part coming from the model GOCE-DIR5 up to degree and order 300.
- ❖ The TVG part is modeled for each year between August 2002 and June 2016 as an annual bias + slope + annual and semi-annual periodic components.
- ❖ For the low degrees of the gravity field, the TVG part prior to August 2002 will either :
 - Be modeled, for degree 2 only, by SLR data from January 1985 to July 2002
 - Or be modeled in a more ambitious way, i.e. by a 5x5 SH model or through a “mascon” approach (John Moyard’s presentation at the IDS workshop).

Data processing in the RL04 reprocessing

(June – December 2017)

GRACE (L-1B “Version2” data)

- K-Band Range-Rate data ($\sigma_{\text{apriori}} = .1 \mu\text{m/s}$)
- GPS data (1-day arcs, $\sigma_{\text{code}} = 80 \text{ cm}$, $\sigma_{\text{phase}} = 20 \text{ mm} / 30\text{s}$ resolution)
- ACC and SCA data (KBR CoP coordinates solved once / day)



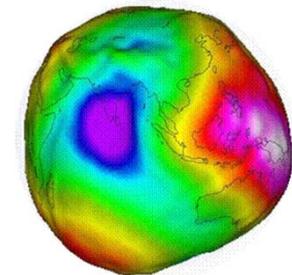
SLR

- Lageos1/2 data (10-day arcs, $\sigma_{\text{apriori}} = 6 \text{ mm}$)
- Starlette/Stella data (5-day arcs, $\sigma_{\text{apriori}} = 10 \text{ mm}$)



Physical parameters present in the normal equations

- Gravity spherical harmonic coefficients complete to **degree and order 90** (truncated to **30 for LAGEOS** and **40 for GPS** data)
- Ocean tides s. h. coefficients for 14 tidal waves with maximum degree/order ≤ 30 (not used yet)



Models used

Dynamical models

Gravity	<i>EIGEN-GRGS.RL03-v2.MEAN-FIELD</i>
Ocean tide	<i>FES2014 (Legos)</i>
Atmosphere	<i>3-D ECMWF ERA-interim pressure grids / 3 hrs</i>
Ocean mass model	<i>TUGO (Legos) / 3 hrs</i>
Atmospheric tides	<i>→ Not necessary because of the 3 hrs dealiasing time sampling</i>
3 rd body	<i>Sun, Moon, 6 planets (DE405)</i>
Solid Earth tides	<i>IERS Conventions 2010</i>
Pole tides	<i>IERS Conventions 2010</i>
Non gravitational	<i>Accelerometer data (+biases and scale factors)</i>

Geometrical models

SLR stations	<i>ITRF2014 coordinates</i>
GPS	<i>IGS Repro-2 orbits & clocks</i>

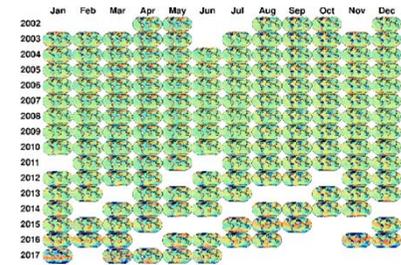
Other models

Hydrology	Taken into account by the a priori gravity field
Glacial Isostatic Adjustment	

From GRACE monthly solutions to mean gravity models

❖ Using directly GRACE/GRACE-FO monthly solutions is not appropriate for POD because of:

- Data gaps in the GRACE time series (e.g. after 2011 and between GRACE and GRACE-FO)
- The problem of extrapolation before 2002 and after 2016



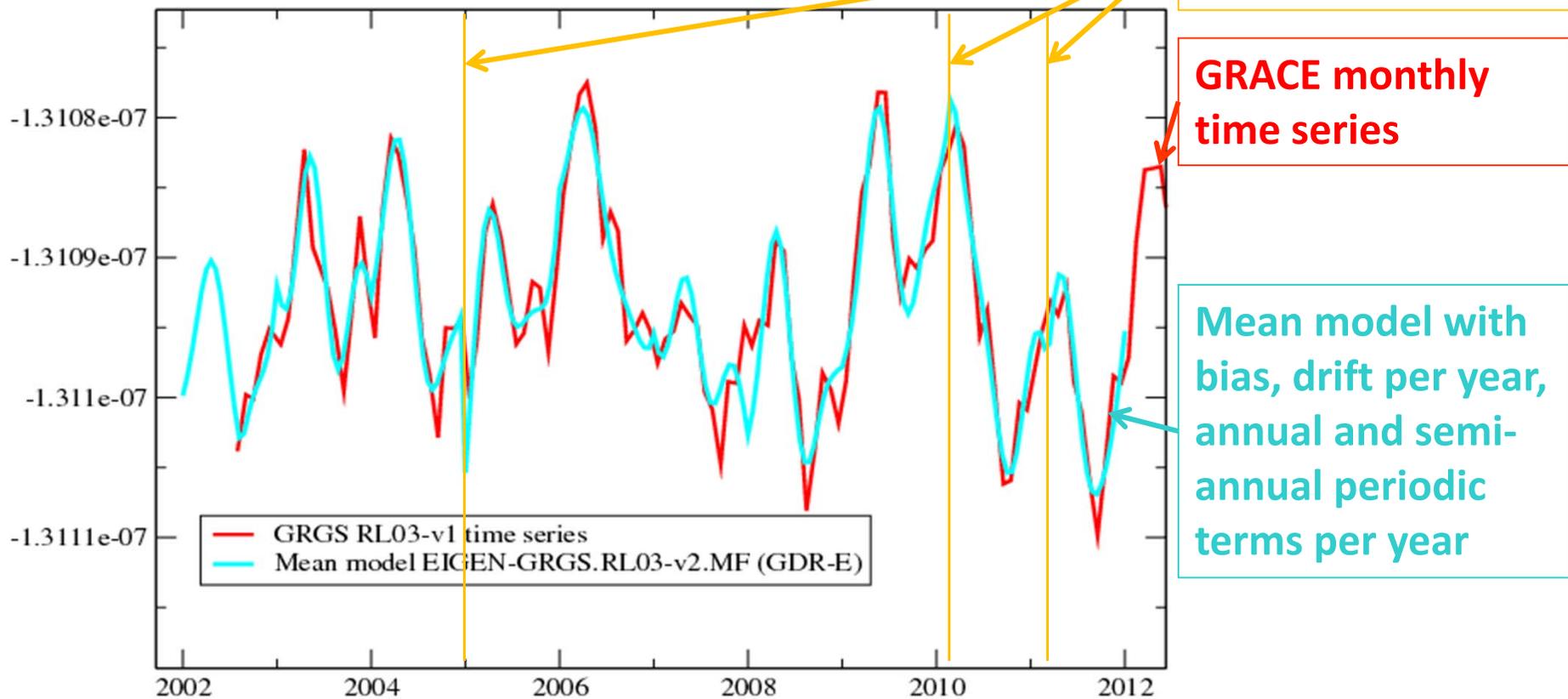
❖ Mean models are now generated from time series

- Fitting each series of monthly spherical harmonic coefficients by a set of 6 parameters :
 - Yearly bias and slope : piecewise linear function except in case of ...
 - Jumps caused by big earthquakes (so far : Sumatra/2005.0, Concepcion/2010.2 and Tohoku/2011.2)
 - Annual and semi-annual sine/cosine functions (with continuity constraints at hinge epochs)
- It means 750 000 coefficients for a 90x90 spherical harmonic model

RL04 mean model

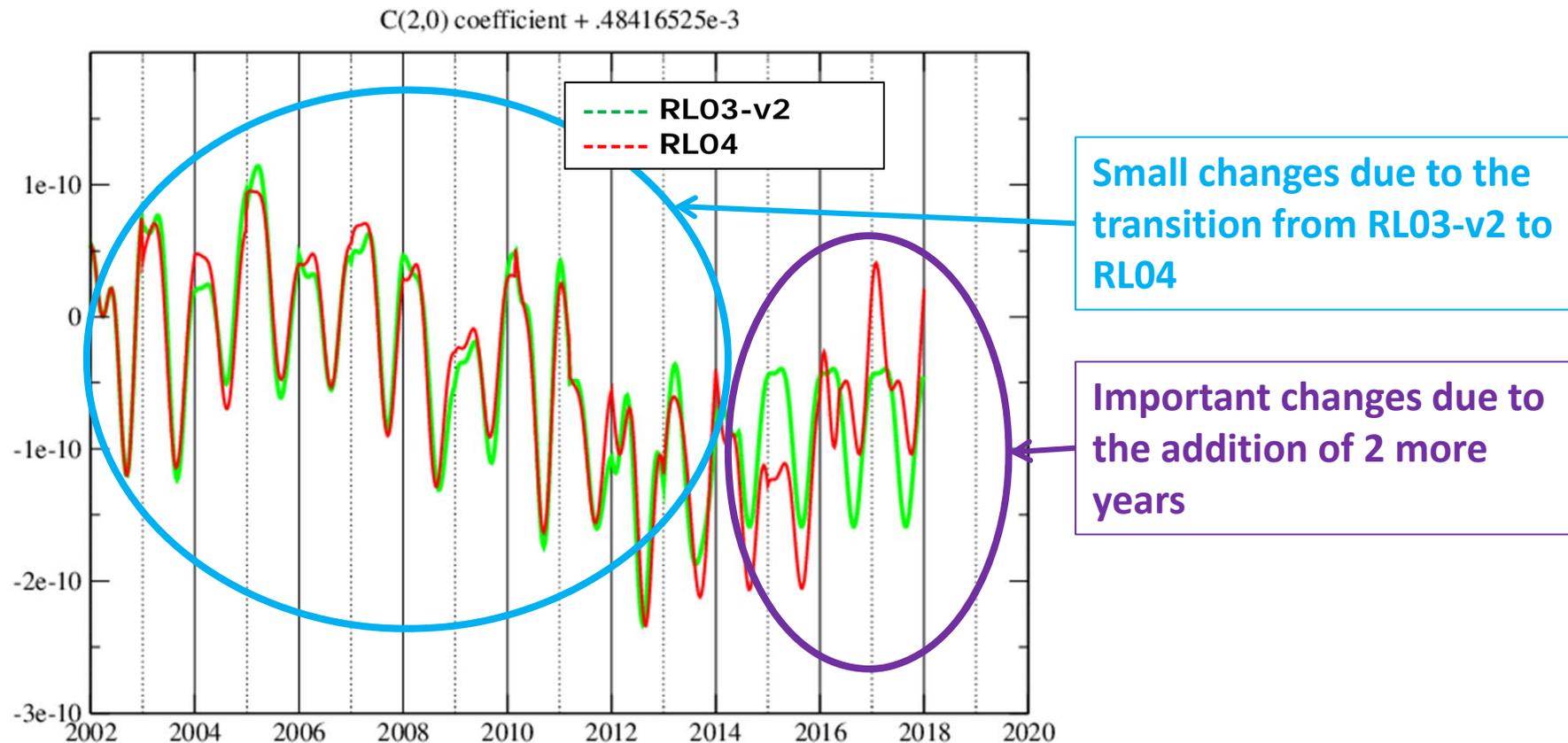
- Example for one spherical harmonic coefficient:

Normalized S (10,01) coefficient



Mean model: from RL03-v2 to RL04

- The new mean field updates the previous one over 2 years: mid-2014 to mid-2016.
- Example for the C(2,0) spherical harmonic coefficient:



Update of the mean model from -v2 to v3

v3

- Extrapolation vs. real data after 2 years: difference between mean-field -v2 and mean field -v3 at mid-2016

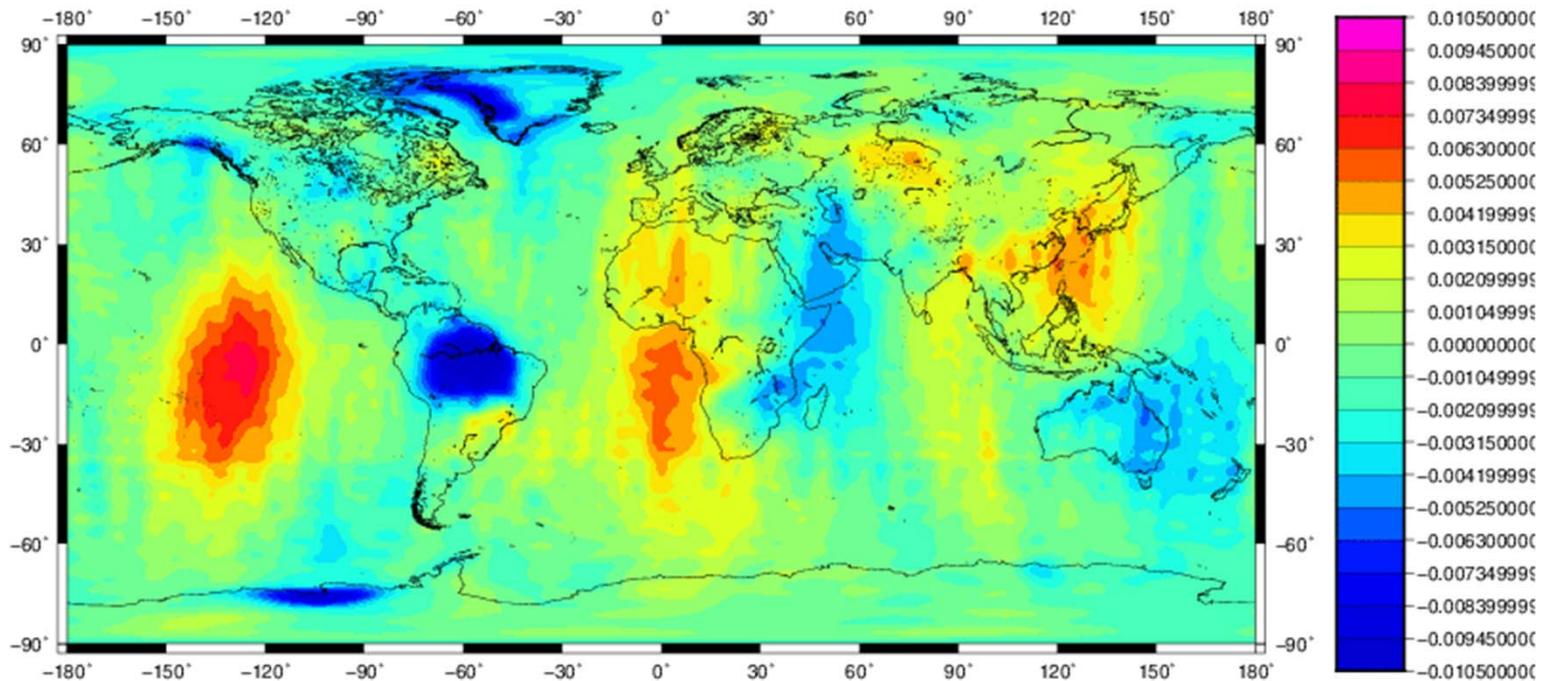
Geoid comparison at date: April 1st, 2016

EIGEN-GRGS.RL03-v3.MEAN-FIELD - EIGEN-GRGS.RL03-v2.MEAN-FIELD

degree 0002 to 0080

(unit : m)

(mean: 0.0000 / st.dev: 0.0027 / min: -0.0185 / max: 0.0079)



Improvement of RL04 wrt RL03-v2

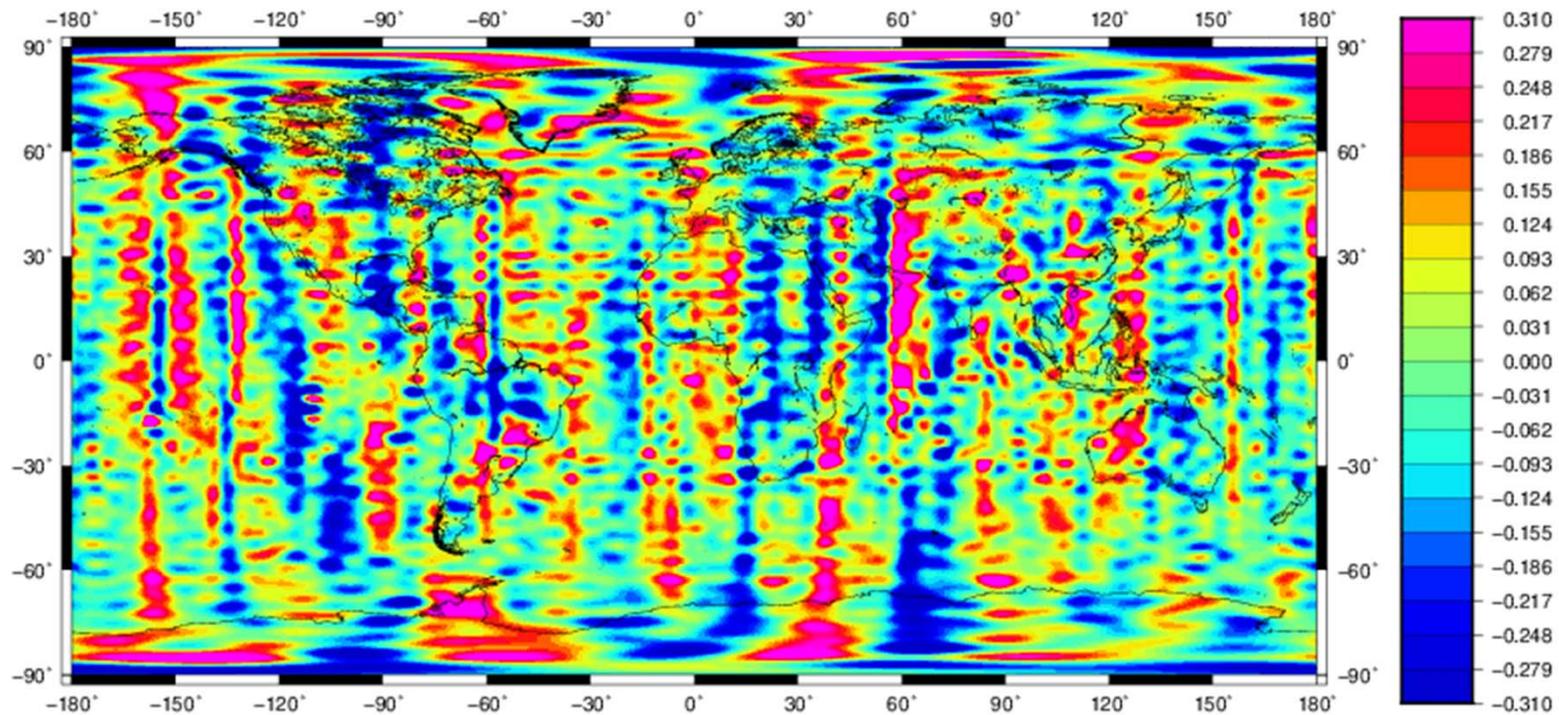
Equivalent Water Height comparison:

20170515 solution-RL04 test-6 w/o spher. cap – EIGEN-GRGS.RL03-v2.MF.MSE

degree 0002 to 0090

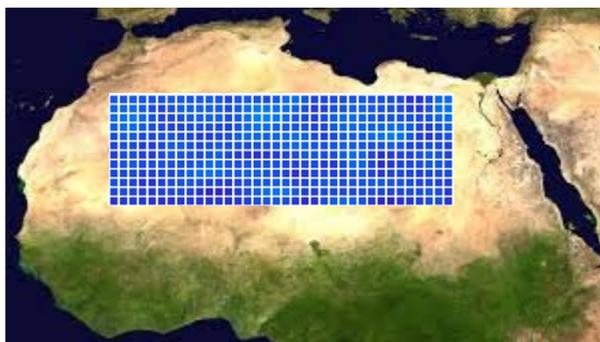
(unit : m)

(mean: -0.0000 / st.dev: 0.1392 / min: -0.6185 / max: 0.7640)



RL04 evaluation

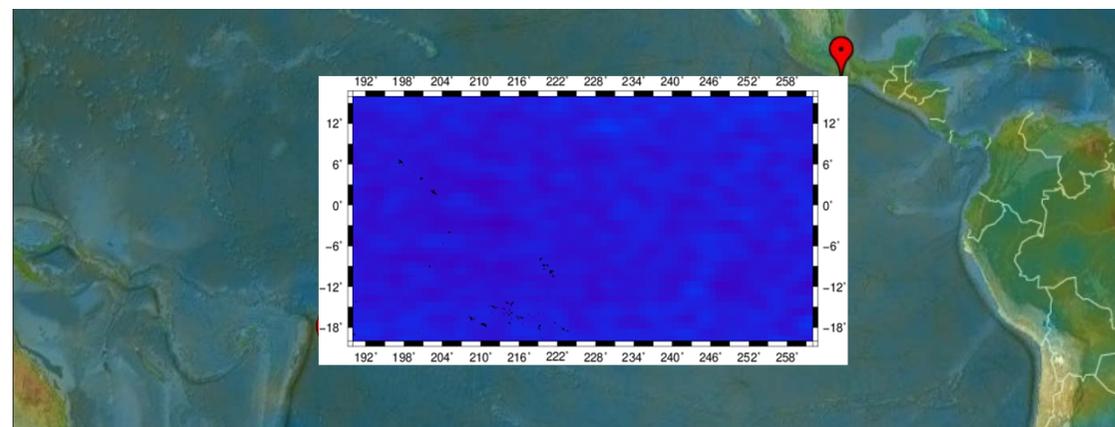
Noise assessment can be made in areas with no or very little mass variations: Sahara and Gobi deserts, East Antarctica, South and Equatorial Pacific



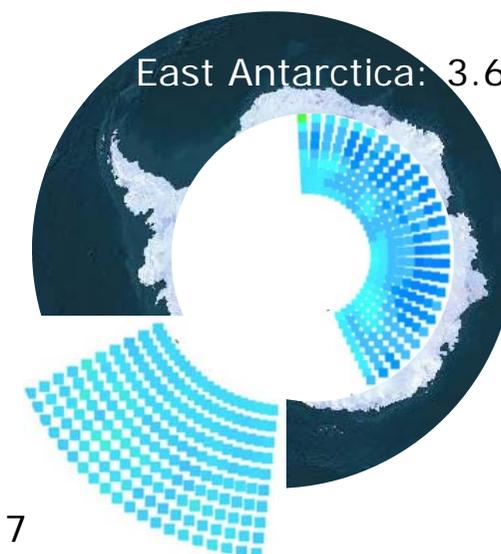
Sahara desert: 2.2 Mkm²



Gobi desert: 1.6
Mkm²



Equatorial Pacific: 31.6 Mkm²



East Antarctica: 3.6 Mkm²

South Pacific 6.7
Mkm²

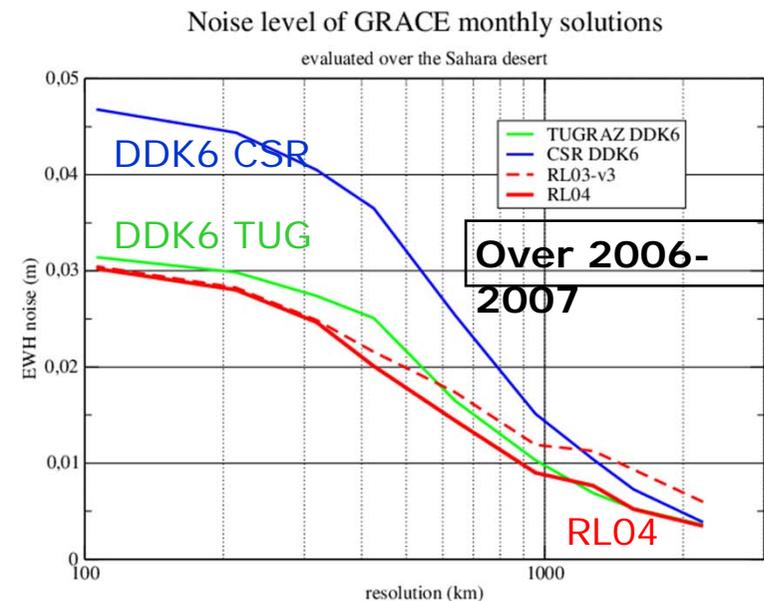
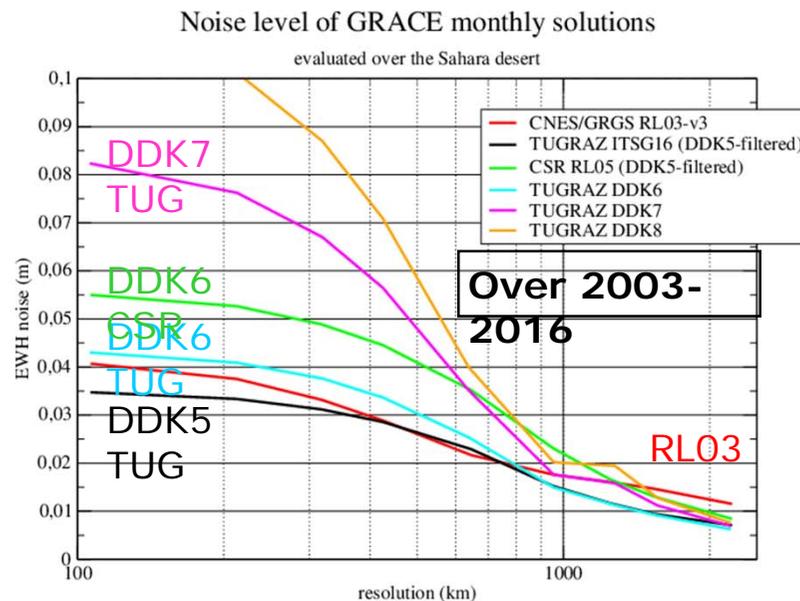
Noise assessment over the Sahara

The Sahara desert shows very little hydrological variations. We have delimited a rectangular zone of 2.2 Mkm² where almost no gravity variation is suspected (except a small depletion of 1.3 mm/yr in South Libya).

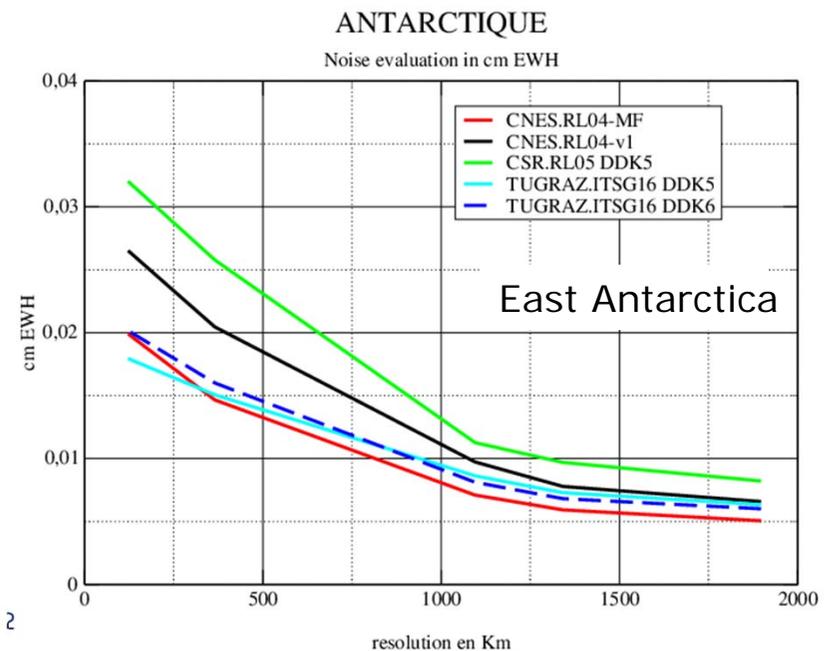
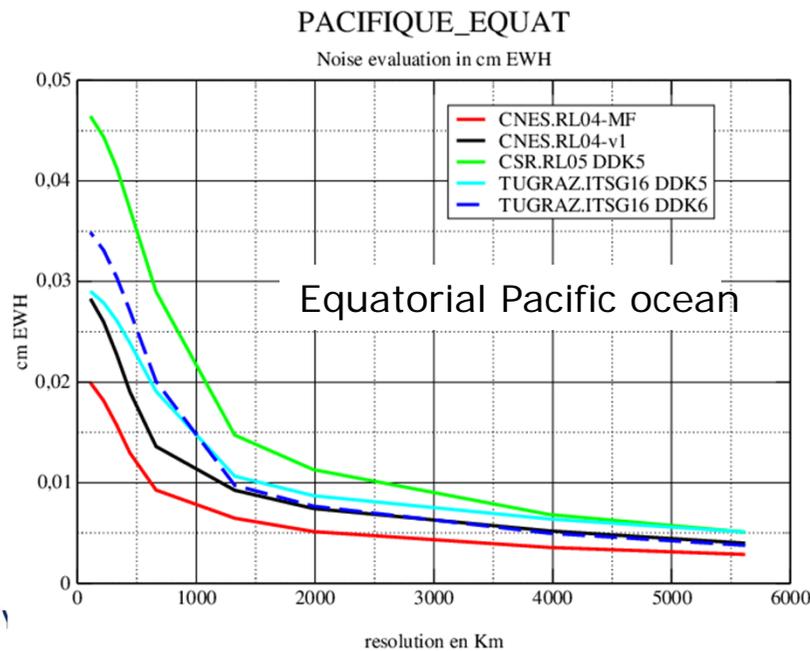
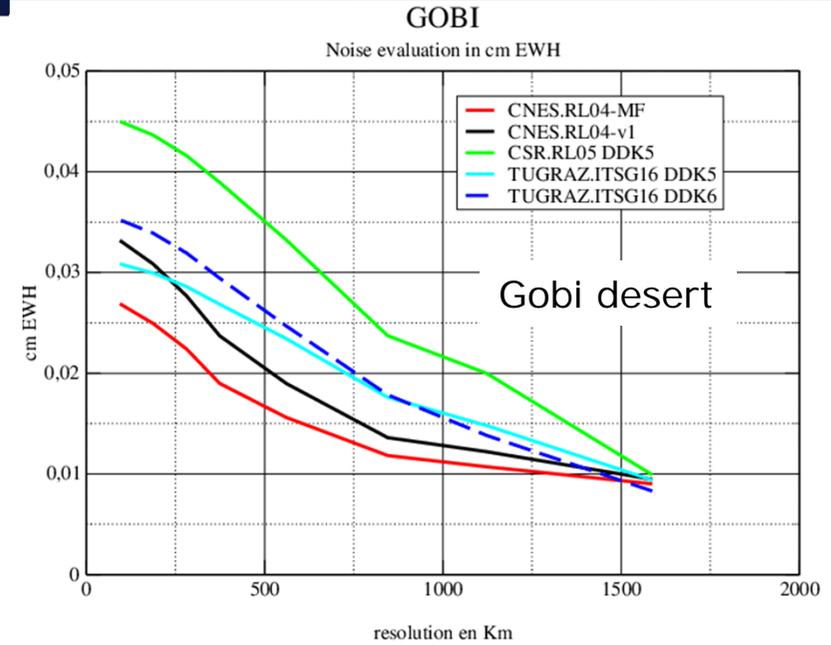
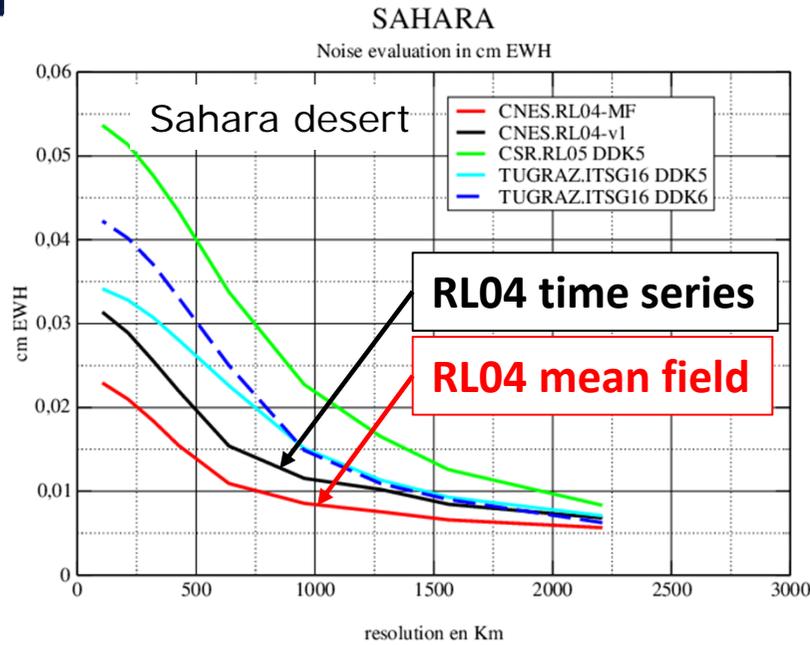
It is hence well dedicated to control the quality of gravity field variation models.

The surface is first divided in 2 deg.*2 deg. blocs (↔ degree/order 90), then averaged in blocs of larger size up to 20 deg.*20 deg. Drift and annual/semi-annual variations are fitted a priori.

Different time-varying gravity models with various DDK filters (Kusche et al.) are compared spectrally in this way from 100 km to 2200 km.

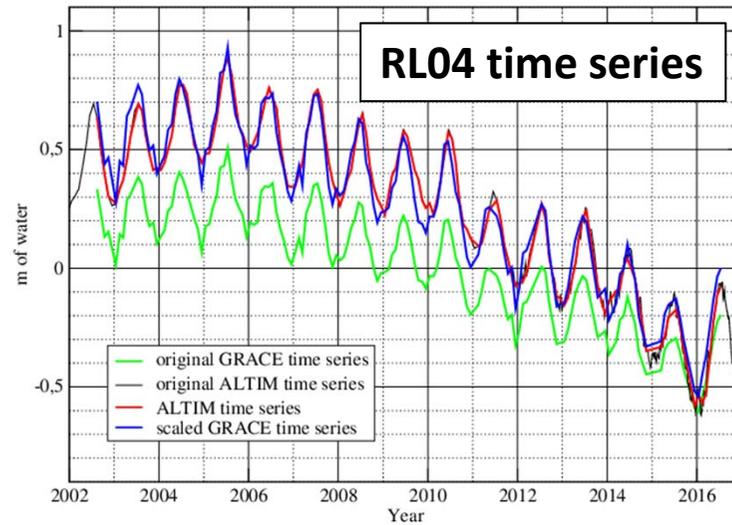


Noise assessment over "deserts"

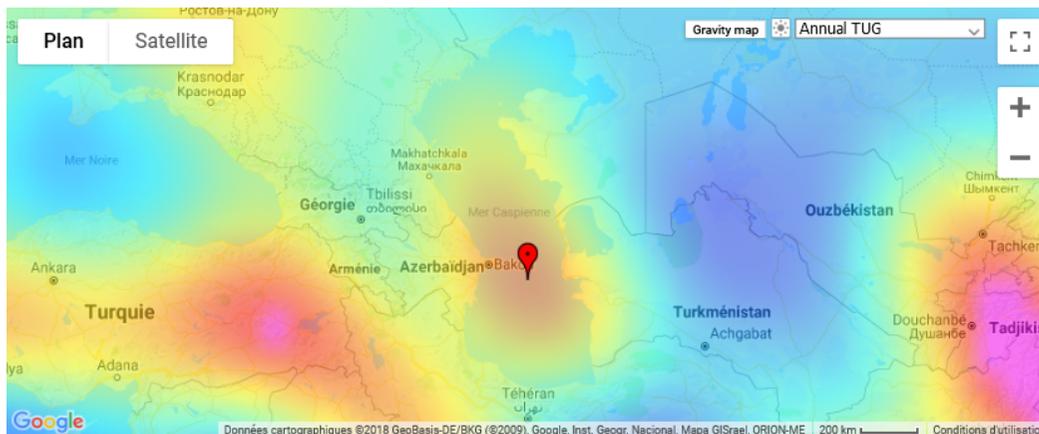
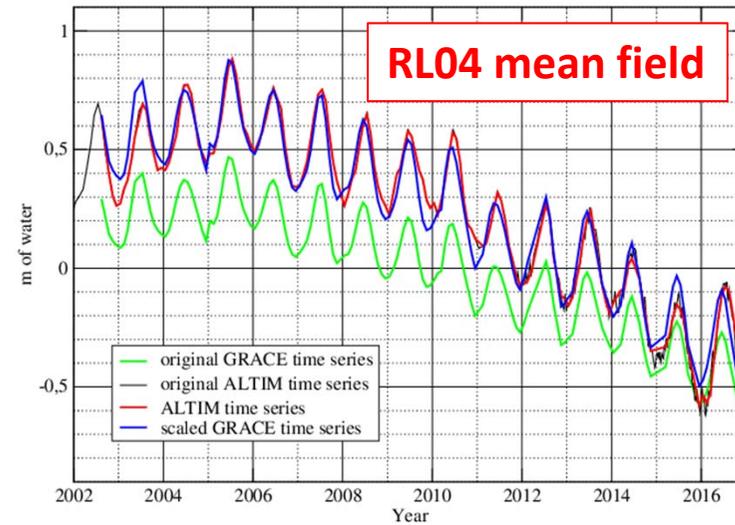


Signal assessment by comparison to altimetry (Caspian Sea)

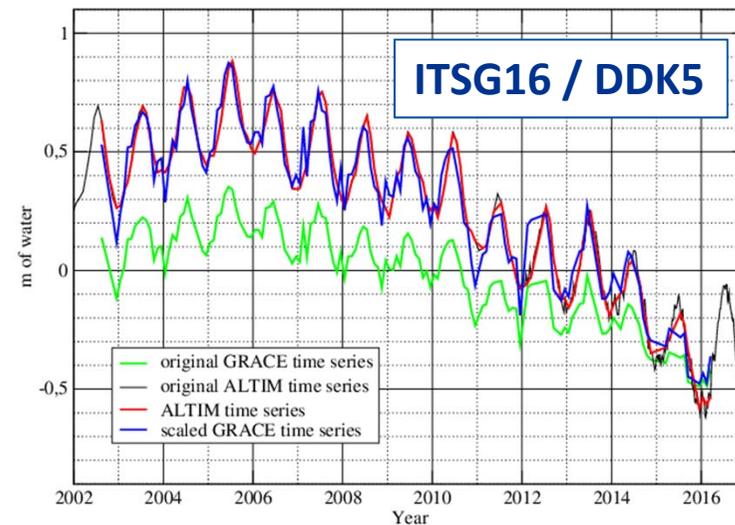
Caspian sea



Caspian sea

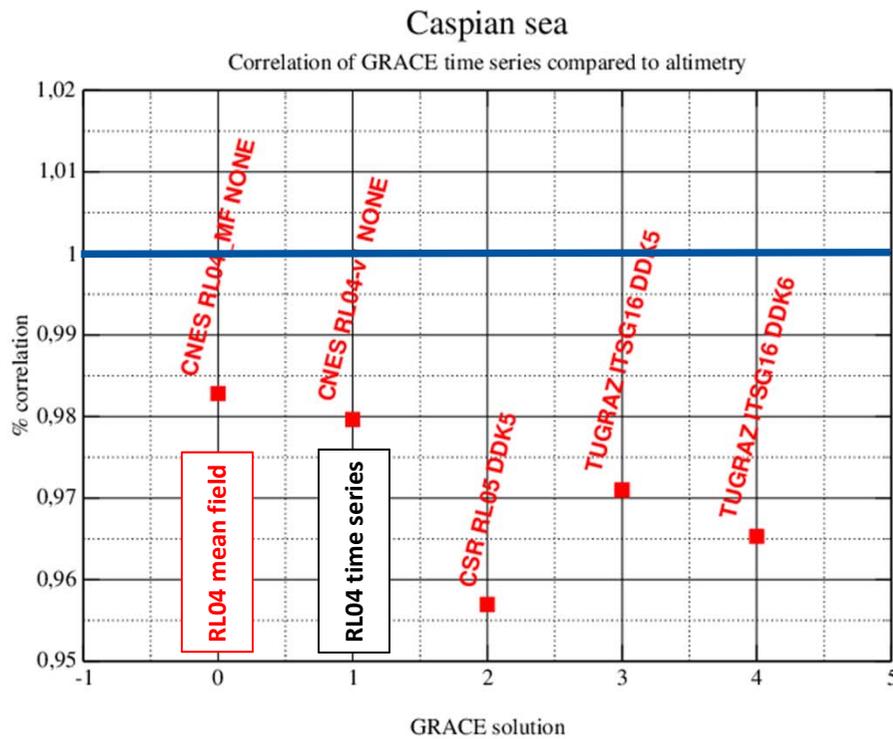


Altimeter time series from Hydroweb (<https://sso.theia-land.fr>)

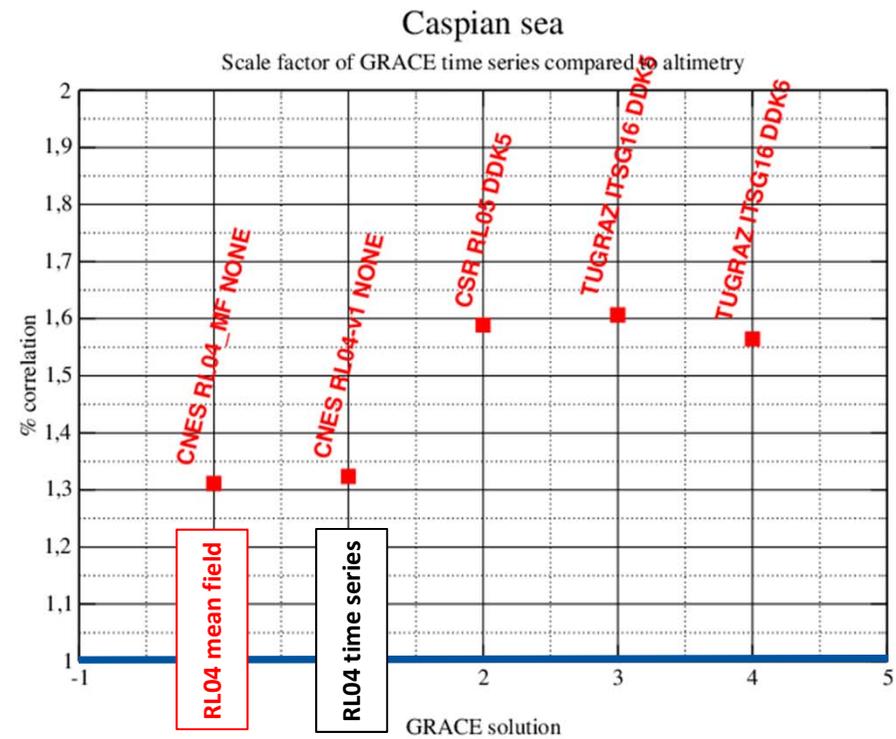


Signal assessment by comparison to altimetry (Caspian Sea)

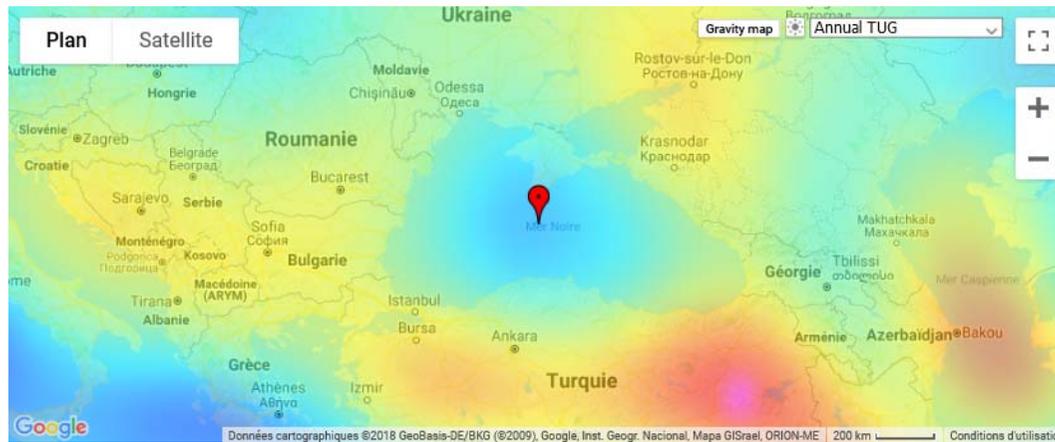
CORRELATION



SCALE FACTOR



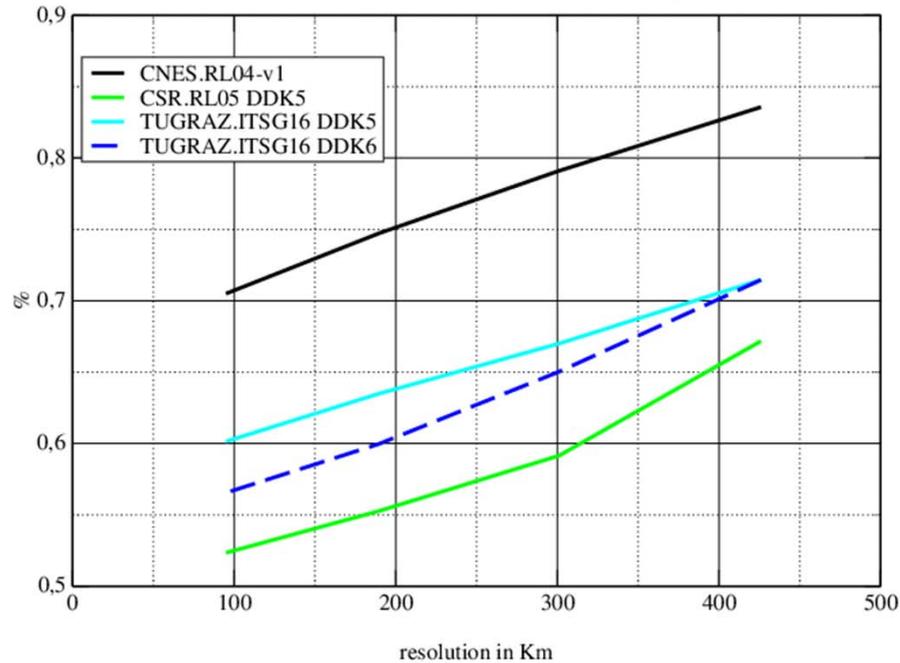
Signal assessment by comparison to altimetry (Black Sea)



Altimeter time series from Copernicus (<http://marine.copernicus.eu>).

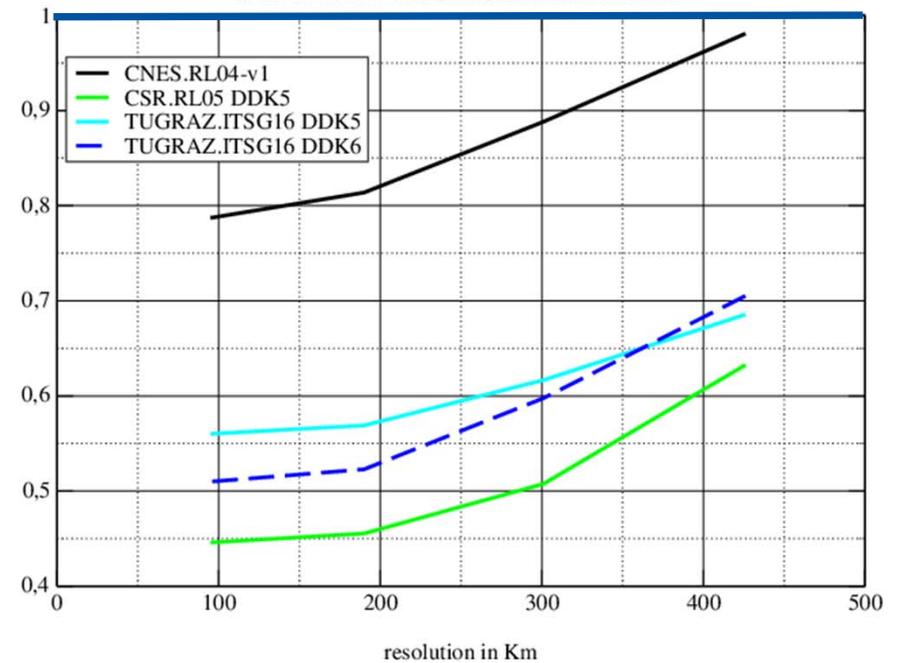
CORRELATION

Correlation between GRACE and ALTIMETRY, in %



SCALE FACTOR

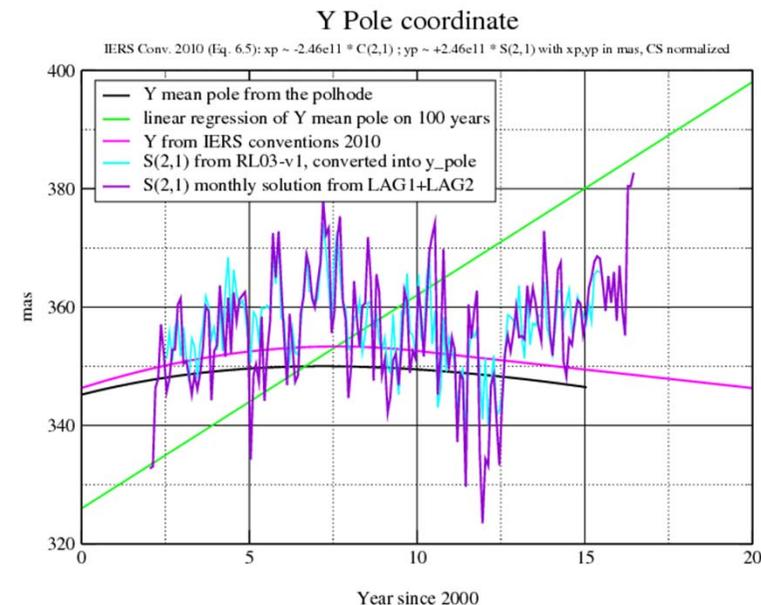
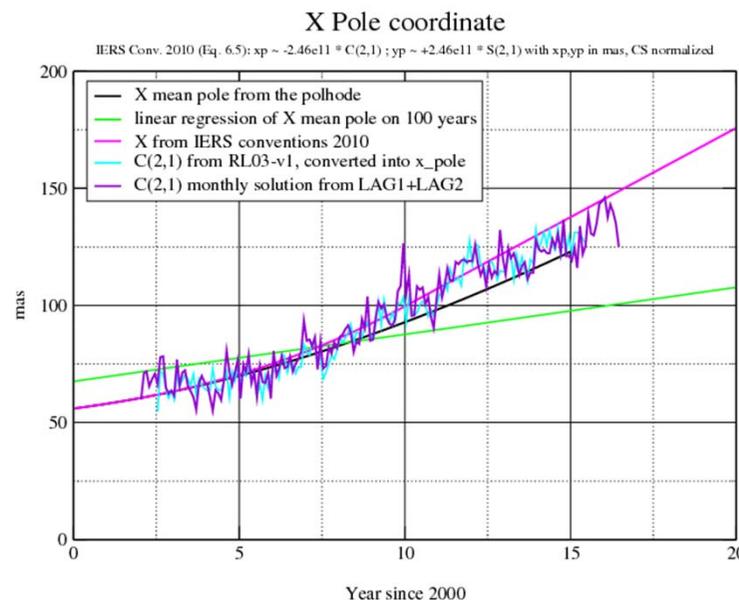
Scale factor between GRACE and ALTIMETRY



Need for consistency

MEAN POLE MODEL and GRAVITY FIELD

- When using the $C(2,1)/S(2,1)$ values of a gravity field model, one must adopt the **same mean pole convention** as the one used for the computation of the model.
- CNES/GRGS is using the mean pole of the **IERS2010 conventions**. If the conventions change for a **linear mean pole**, then the $C(2,1)/S(2,1)$ coefficients of the mean gravity model will have to be adapted to this new convention.



DEALIASING MODELS and GRAVITY FIELD

- The same goes for the **dealiasing models** : CNES uses 3-hour ERA-Interim & TUGO models → the same models should be used for POD

Conclusions and perspectives



- The new mean gravity field model based on CNES/GRGS RL04 is available for the GRACE period (2002 – 2016)
- Validation tests (noise and signal w.r.t. altimetry) show a good performance of this RL04 mean field
- Extrapolated periodic terms (before August 2002 and after May 2016) are based on global fits of monthly coefficients over 14 years of GRACE data
- It still needs to be completed before and after the GRACE period by additional data coming from SLR data (and DORIS data ?) in order to follow the long-term evolution of the lower degrees
- Possibly an accurate modeling of only the degree 2 through SLR and DORIS data is sufficient to achieve good POD performances
- When doing POD one must ensure that the models used are not only “good” but also consistent !!!
 - Consistency between the gravity field model and the HF dealiasing models
 - Consistency between the C(2,1)/S(2,1) time series and the mean pole model that was used to compute it.