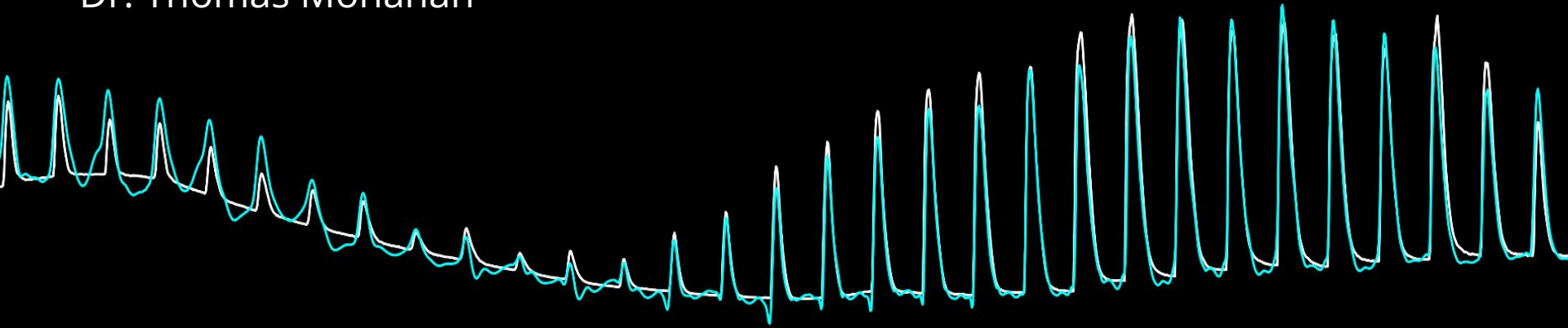


Progress towards a general altimetry derived tidal process model

Dr. Thomas Monahan



In collaboration with: Michael Hart-Davis, Jeffrey Polton, Stephen Roberts, and Thomas Adcock



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and Water Management

Eric and Wendy Schmidt

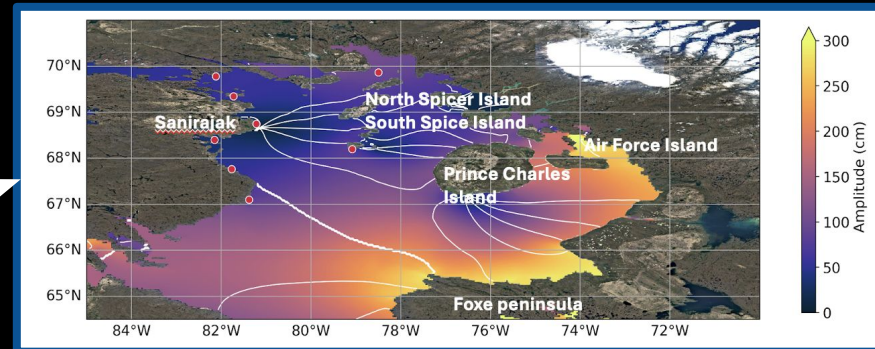
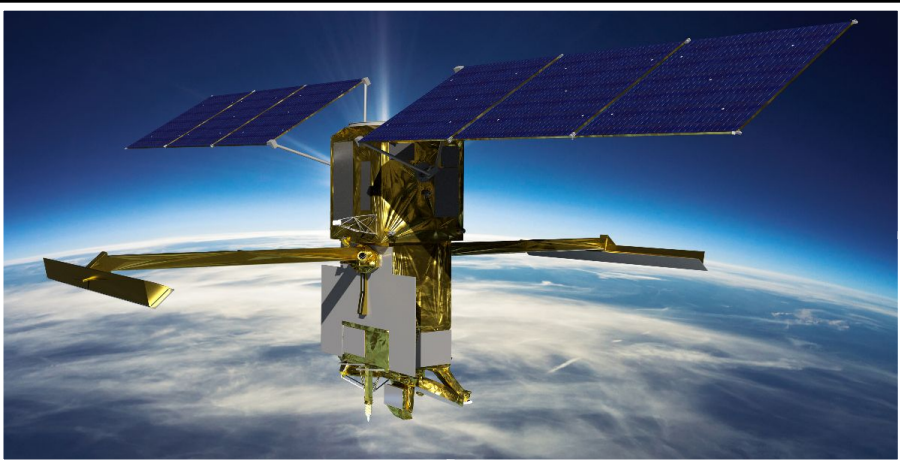
AI in Science Postdoctoral Fellowship,

a program of  SCHMIDT FUTURES

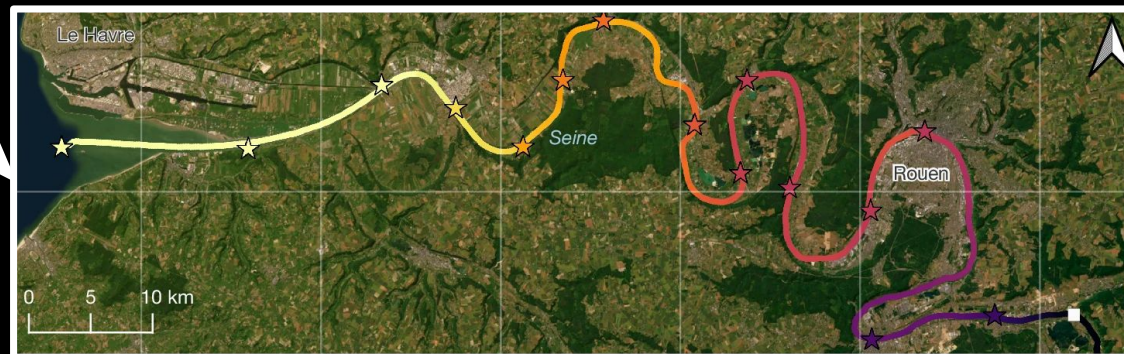
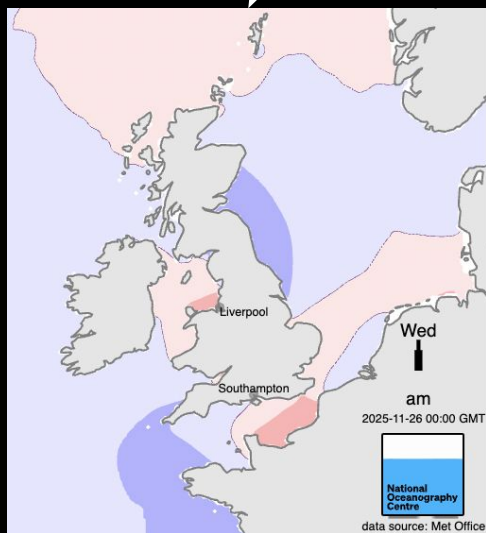


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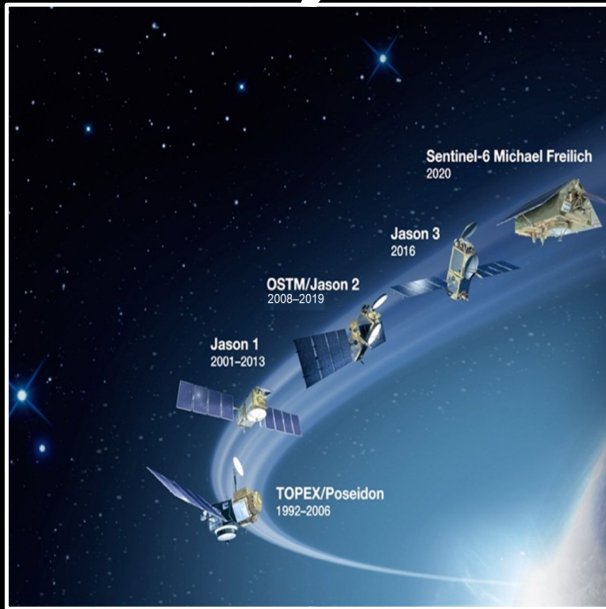


See our poster!



Hart-Davis, Michael G., et al. "Observing the tidal pulse of rivers from wide-swath satellite altimetry." *Nature* (2026): 1-8.

How do we use altimetry to improve our representation of high-frequency processes?



GOT

:

FES/DAC

EOT

Residual analysis

Existing Model

—

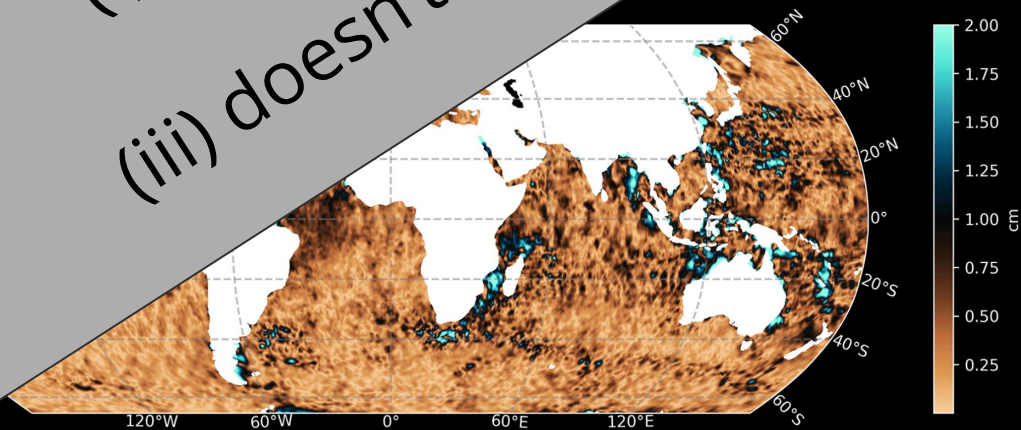
Observations

=

Final Forecast =

Existing Model

Challenges:
(i) Just applies to tides
(ii) Aliasing
(iii) doesn't use the model

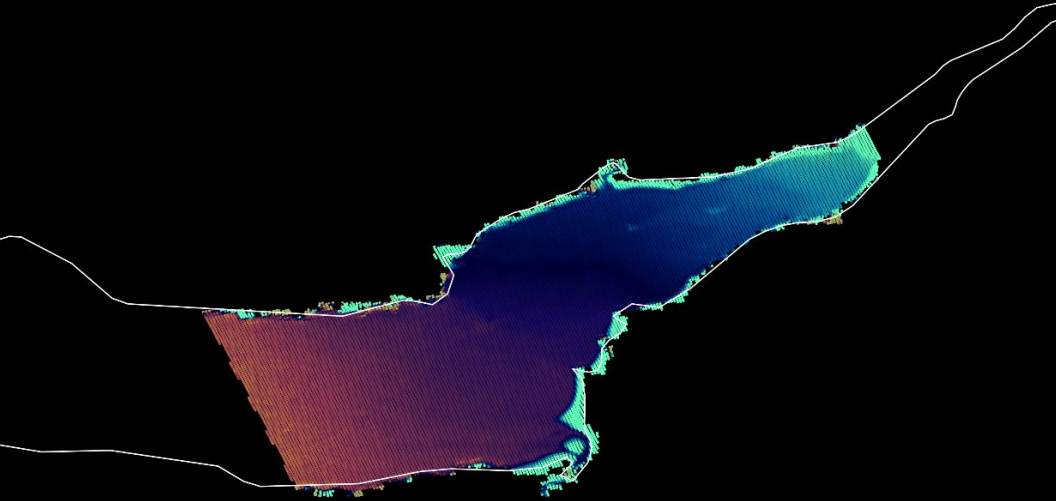


How can we use data to improve these models in challenging regions?

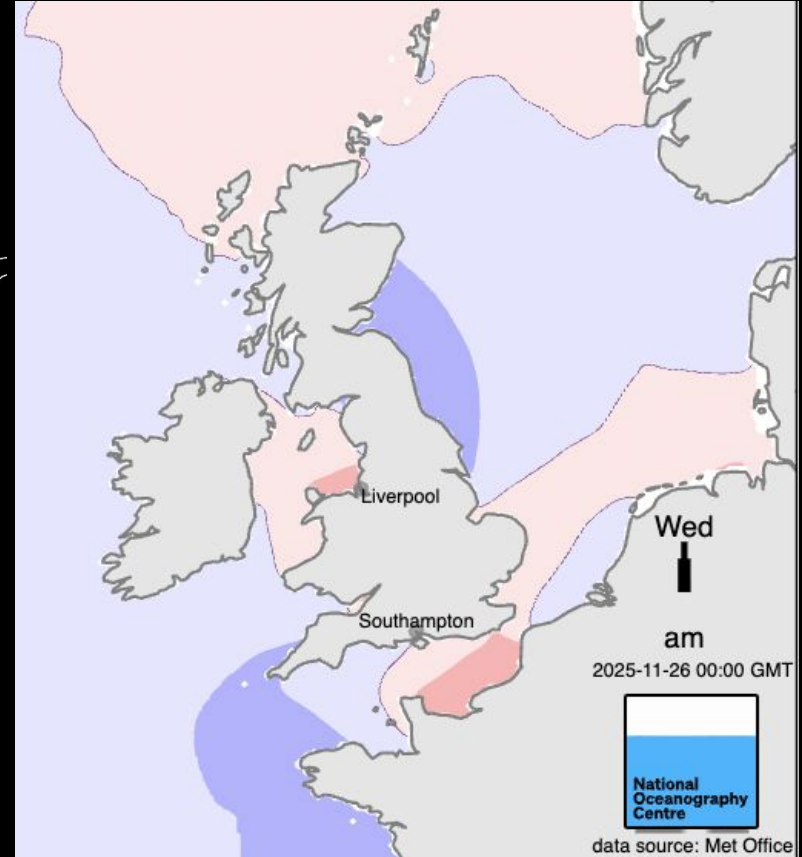
Observation:

Sea-levels smoothly vary between nearby locations.

$$\text{As } A \rightarrow B: \zeta_A \rightarrow \zeta_B$$



*ignoring hydraulic jumps



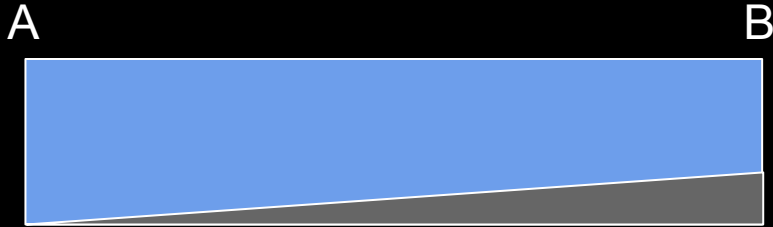
Claim:

IF, no exogenous forcing exists between stations, the sea-level at station **A** can be related to that of station **B** by delays and distortions of station A.

$$\zeta(x_B, t_n) \approx \sum_{\ell=-L}^L \alpha_{\ell} \zeta(x_A, t_n - \ell \Delta t)$$

Claim:

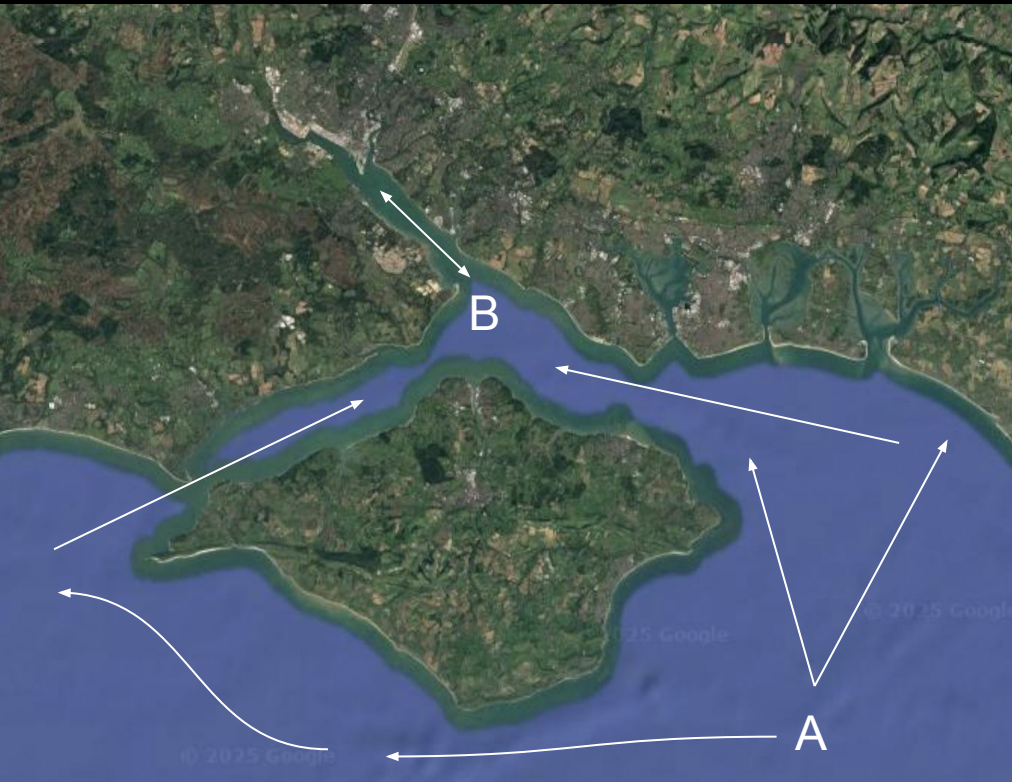
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$$\zeta(x_B, t_n) \approx \alpha_\ell \zeta(x_A, t_n - \ell \Delta t)$$

Claim:

IF, no exogenous forcing exists between stations, the sea-level at station **A** can be related to that of station **B** by delays and distortions of station A.

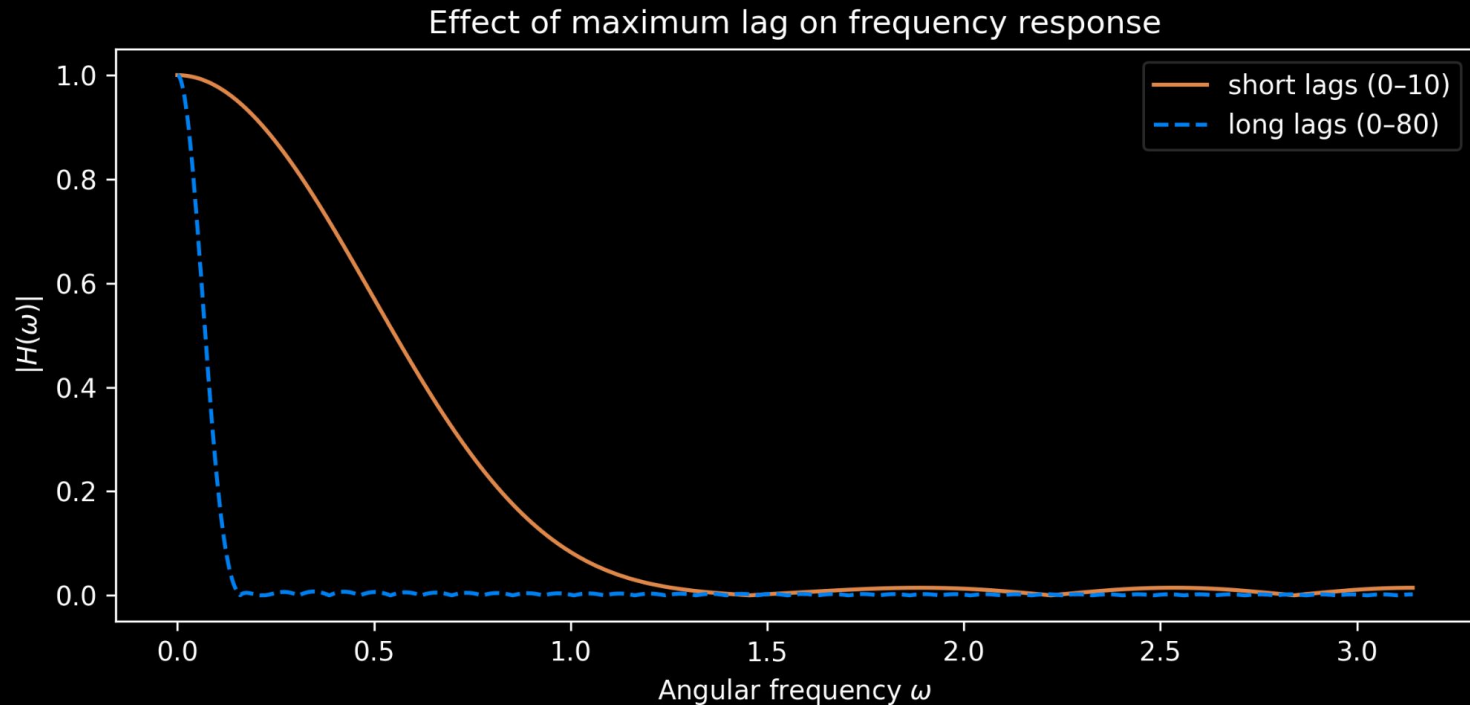


$$\zeta(x_B, t_n) \approx \sum_{\ell=-L}^L \alpha_{\ell} \zeta(x_A, t_n - \ell \Delta t)$$

Matte, Pascal, and Silvia Innocenti.
"Filling gaps and extending inland tide
gauge records to support accurate
long-term and extreme water level
analyses." *Continental Shelf Research*
(2026): 105668.

A frequency interpretation

$$H(\omega) = \sum_{\ell \in \mathcal{L}} \alpha_{\ell} e^{-i\omega \ell \Delta t}$$



But linear kernels aren't always enough...

Bed friction

$$\tau_b = C_d |u| u \sim u^2$$

Primary source of overtides
in rivers and estuaries

Depth–transport coupling

$$\frac{\partial}{\partial x}(\eta u)$$

Elevation $\times$ velocity:
tide–surge interaction

Depth–wind modulation

$$\frac{\tau_w}{\rho(h + \eta)} \approx \frac{\tau_w}{\rho h} \left(1 - \frac{\eta}{h}\right)$$

Surge modulates the
effective wind response

A **linear** kernel can only redistribute existing frequencies.

The **bilinear** kernel

$$\sum_{\tau_1, \tau_2} w_2(\tau_1, \tau_2) \hat{\zeta}(t - \tau_1) \hat{\zeta}(t - \tau_2)$$

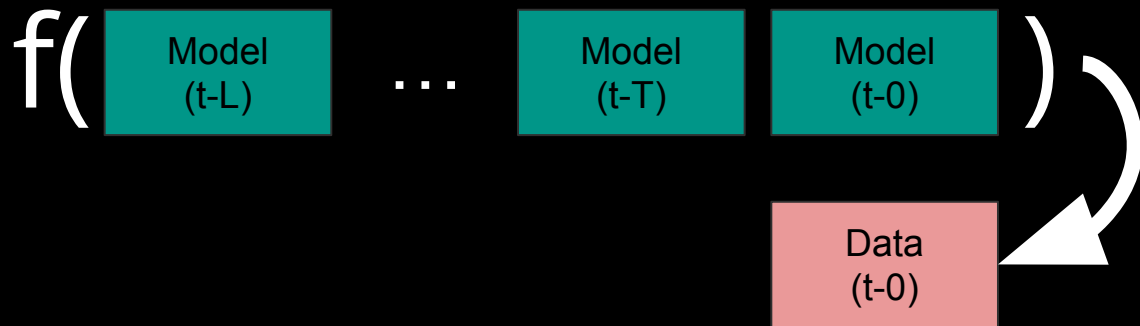
generates new harmonics

— it is the data-driven analogue of
weakly nonlinear shallow-water theory.

A time-invariant framework for post-processing shallow water models

$$\zeta(t) \approx b + \sum_{\tau=0}^L w_1(\tau) \hat{\zeta}(t - \tau) + \sum_{\tau_1=0}^L \sum_{\tau_2=0}^L w_2(\tau_1, \tau_2) \hat{\zeta}(t - \tau_1) \hat{\zeta}(t - \tau_2),$$

Recipe:

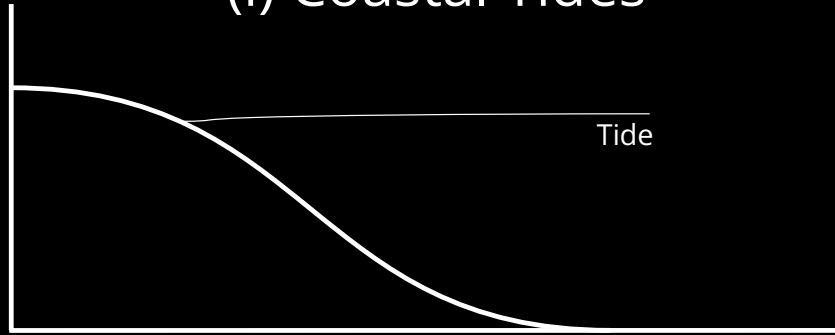


**Time-invariant
postprocessing**

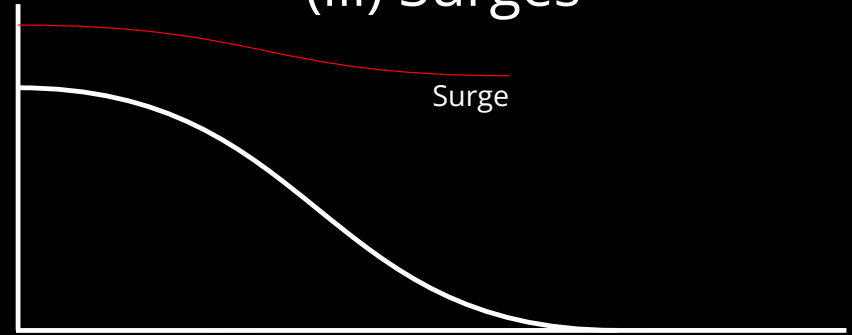


3 contexts.

(i) Coastal Tides



(iii) Surges



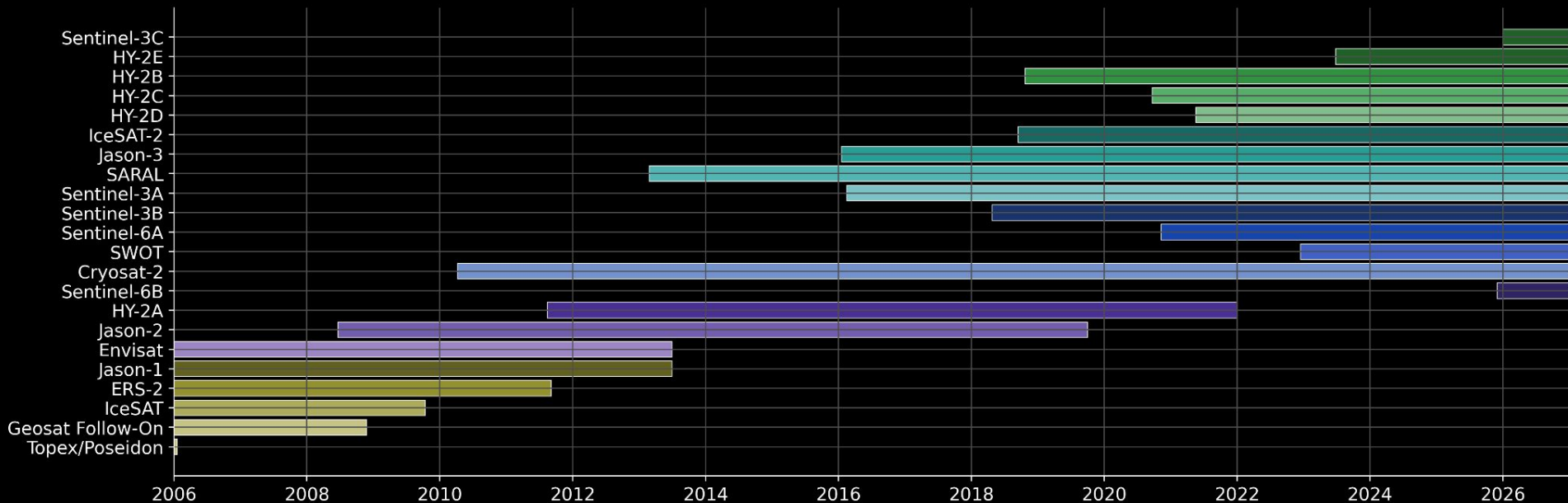
(ii) Tidal Rivers





OpenADB

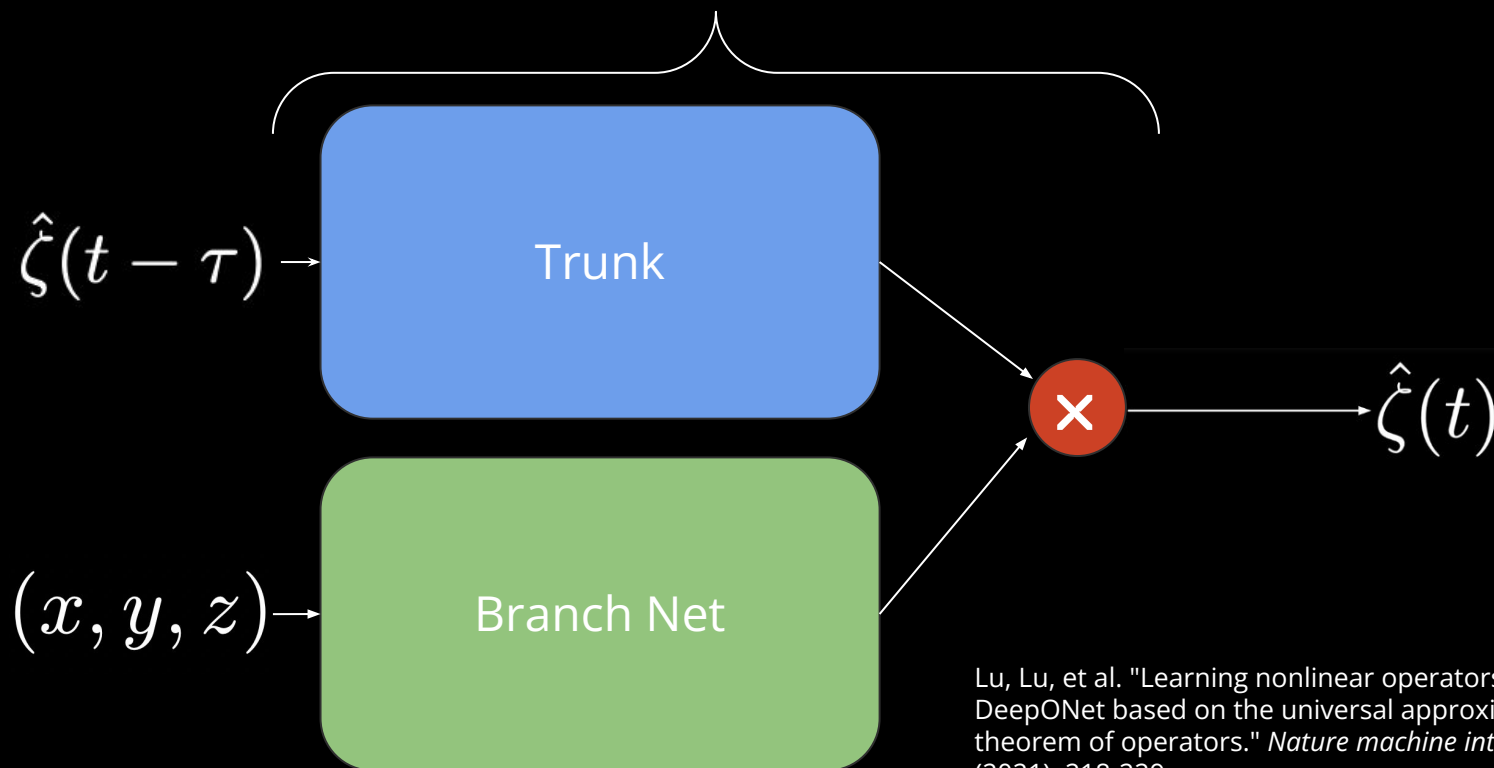
Questions (Ask Mike!)



$$\zeta(x, y, t) \approx b + \sum_{\tau=0}^{L^{(1)}} w_1(x, y, \tau) \hat{\zeta}(t - \tau) + \sum_{\tau_1=0}^{L^{(2)}} \sum_{\tau_2=\tau_1}^{L^{(2)}} w_2(x, y, \tau_1, \tau_2) \hat{\zeta}(t - \tau_1) \hat{\zeta}(t - \tau_2)$$

A spatially coherent response operator

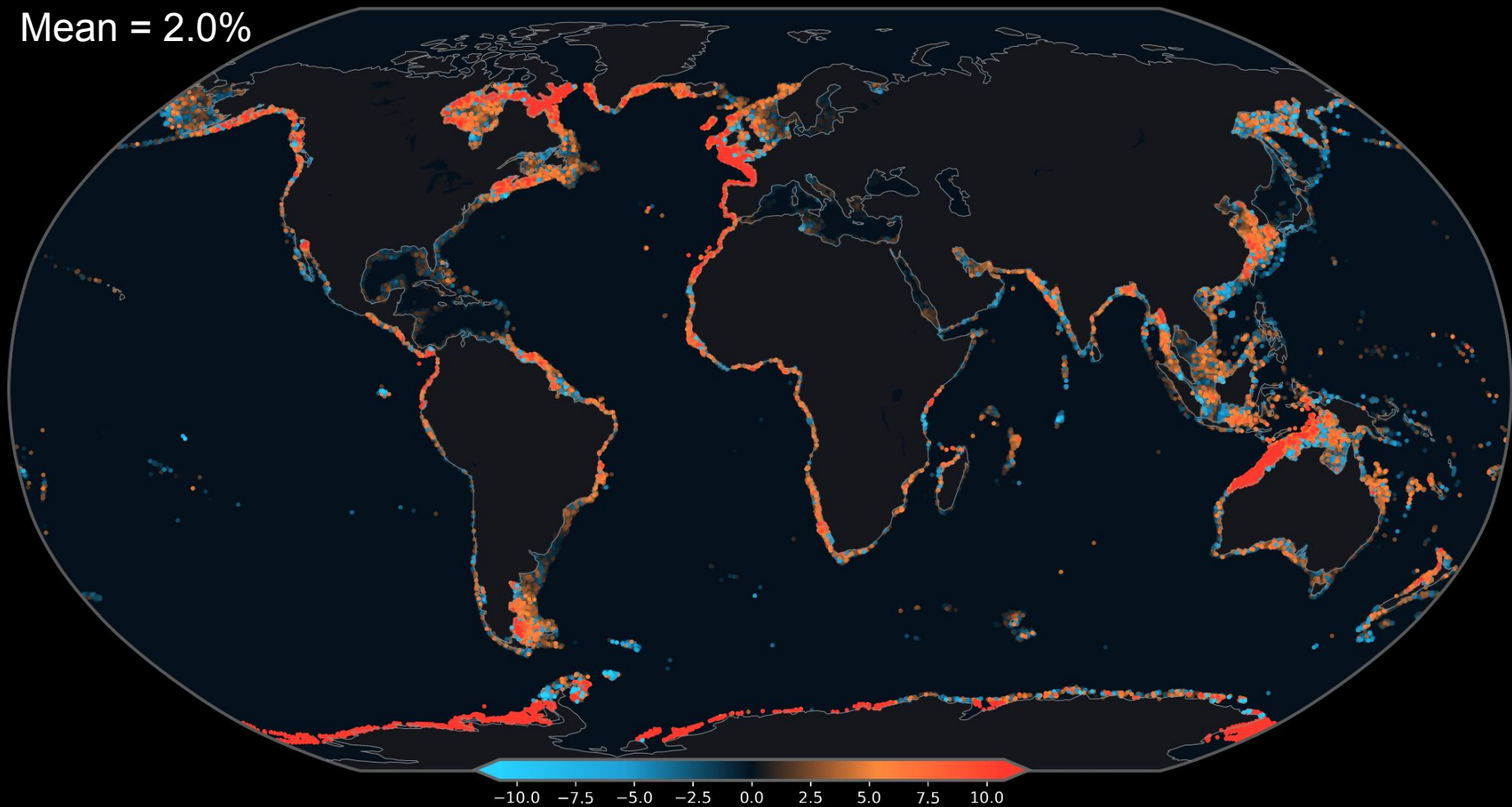
$$\zeta(x, y, t) \approx b + \sum_{\tau=0}^{L^{(1)}} w_1(x, y, \tau) \hat{\zeta}(t - \tau) + \sum_{\tau_1=0}^{L^{(2)}} \sum_{\tau_2=\tau_1}^{L^{(2)}} w_2(x, y, \tau_1, \tau_2) \hat{\zeta}(t - \tau_1) \hat{\zeta}(t - \tau_2)$$



Lu, Lu, et al. "Learning nonlinear operators via DeepONet based on the universal approximation theorem of operators." *Nature machine intelligence* 3.3 (2021): 218-229.

GOAT(FES) Variance Reduction

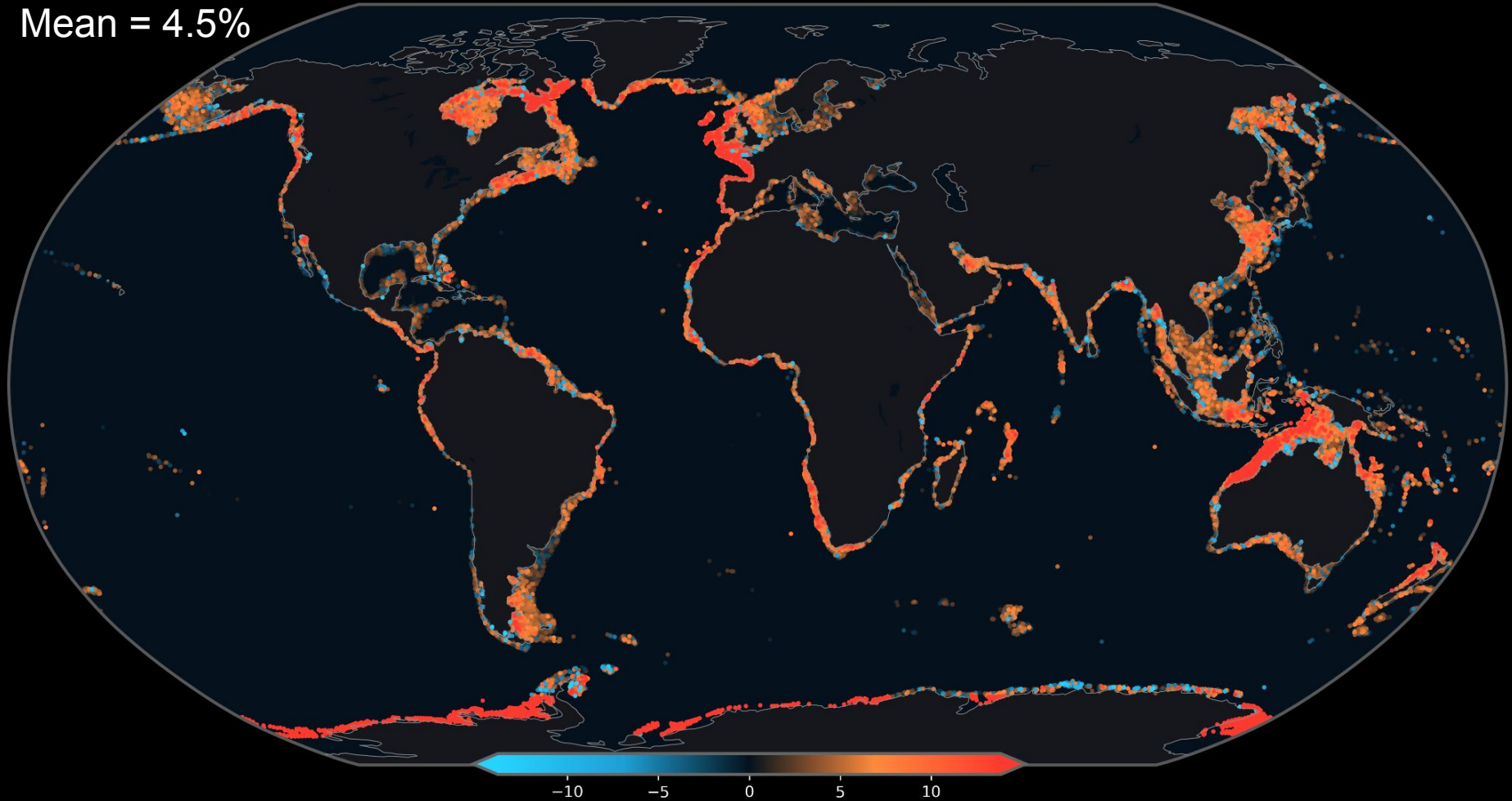
Mean = 2.0%



% Variance Reduction: red = better, blue = worse.

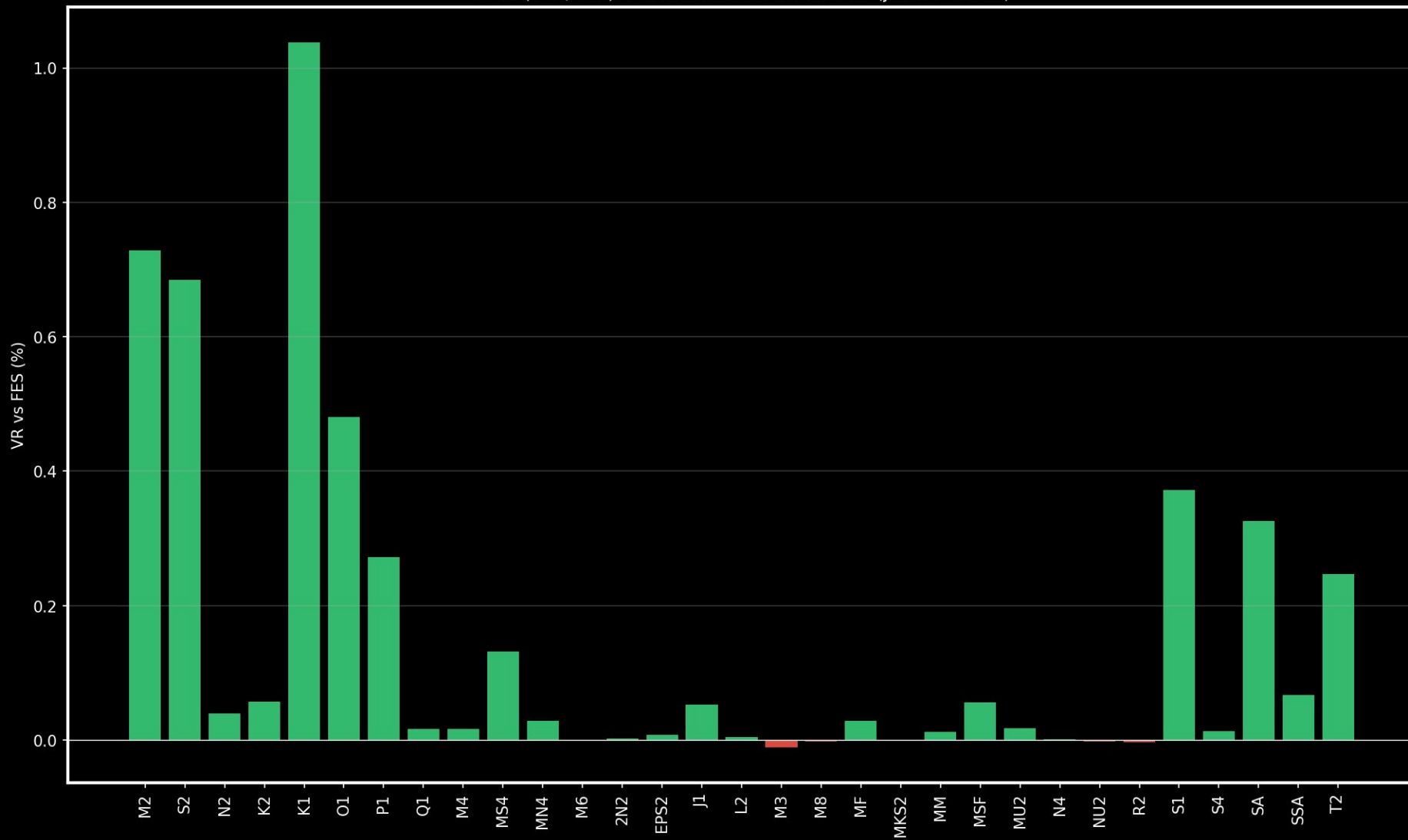
GOAT(FES +EOT20) Variance Reduction

Mean = 4.5%



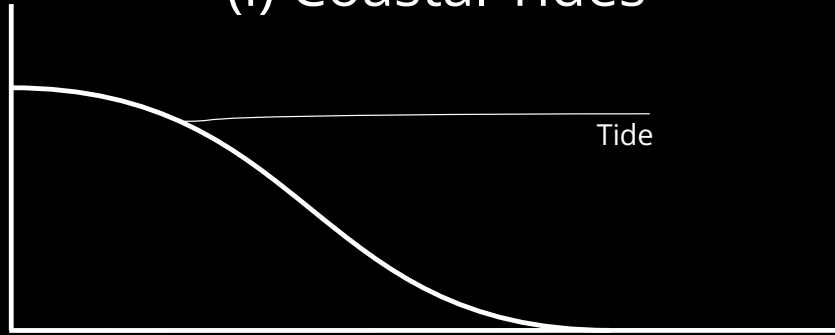
% Variance Reduction: red = better, blue = worse.

GOAT(FES,EOT): Per-constituent VR vs FES (joint +4.49%)

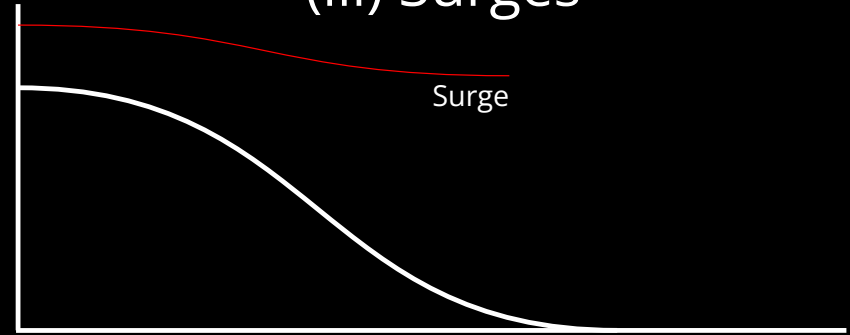


3 contexts.

(i) Coastal Tides



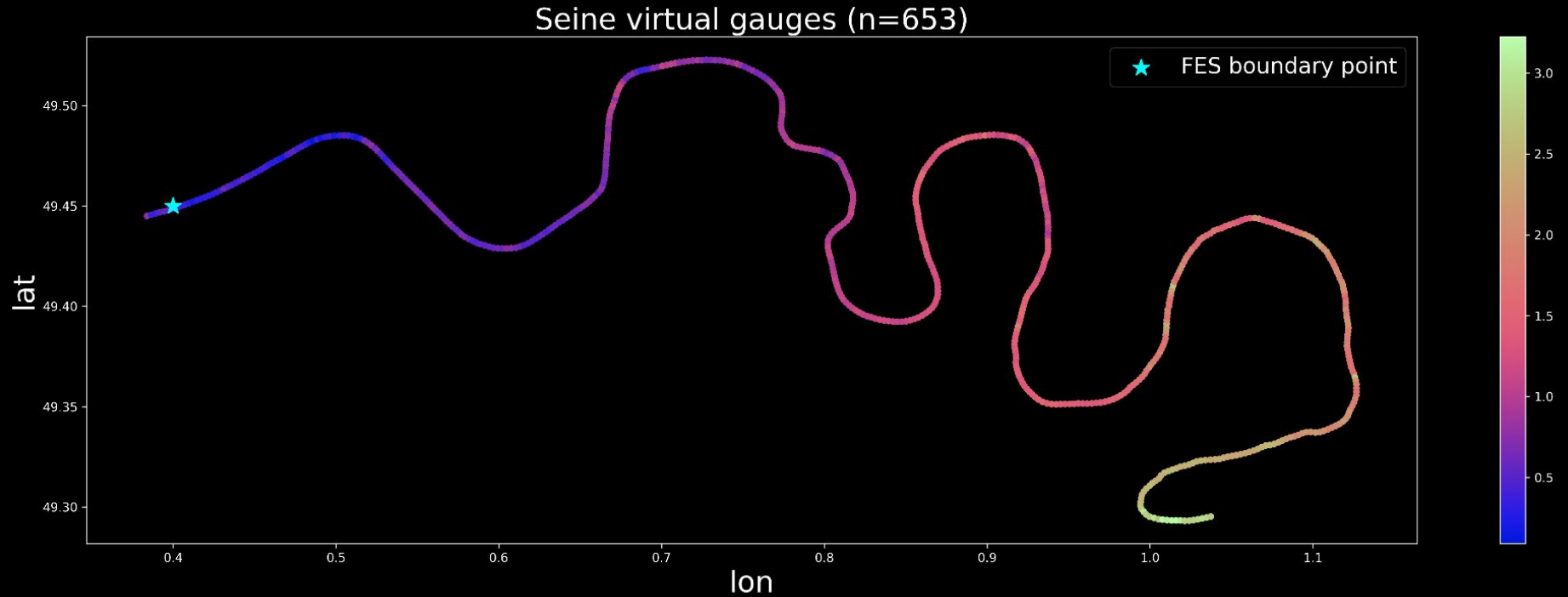
(iii) Surges



(ii) Tidal Rivers

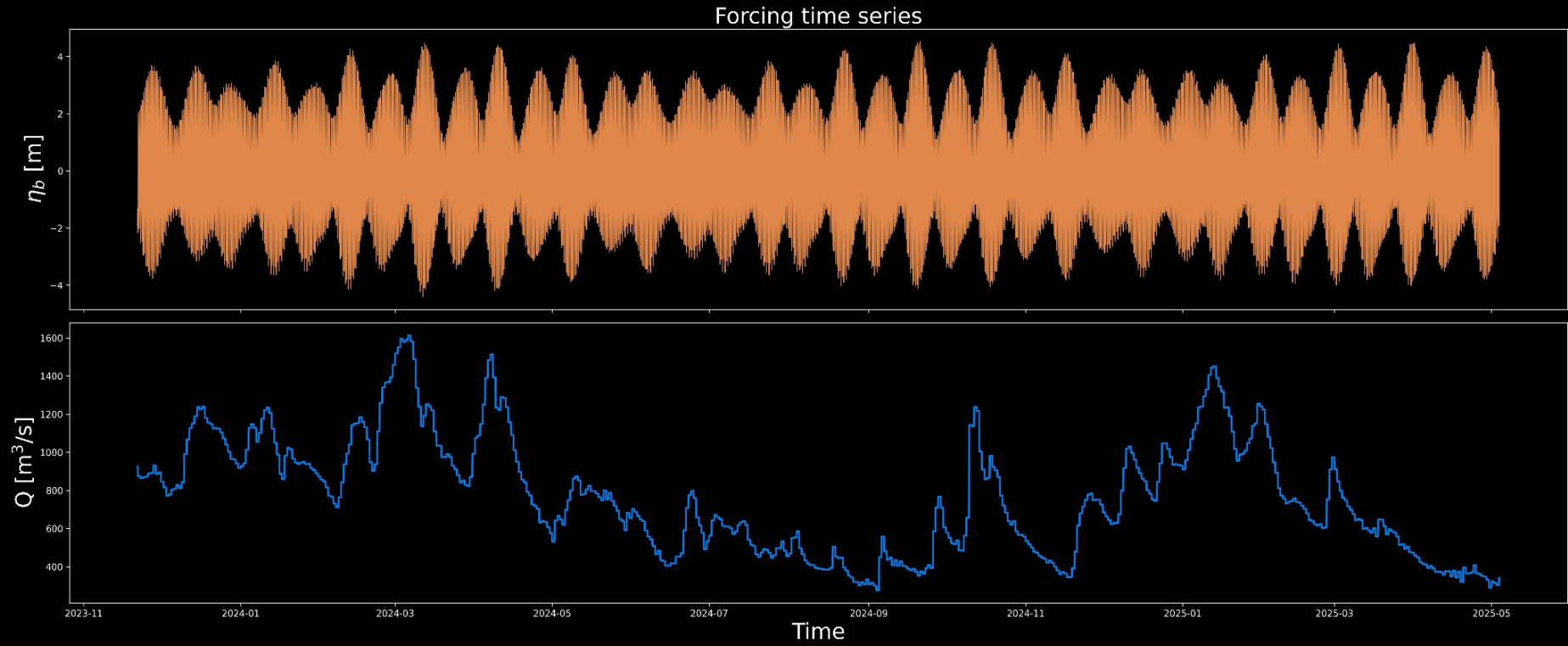


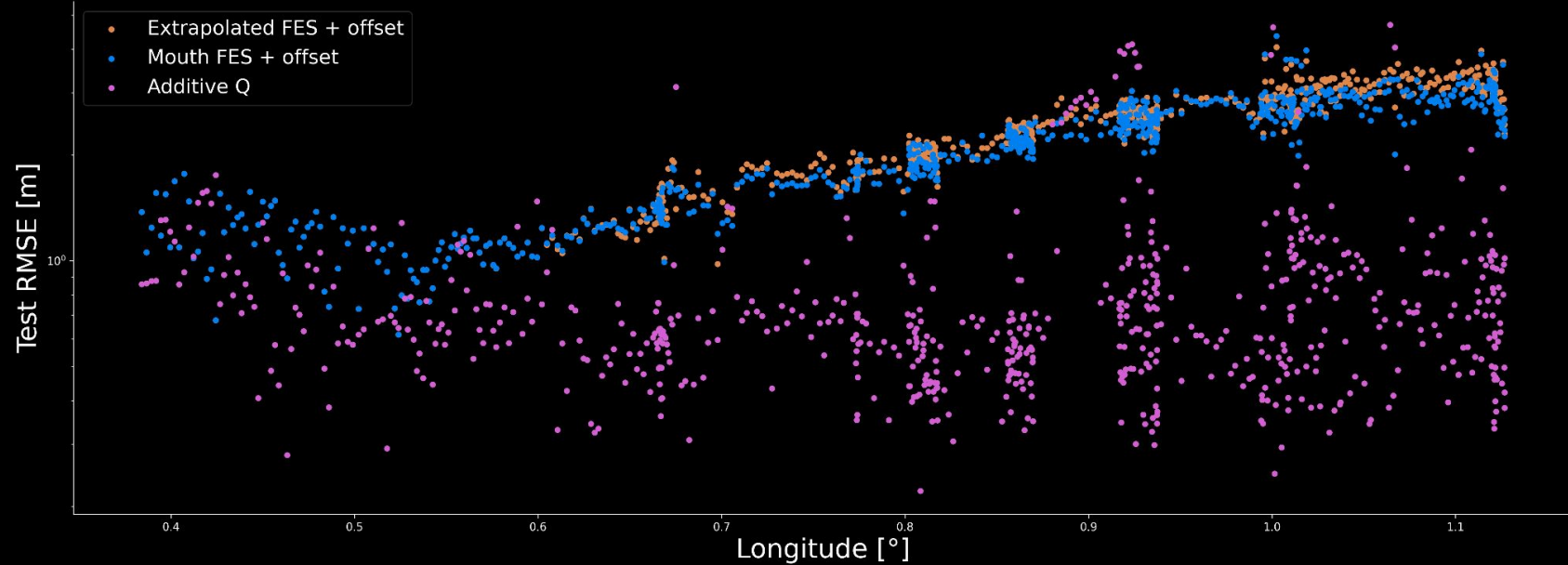
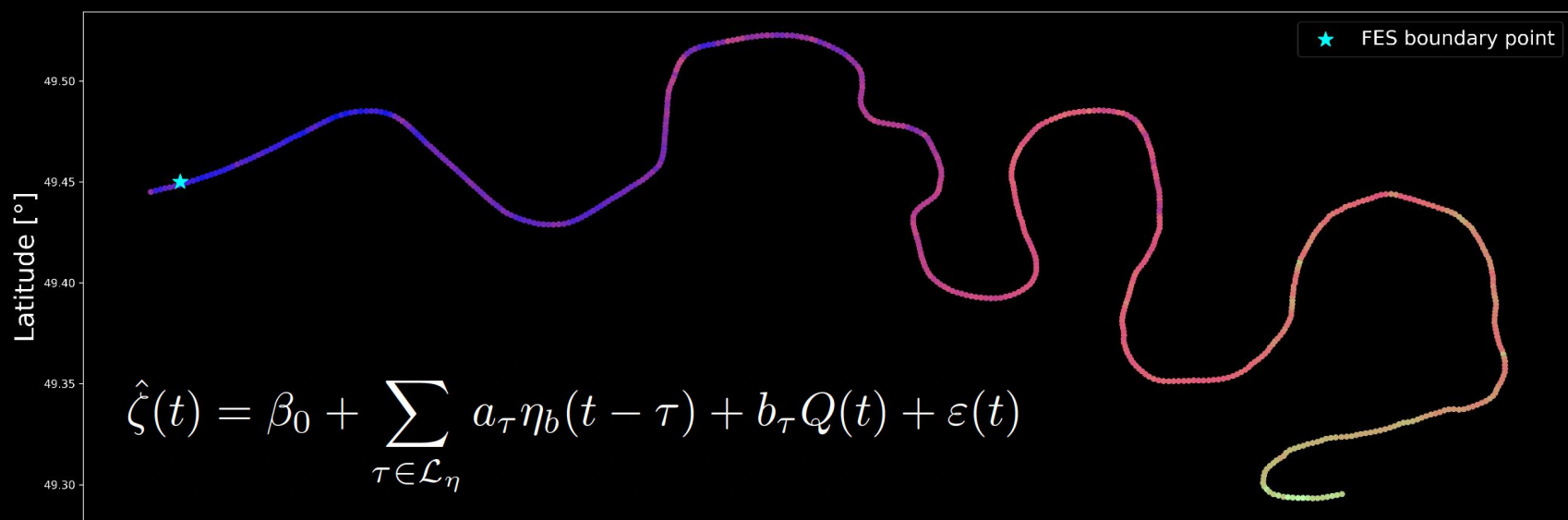
Extending global tide models up tidal rivers



$$\hat{\zeta}(t) = \beta_0 + \sum_{\tau \in \mathcal{L}_\eta} a_\tau \eta_b(t - \tau) + b_\tau Q(t) + \varepsilon(t)$$

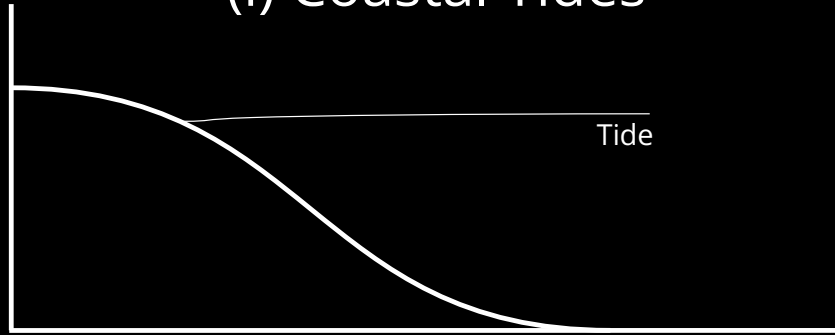
Tidal Rivers



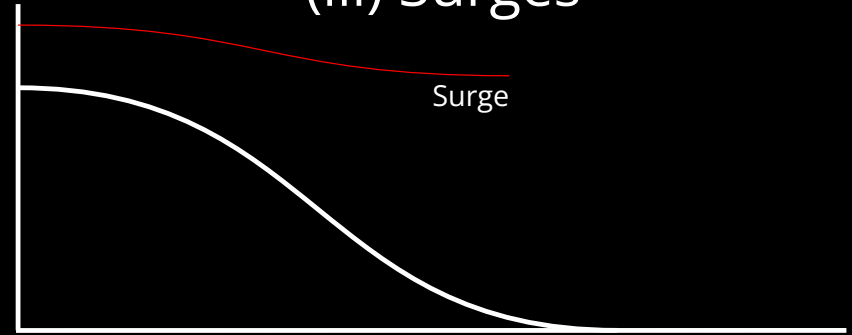


3 contexts.

(i) Coastal Tides



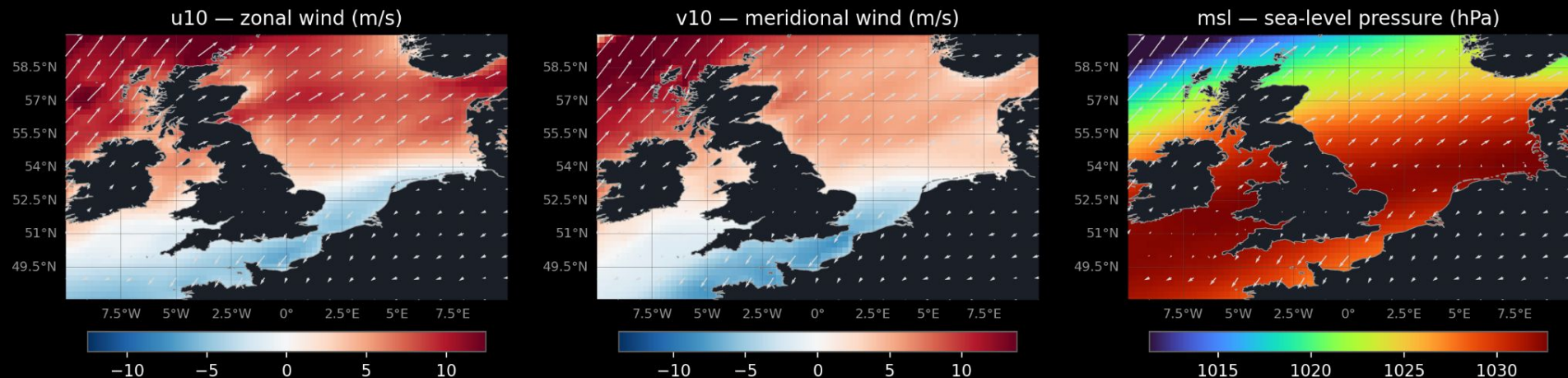
(iii) Surges



(ii) Tidal Rivers



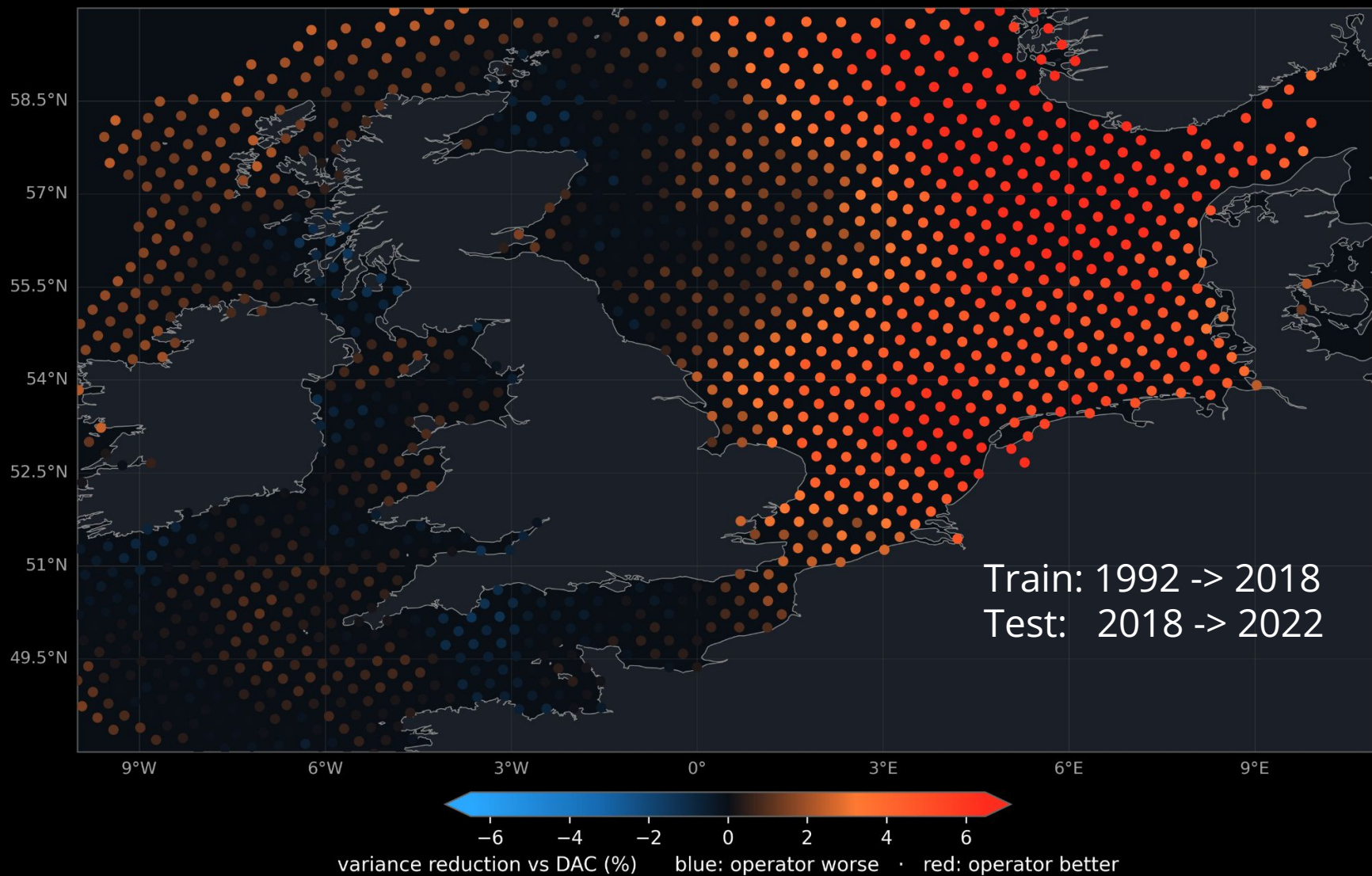
DAC



$$\begin{aligned}\hat{\zeta}(t) = & \sum_{\tau} \alpha_u(\tau) \cdot u(t - \tau) + \sum_{\tau} \alpha_v(\tau) \cdot v(t - \tau) + \sum_{\tau} \alpha_{\text{msl}}(\tau) \cdot \text{MSL}(t - \tau) \\ & + \sum_{\tau} \beta_u(\tau) \cdot \bar{u}_{5^\circ}(t - \tau) + \sum_{\tau} \beta_v(\tau) \cdot \bar{v}_{5^\circ}(t - \tau) + \sum_{\tau} \beta_{\text{msl}}(\tau) \cdot \bar{\text{MSL}}_{5^\circ}(t - \tau)\end{aligned}$$

Atmospheric operator vs DAC — out-of-sample variance reduction

mean +2.07% · median +1.36% · 83% of 1,485 sites improved



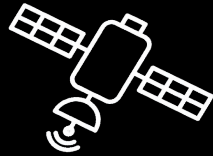
Learned admittances are interpretable.



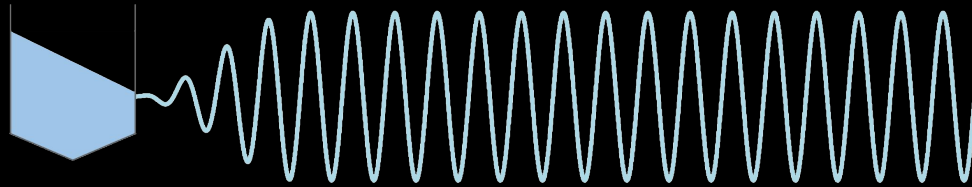
Conclusions.

- Broadband response operators provide an interesting way forward in **augmenting** existing global tide models for hard problems.
- Optimal combinations of global tide models should be considered further
- These operators when modified to allow for river modulation can **enable existing global tide models to extend up tidal rivers.**
 - The broadband operators can be learned from sparse SWOT data.
- Response type approaches may be useful for **identifying**, and in some case **learning**, deficiencies of surge models.

Special thanks to collaborators, partners, and funders... Especially High-freq!



Questions?



thomas.monahan@eng.ox.ac.uk



NOVELTIS



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**Ministry of Infrastructure
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Eric and Wendy Schmidt

AI in Science Postdoctoral Fellowship,

a program of  **SCHMIDT FUTURES**



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OXFORD**

How do we estimate w ?

$$\zeta = \mathbf{w}^T X + \epsilon \quad \longleftarrow \text{Linear Regression!}$$

$$\zeta = [\zeta(t_1), \zeta(t_2), \dots, \zeta(t_N)]$$

$$X = [\hat{\zeta}(t_i - \tau), \dots, \hat{\zeta}(t - L), \hat{\zeta}(t_i - \tau_1) \hat{\zeta}(t_i - \tau_2), \dots]^T$$

Monahan, T., Tang, T., Roberts, S., & Adcock, T. A. A. (2025). Tidal corrections from and for SWOT using a spatially coherent variational Bayesian harmonic analysis. *Journal of Geophysical Research: Oceans*, 130, e2024JC021533. <https://doi.org/10.1029/2024JC021533>

Why not higher order?

$$\zeta(t) \approx b + \sum_{n=1}^P \sum_{\tau_1, \dots, \tau_n=0}^L w_n(\tau_1, \dots, \tau_n) \prod_{j=1}^n \hat{\zeta}(t - \tau_j),$$

Classical response method:

Constructs narrow-band filters over tidal species.

M6 = trilinear product M2 X M2 X M2

Lag-based approach

Constructs broad-band frequencies across entire tidal spectrum.

M6 = product of M4 X M2? Or even redistribution of M6 already in model.