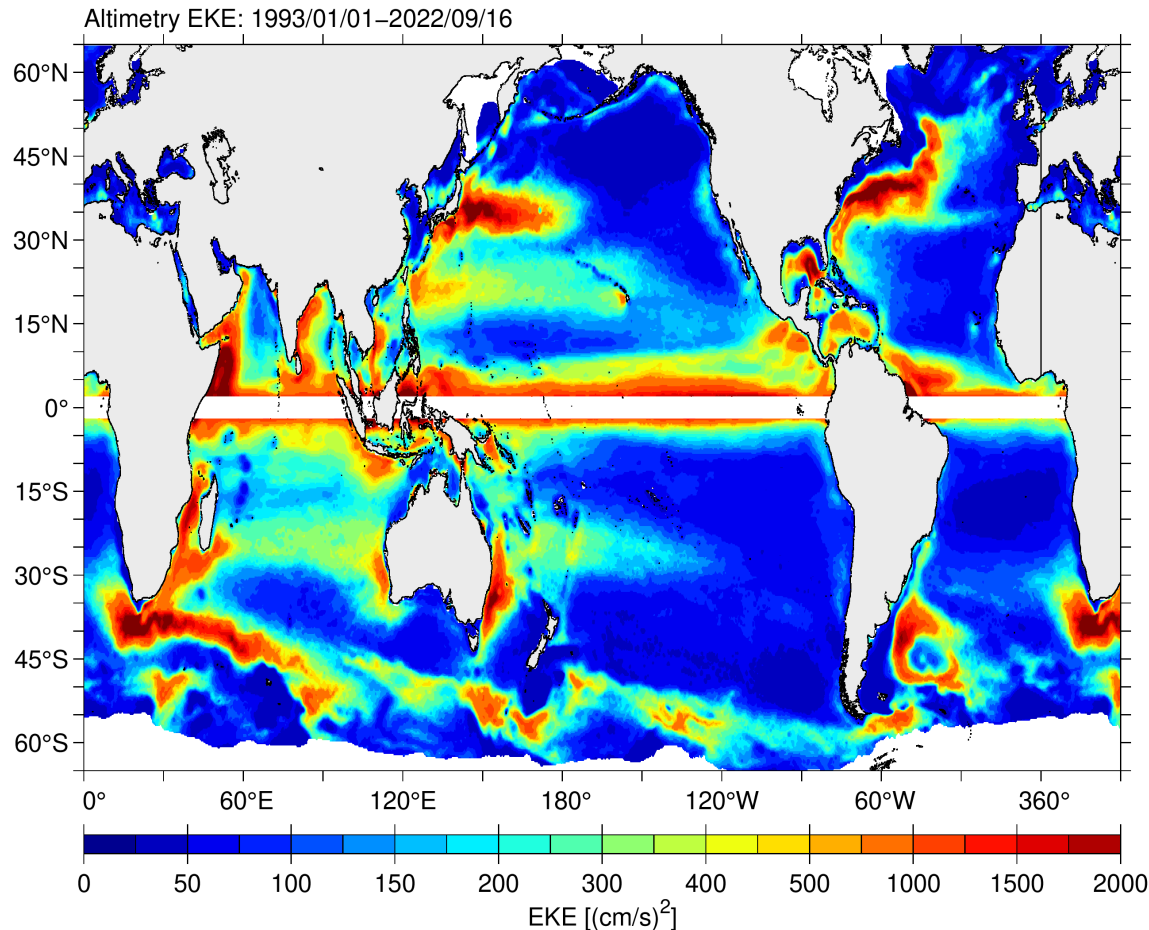


# Surface ocean eddy kinetic energy trend from 3-decade satellite altimetry missions: Observations & causes

**Bo Qiu & Shuiming Chen**

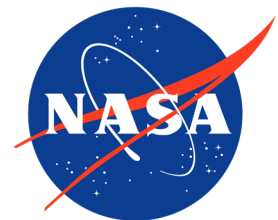
Department of Oceanography, University of Hawaii at Manoa



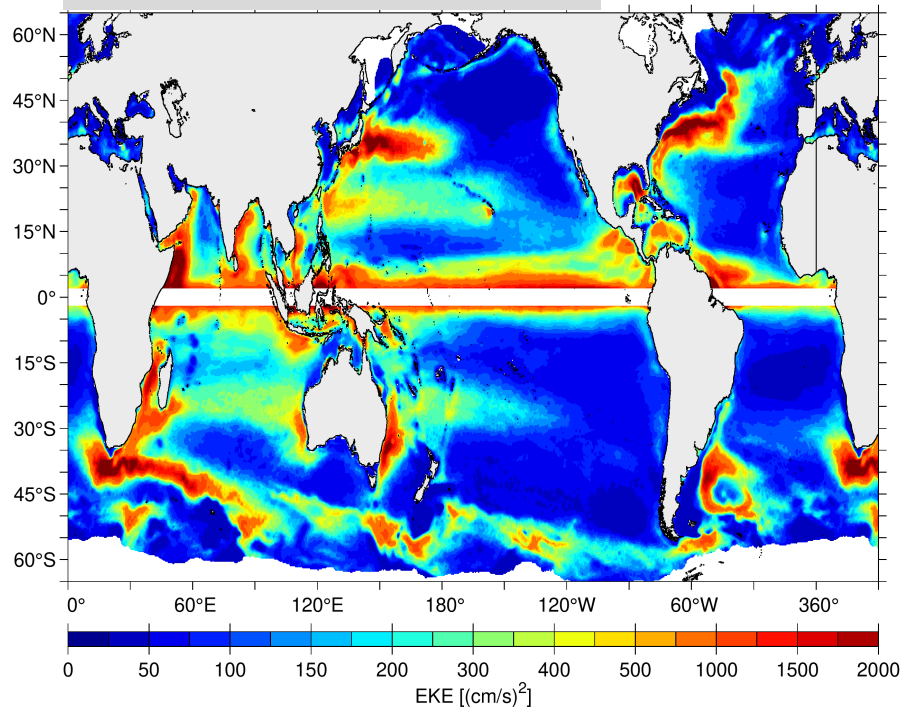
Altimeter-derived EKE  
map in global ocean  
from 1993–2022



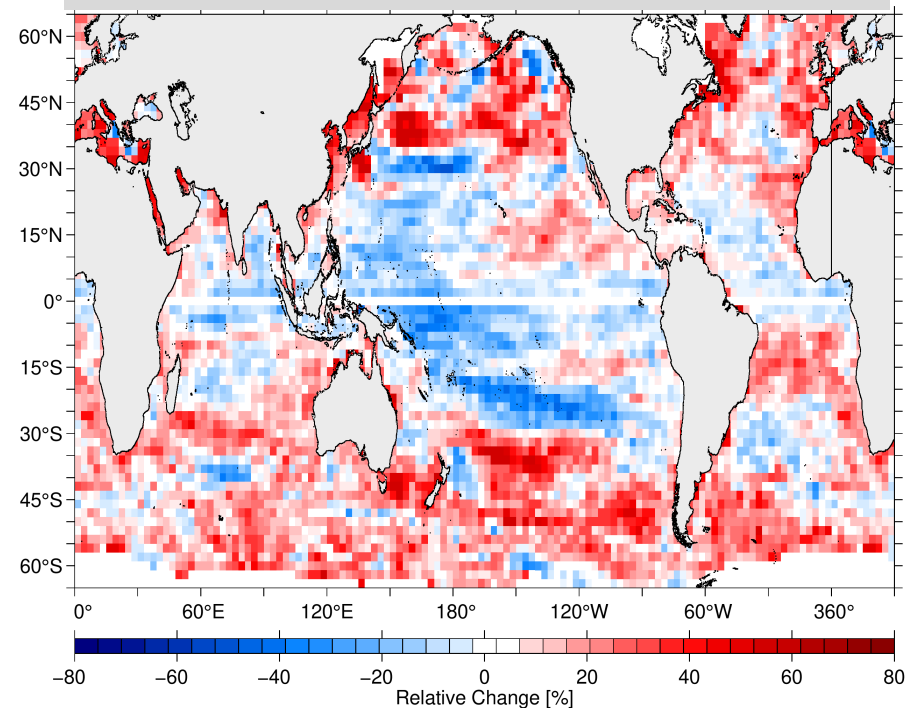
OSTST PI Meeting, Venice, 3 November 2022



Mean EKE map from 1993–2022

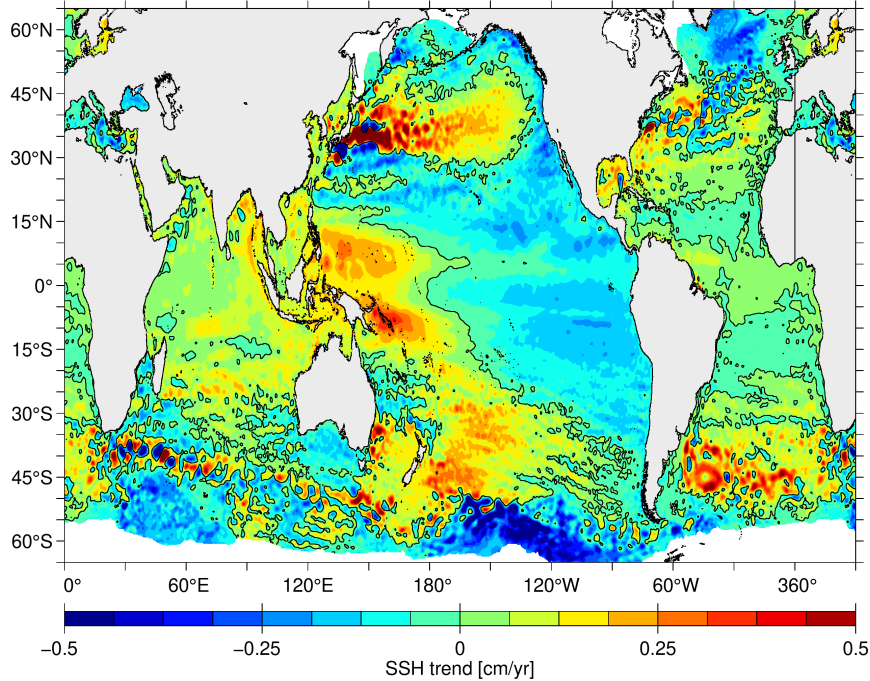


EKE change relative to local mean from 1993–2022

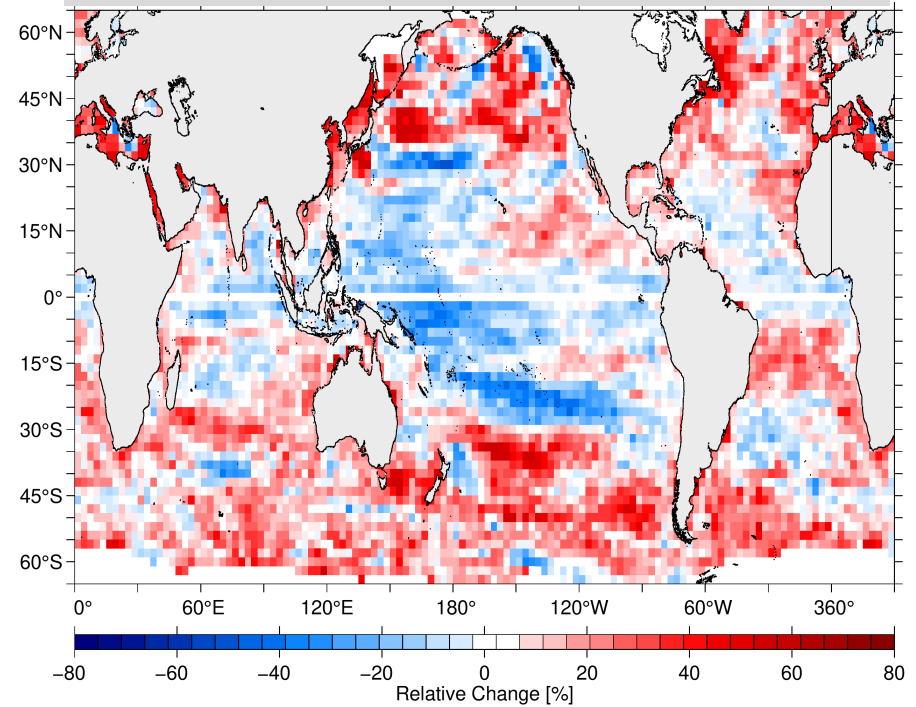


- Surface ocean EKE distribution has been established since beginning of the satellite altimetry era
- 3-decade accumulation of altimetry data allows us now to examine EKE linear trends; right map shows **relative EKE change** over the past 30-year period
- Increasing trends are seen in most of the **interior quiescent world oceans**
- Decreasing trends appear in equatorial & western basins of tropical/subtropical gyres

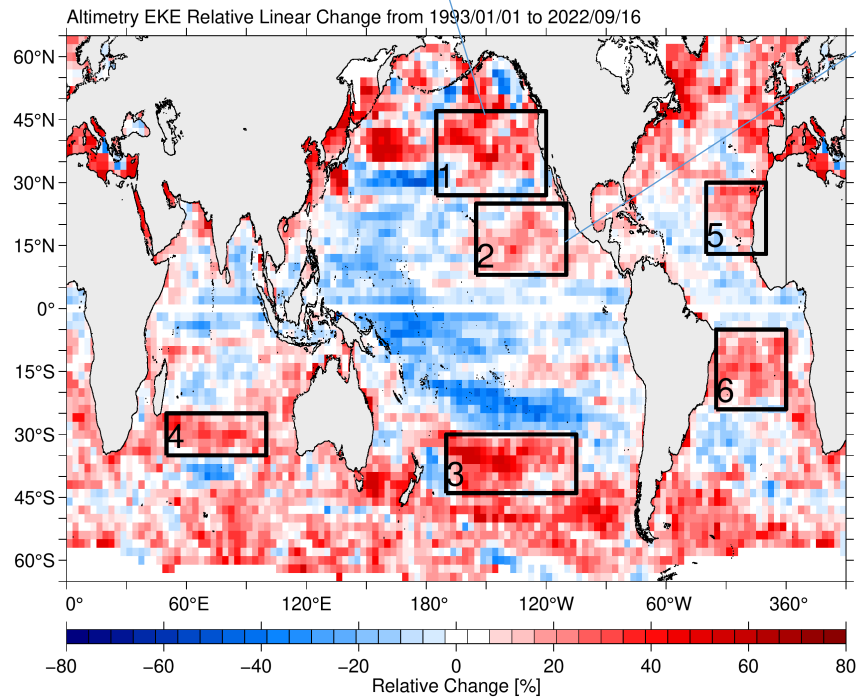
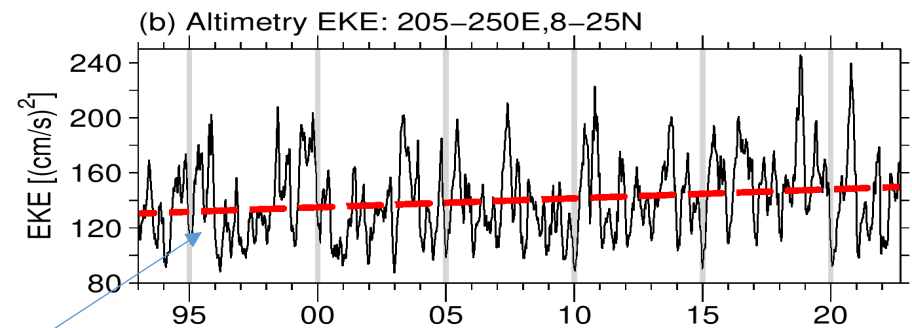
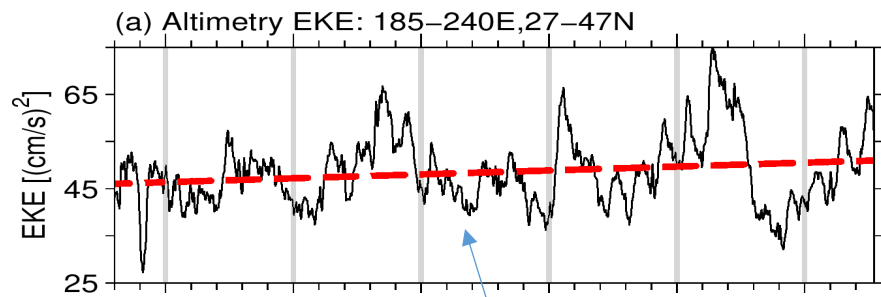
Sea level rise trend (no global mean) from 1993–2022



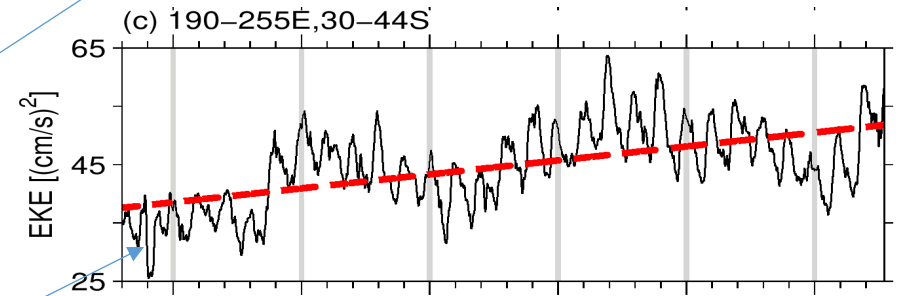
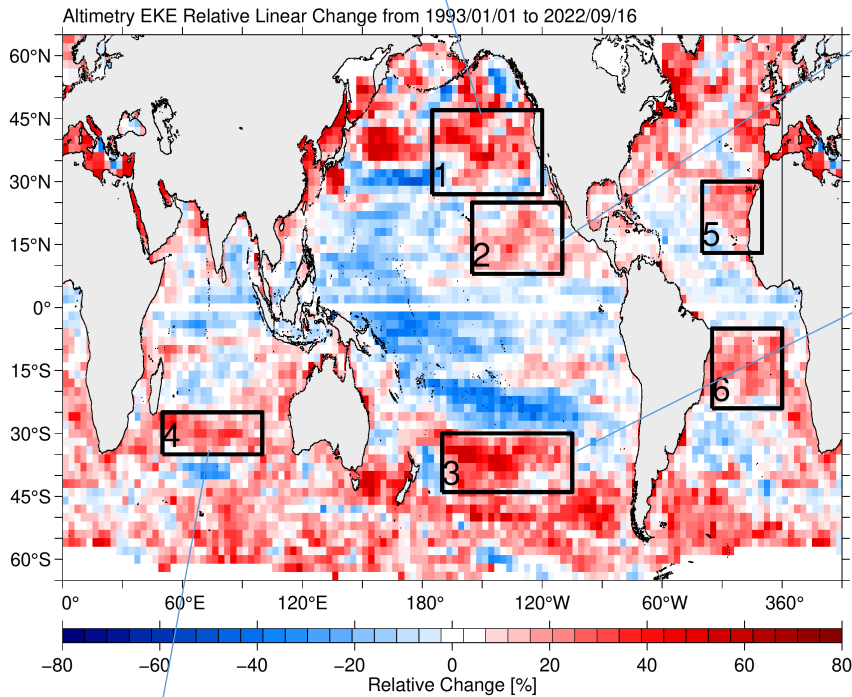
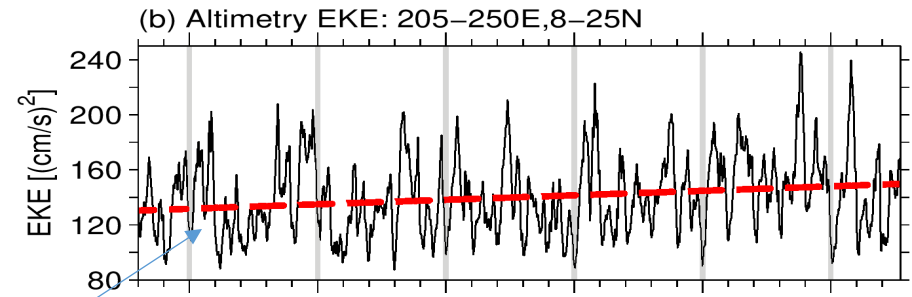
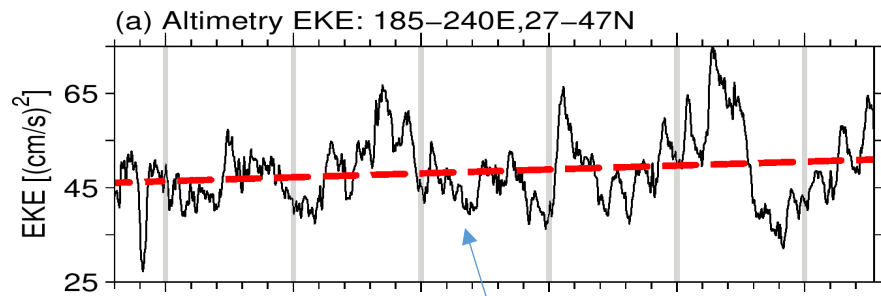
EKE change relative to local mean from 1993–2022



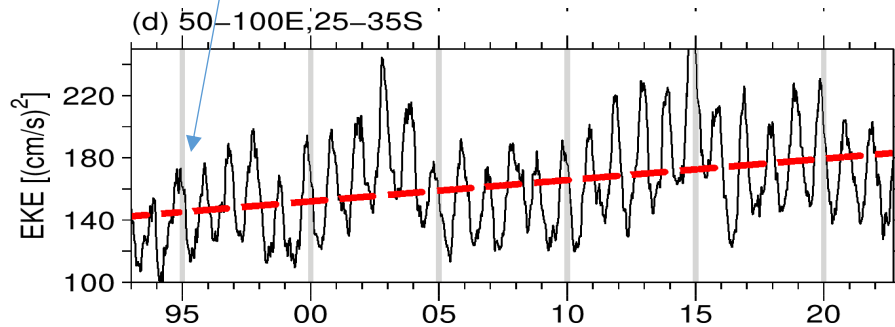
- Surface ocean EKE distribution has been established since beginning of the satellite altimetry era
- 3-decade accumulation of altimetry data allows us now to examine EKE linear trends; right map shows **relative EKE change** over the past 30-year period
- Increasing trends are seen in most of the **interior quiescent world oceans**
- EKE trend has a spatial pattern different from that of **regional sea level rise trend**, even though both patterns (as we will see below) are related to wind forcing

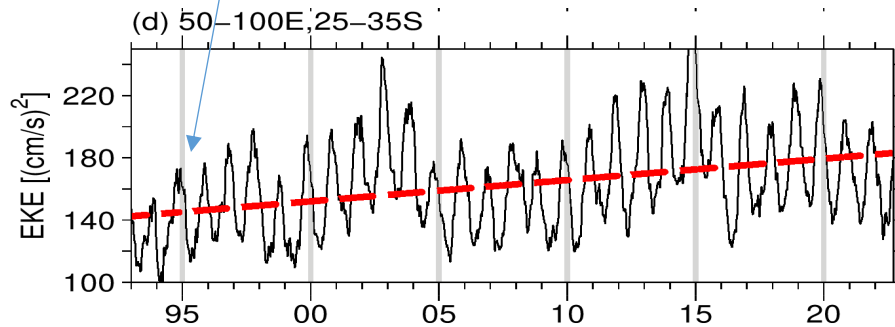
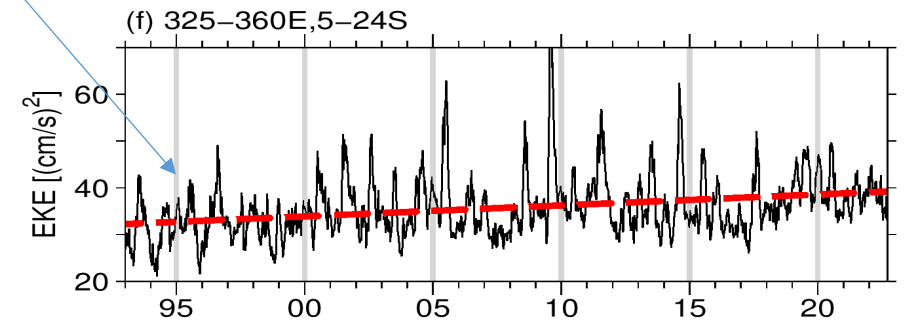
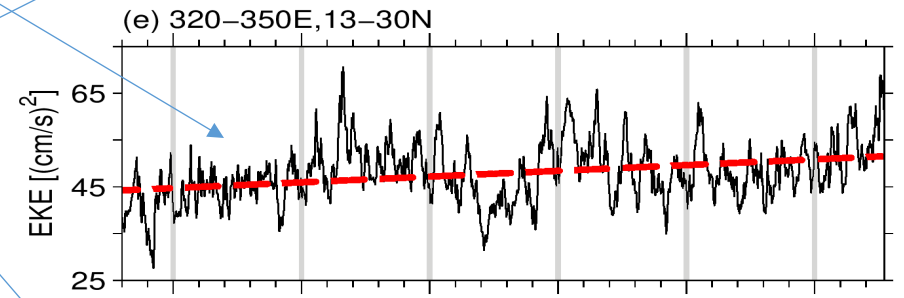
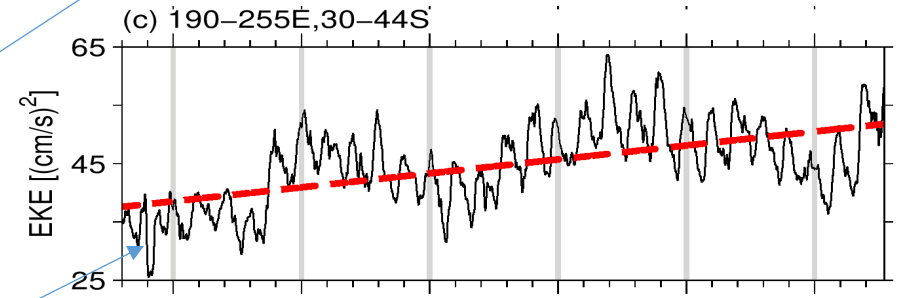
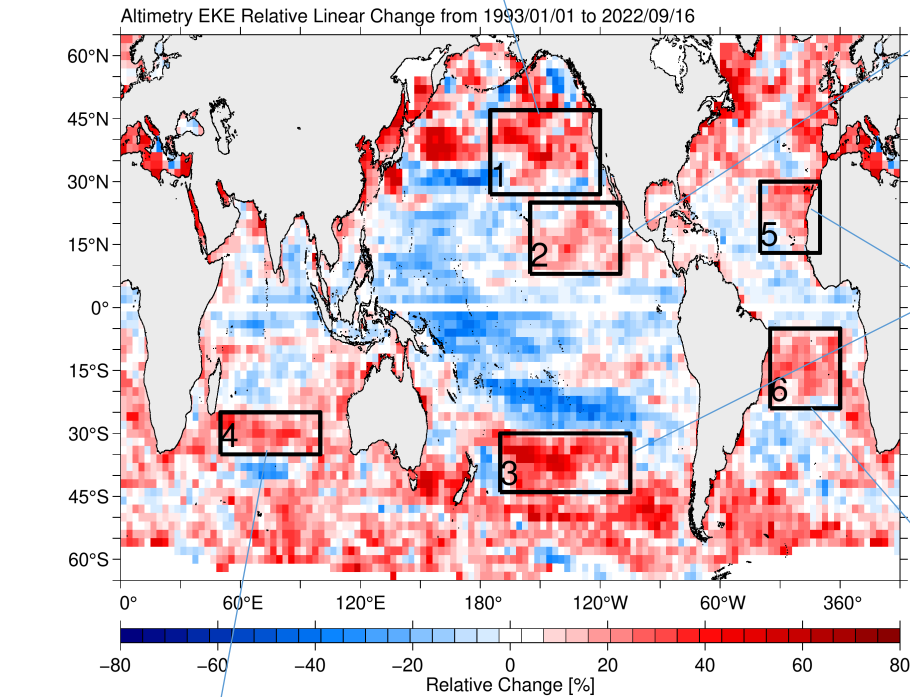
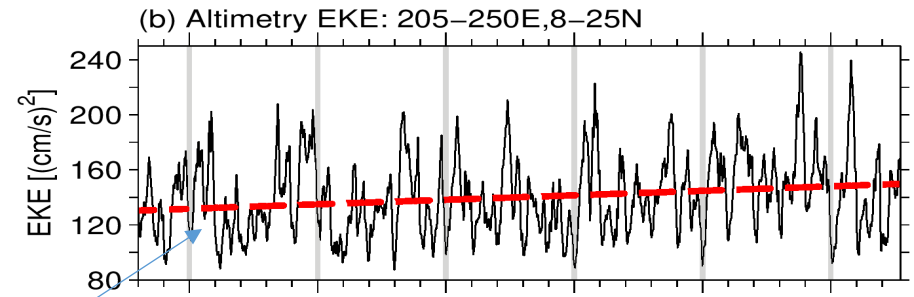
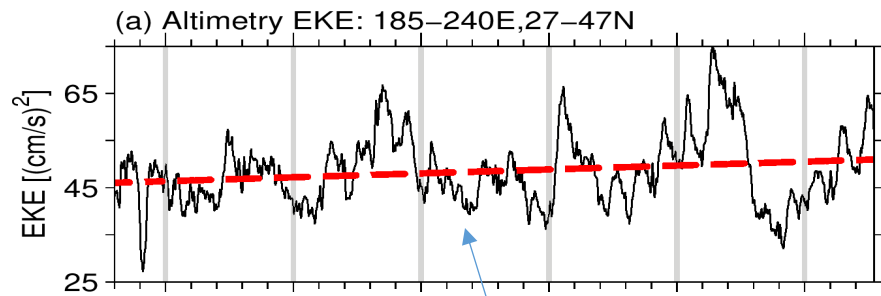


- While showing increasing trends, the EKE characteristics differ among different regions

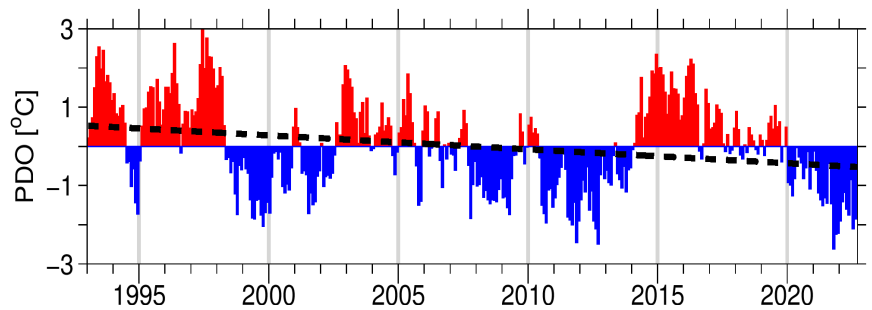
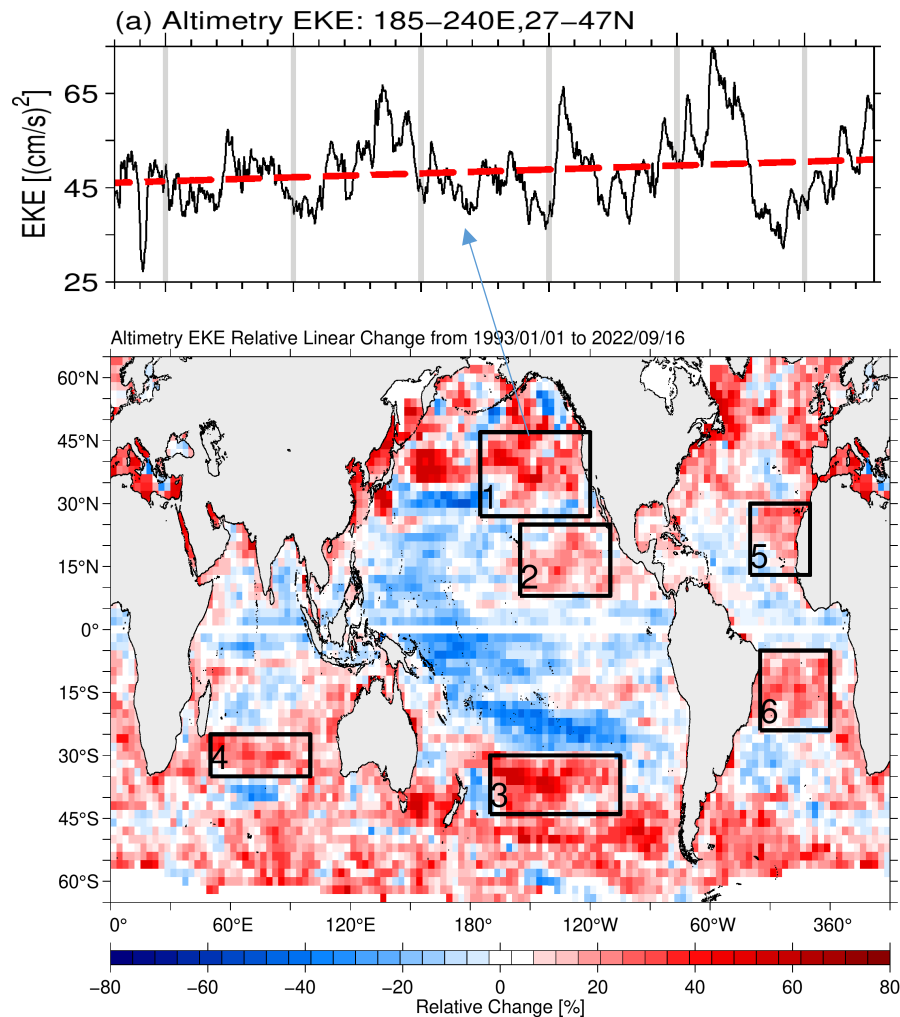


- Interior South Indian & Pacific Oceans reveal dominance of coherent seasonal-to-interannual modulations



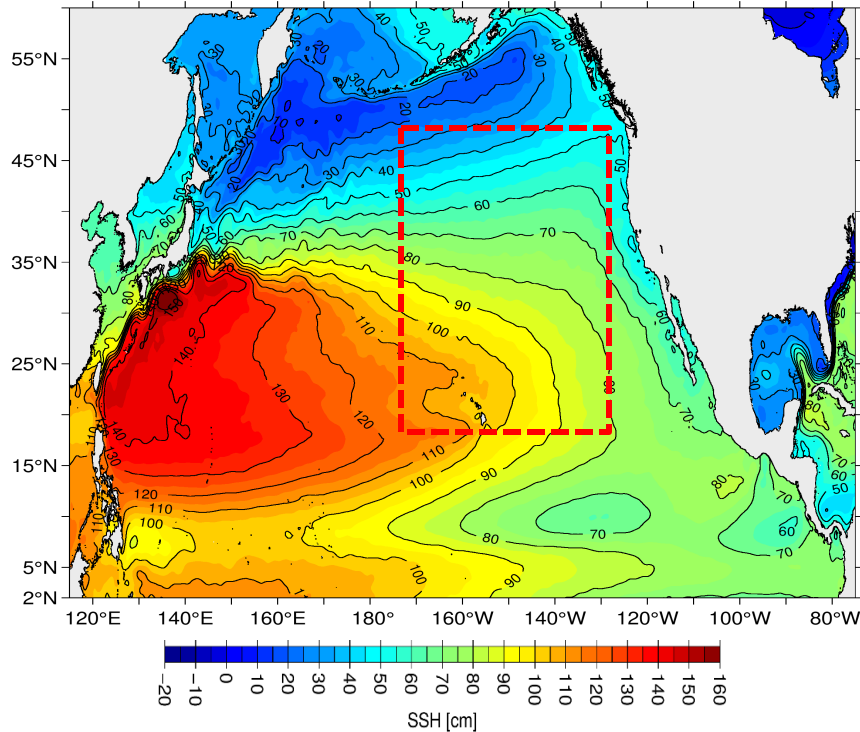


- Higher-frequency signals appear dominant in North & South Atlantic Oceans

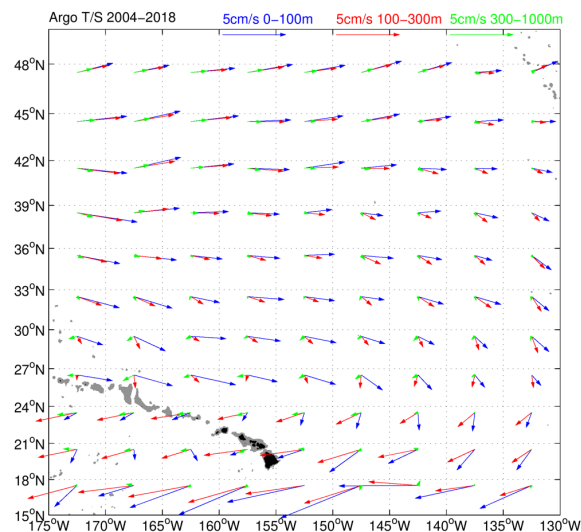


- This presentation will explore causes for the EKE trend in the **NE Pacific region**
- Given the low-frequency EKE signals in the region, it's natural to seek connection with the **PDO-related wind forcing**
- A negative PDO index trend exists, although there is little concurrent interannual-to-decadal EKE variability that matches the PDO forcing

Mean Dynamics Topography (CLS18, 20 Years: 1993–2012)



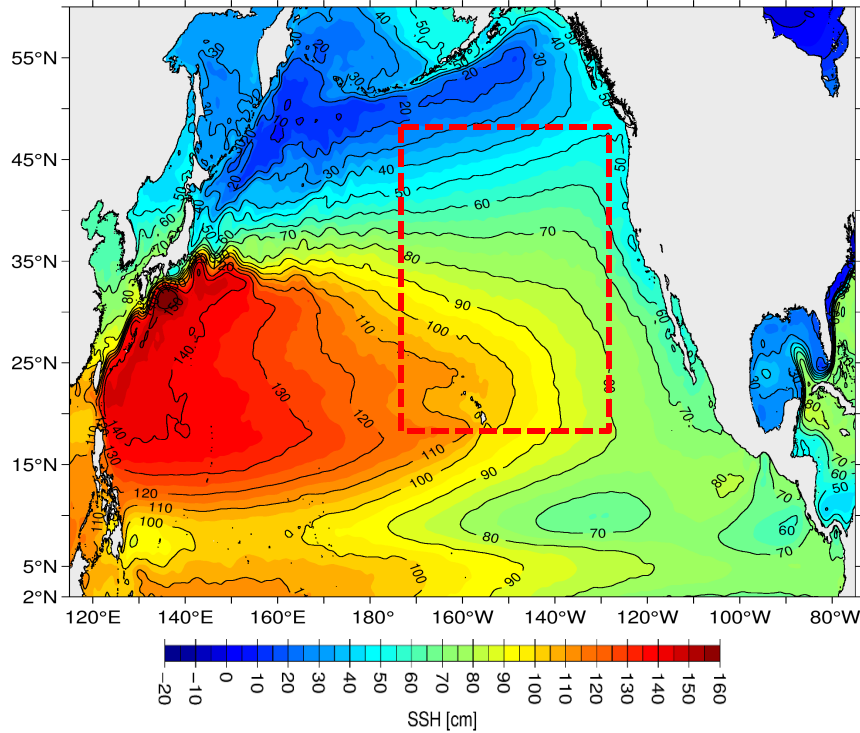
Upper ocean velocities based on 2004-2018 Argo data



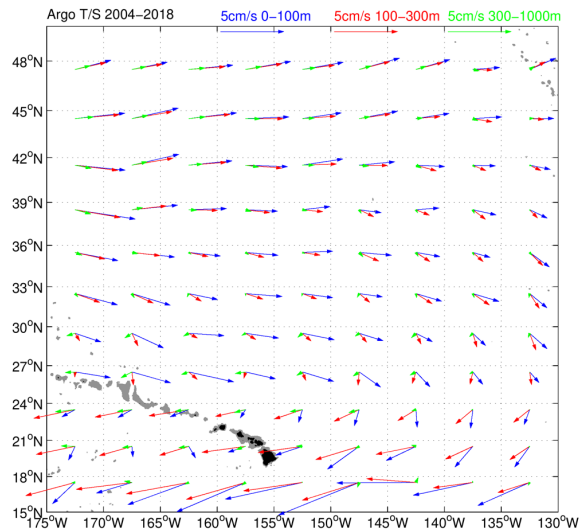
**Question:** What drives the interannual & longer timescale EKE variability in the NE Pacific Ocean?

- Located in the NE quadrant of the wind-driven subtropical gyre, the upper ocean flows veer clockwise with increasing depth →  $\beta$  spiral structure
- Although the time-mean circulation in the region is weak,  $O(\text{a few cm/s})$ , the  $\beta$  spiraled flow system is baroclinically unstable, especially for the meridional components

Mean Dynamics Topography (CLS18, 20 Years: 1993–2012)



Upper ocean velocities based on 2004-2022 Argo data



- Regional EKE growth can be quantified by stability analysis using a **3½-layer QG model** by specifying observed upper ocean velocity & stratification data:

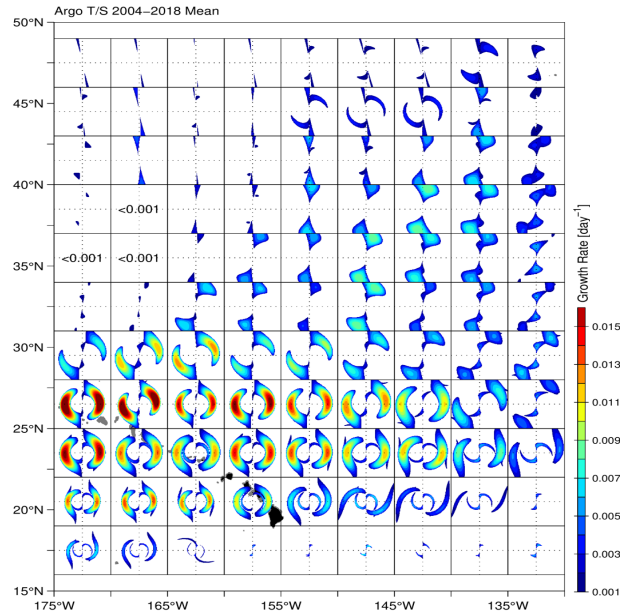
$$\left( \frac{\partial}{\partial t} + \vec{U}_n \cdot \nabla \right) q_n + J(\phi_n, \Pi_n) = 0, \quad (n = 1, 2, 3)$$

where  $q_n$  is perturbation PV,  $\Pi_n$  mean PV,  $\phi_n$  perturbation streamfunction, &

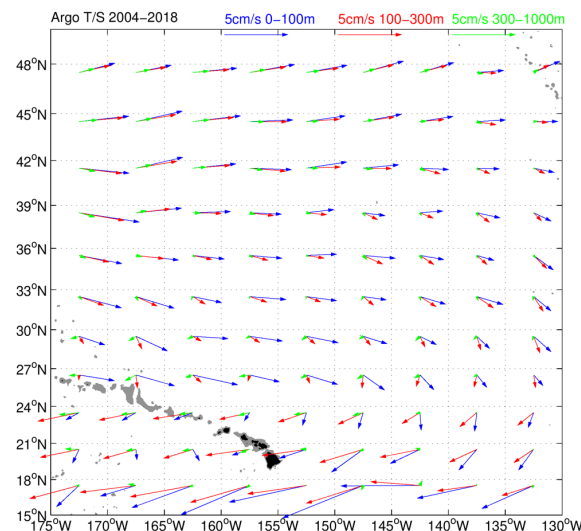
$$\begin{cases} q_1 = \nabla^2 \phi_1 + F_1(\phi_2 - \phi_1) \\ q_2 = \nabla^2 \phi_2 + F_2(\phi_1 - \phi_2) + G_2(\phi_3 - \phi_2) \\ q_3 = \nabla^2 \phi_3 + G_3[\phi_2 - (1 + \gamma)\phi_3], \end{cases}$$

$$g'_{ji} = g \frac{(\rho_j - \rho_i)}{\rho_0}, \quad F_i = \frac{f_0^2}{g'_{21} H}, \quad G_i = \frac{f_0^2}{g'_{32} H_i}, \quad \gamma = \frac{\rho_4 - \rho_3}{\rho_3 - \rho_2},$$

## Unstable wave growth rate $\omega_i$ as a function of $k$ & $l$ under 2004-2022 mean condition



## Upper ocean velocities based on 2004-2022 Argo data



- Regional EKE growth can be quantified by stability analysis using a **3½-layer QG model** by specifying observed upper ocean velocity & stratification data:

$$\left( \frac{\partial}{\partial t} + \vec{U}_n \cdot \nabla \right) q_n + J(\phi_n, \Pi_n) = 0, \quad (n = 1, 2, 3)$$

where  $q_n$  is perturbation PV,  $\Pi_n$  mean PV,  $\phi_n$  perturbation streamfunction, &

$$\begin{cases} q_1 = \nabla^2 \phi_1 + F_1(\phi_2 - \phi_1) \\ q_2 = \nabla^2 \phi_2 + F_2(\phi_1 - \phi_2) + G_2(\phi_3 - \phi_2) \\ q_3 = \nabla^2 \phi_3 + G_3[\phi_2 - (1 + \gamma)\phi_3], \end{cases}$$

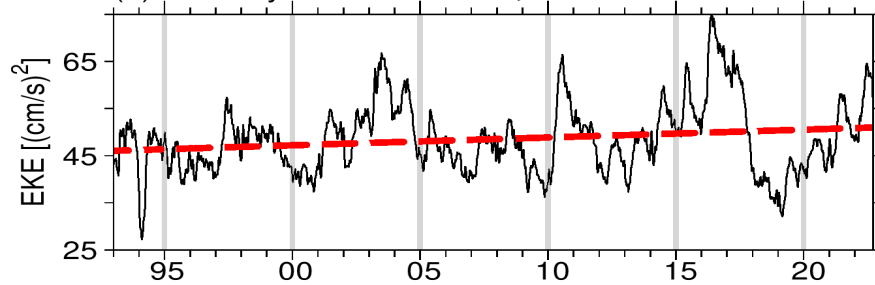
$$g'_{ji} = g \frac{(\rho_j - \rho_i)}{\rho_0}, \quad F_i = \frac{f_0^2}{g'_{21} H_i}, \quad G_i = \frac{f_0^2}{g'_{32} H_i}, \quad \gamma = \frac{\rho_4 - \rho_3}{\rho_3 - \rho_2},$$

- Assume normal-mode solutions for  $\phi_n$  :

$$\phi_n = A_n e^{i(kx + ly - \omega t)},$$

growth rate  $\omega_i$  of unstable waves can be found by solving an eigen-value problem

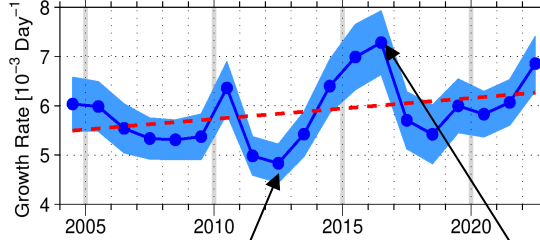
(a) Altimetry EKE: 185–240E, 27–47N



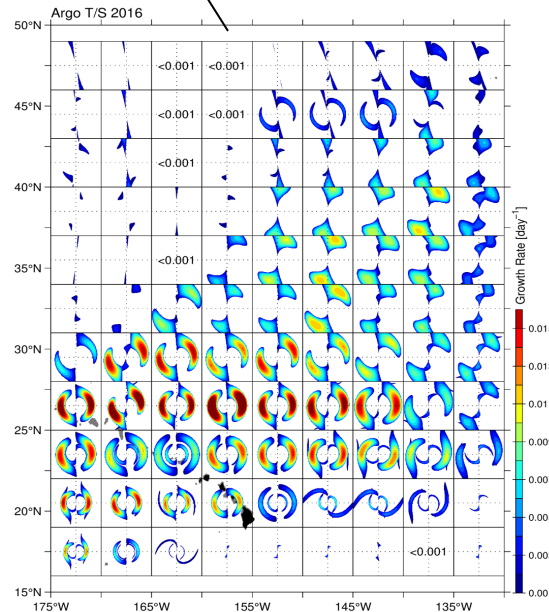
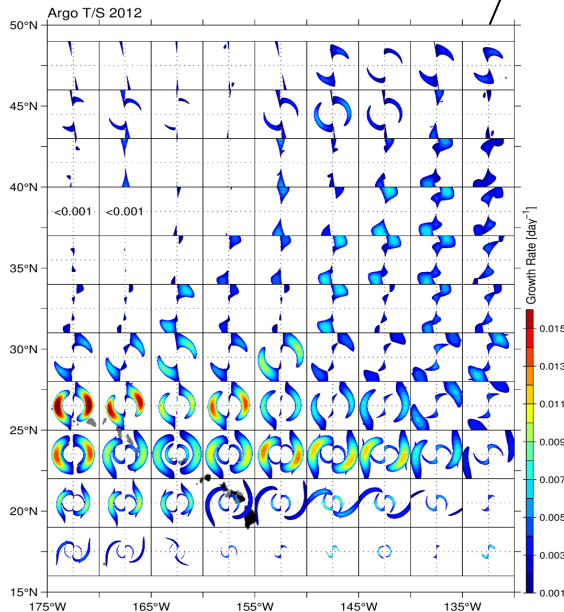
- Using yearly-varying upper ocean velocity & stratification data from Argo, we evaluated the growth rate variations in the NE Pacific

Growth rate  
averaged in  
NE Pacific

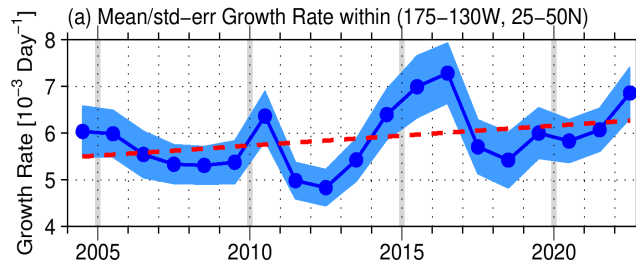
(a) Mean/std-err Growth Rate within (175–130W, 25–50N)



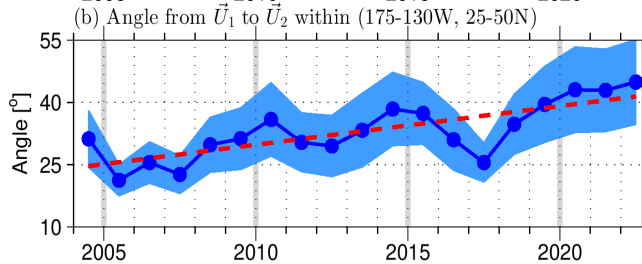
→ Favorable correspondence exists between the area-averaged growth rate & altimeter-measured EKE changes



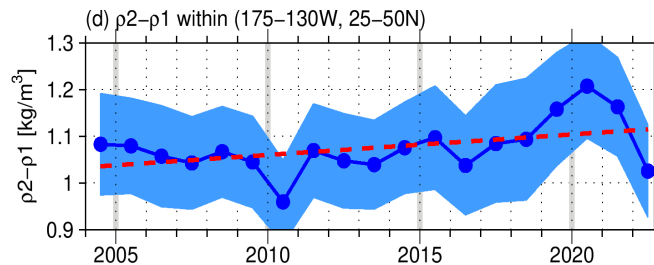
growth rate  
averaged in  
NE Pacific



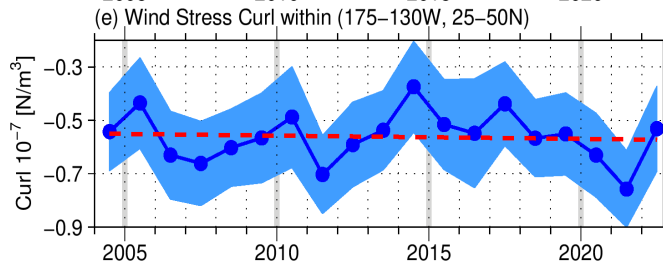
angle  $\theta$   
between  $U_1$   
&  $U_2$



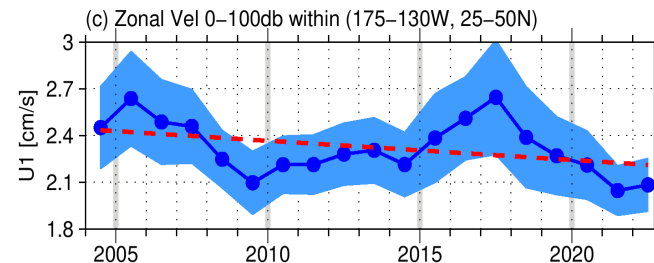
stratification  
 $\rho_2 - \rho_1$



wind  
stress curl



magnitude  
of  $U_1$



- Sensitivity study reveals the increased instability is largely due to the increase in angle  $\theta$  between surface & subsurface flow vectors in the past 2 decades

**Question:** What caused the angle in the upper ocean velocity field to increase during the past 2 decades?

- Dynamically,  $\theta$  change with depth is given by

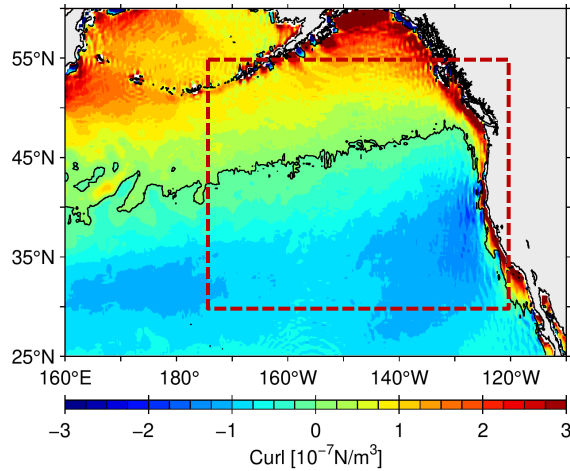
$$\frac{\partial \theta}{\partial z} = \frac{gw}{fU^2} \frac{\partial \rho}{\partial z}$$

where  $w$  is vertical velocity ( $\sim w_{Ek} < 0$  in NE Pacific), &  $U$  is flow speed

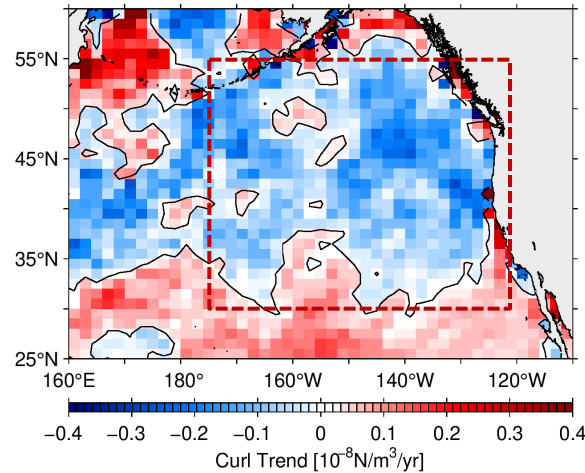
- In past 2 decades, amplitudes of  $w$  &  $\partial \rho / \partial z$  increased while that of  $U$  decreased  
 → all of which contributed to the increase in  $\partial \theta / \partial z$ , hence the regional instability

- Increase in  $\partial \rho / \partial z$  is related to the upper ocean warming

## ERA5 wind stress curl climatology

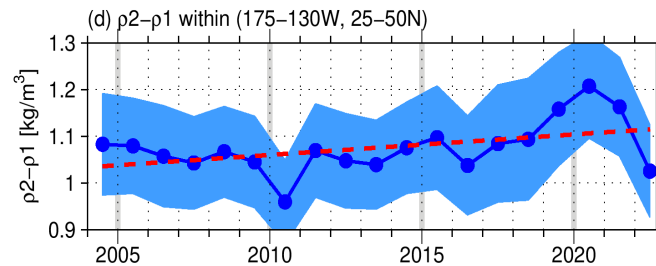


## ERA5 wind stress curl trend

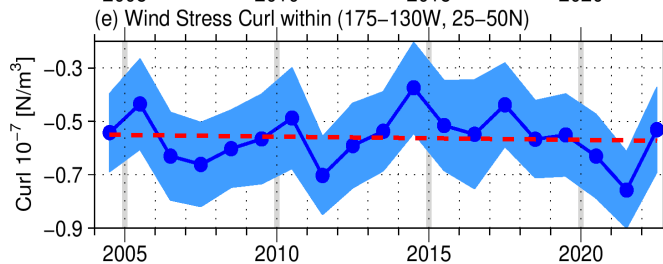


- Increase in  $|w|$  is related to the PDO phase transition from **positive** to **negative** (i.e., weakening of the Aleutian Low)

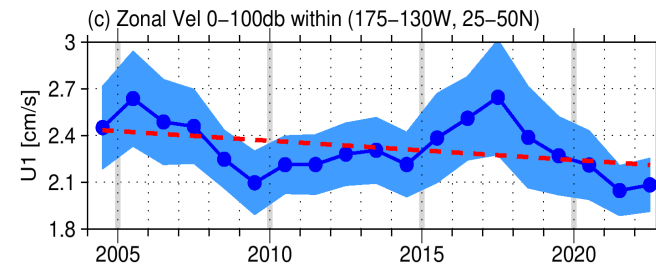
stratification  
 $\rho_2 - \rho_1$



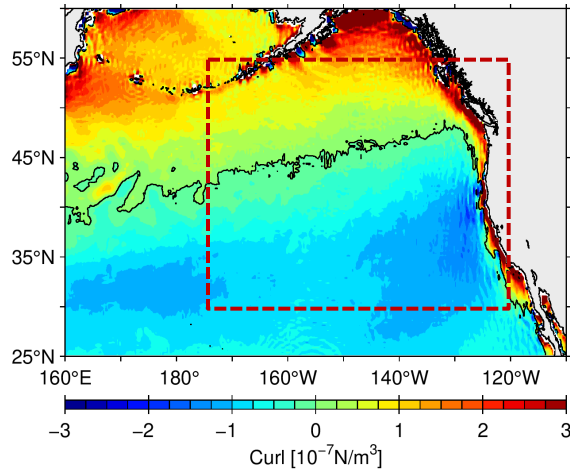
wind  
stress curl



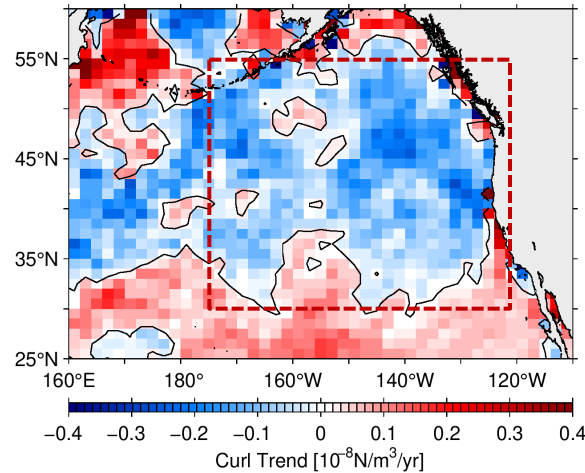
magnitude  
of  $U_1$



## ERA5 wind stress curl climatology



## ERA5 wind stress curl trend

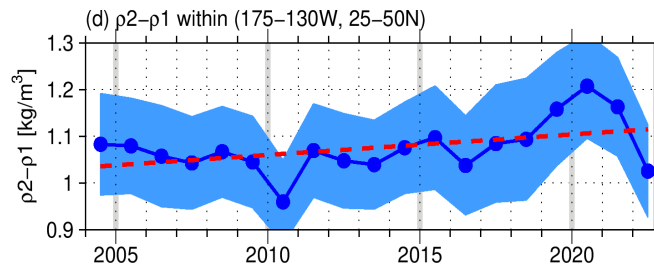


- Increase in  $|w|$  is related to the PDO phase transition from **positive** to **negative** (i.e., the weakening of the Aleutian Low)

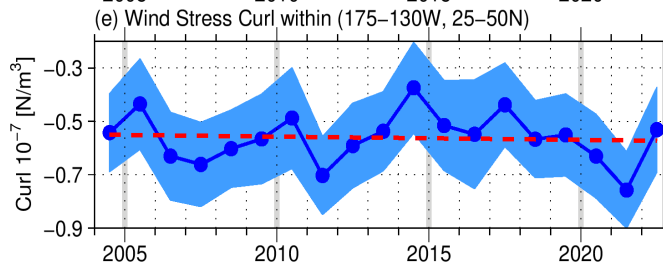
**Question:** Given the increase in negative wind stress curl (or Ekman pumping), why did the upper ocean velocity decrease?

stratification

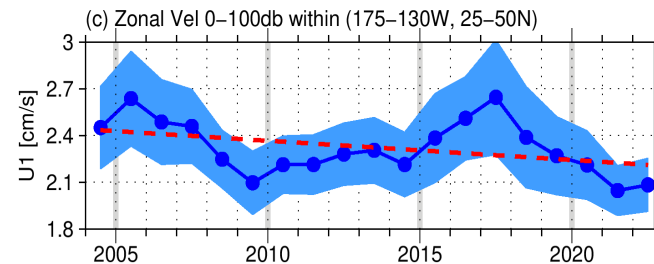
$$\rho_2 - \rho_1$$



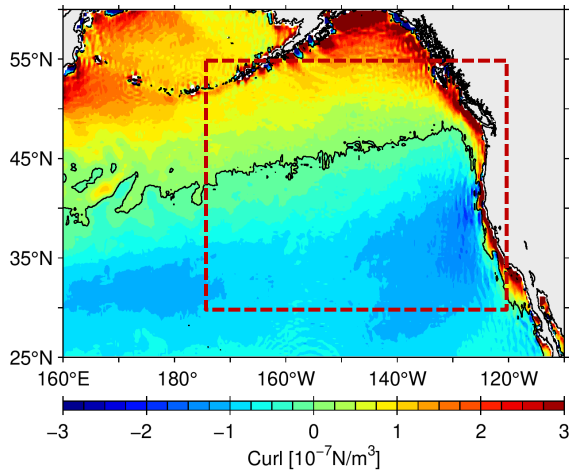
wind  
stress curl



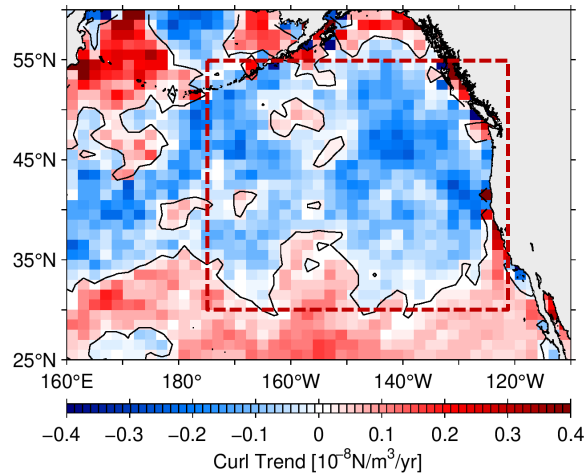
magnitude  
of  $U_1$



## ERA5 wind stress curl climatology



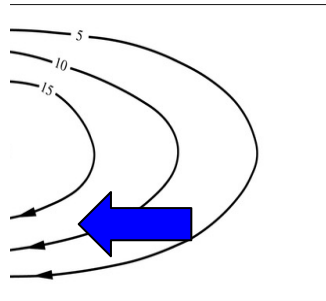
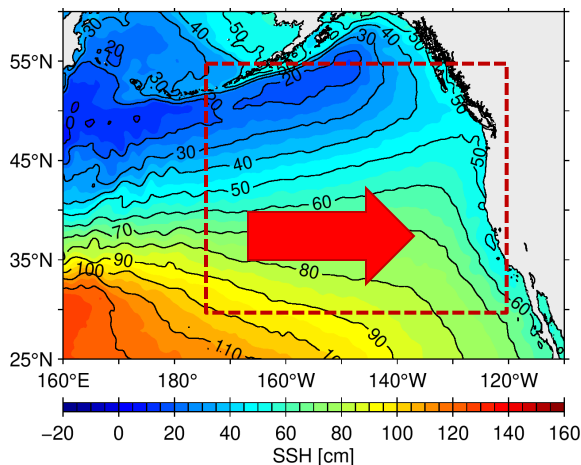
## ERA5 wind stress curl trend



- Increase in  $|w|$  is related to the PDO phase transition from **positive** to **negative** (i.e., the weakening of the Aleutian Low)

**Question:** Given the increase in negative wind stress curl (or Ekman pumping), why did the upper ocean velocity decrease?

## CLS18 mean dynamic topography

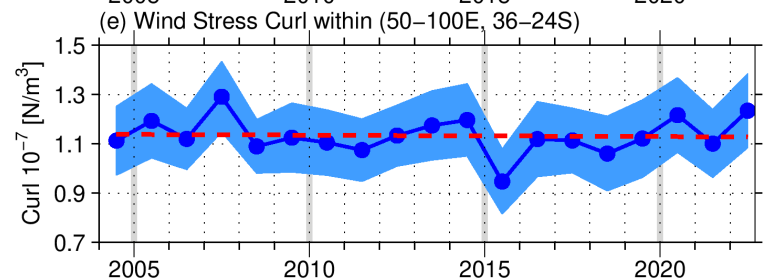
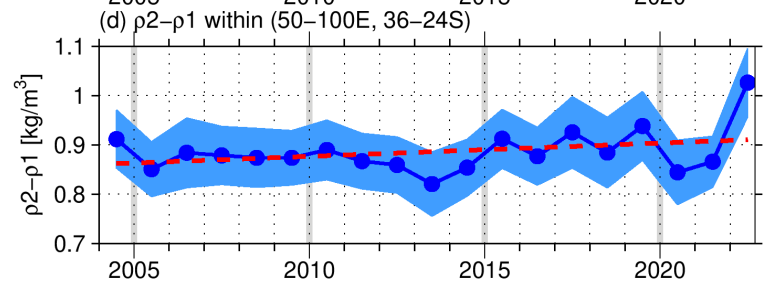
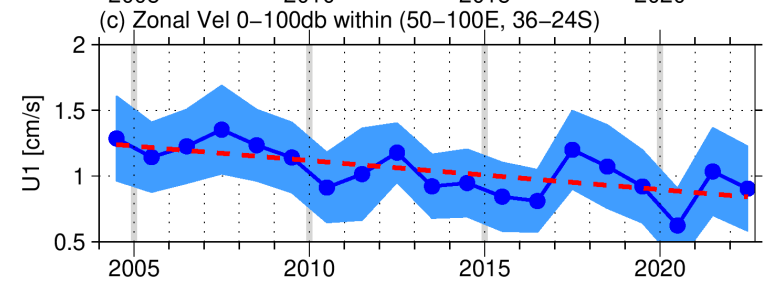
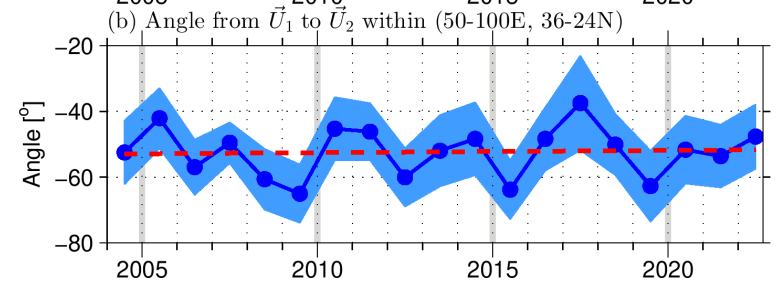
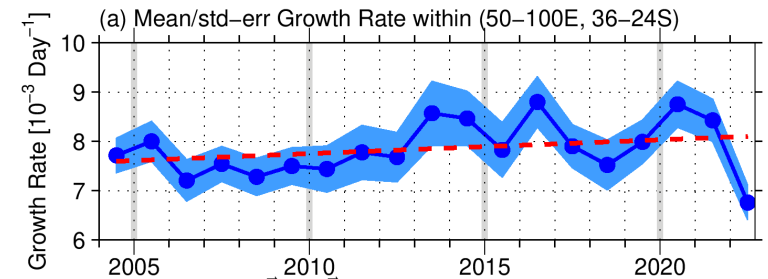
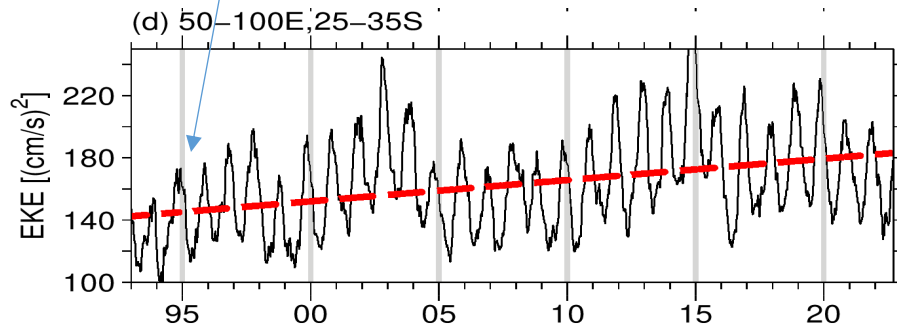
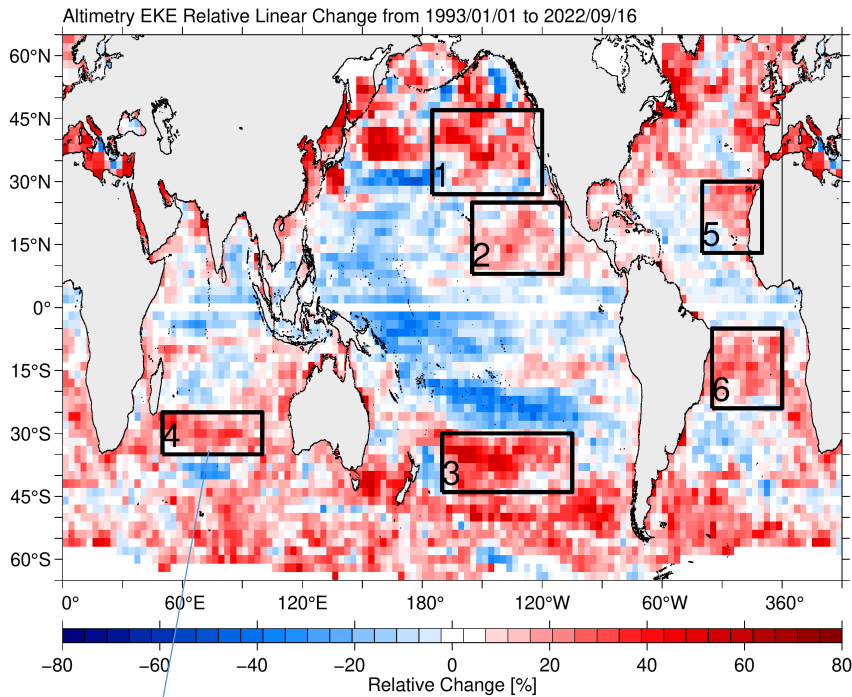


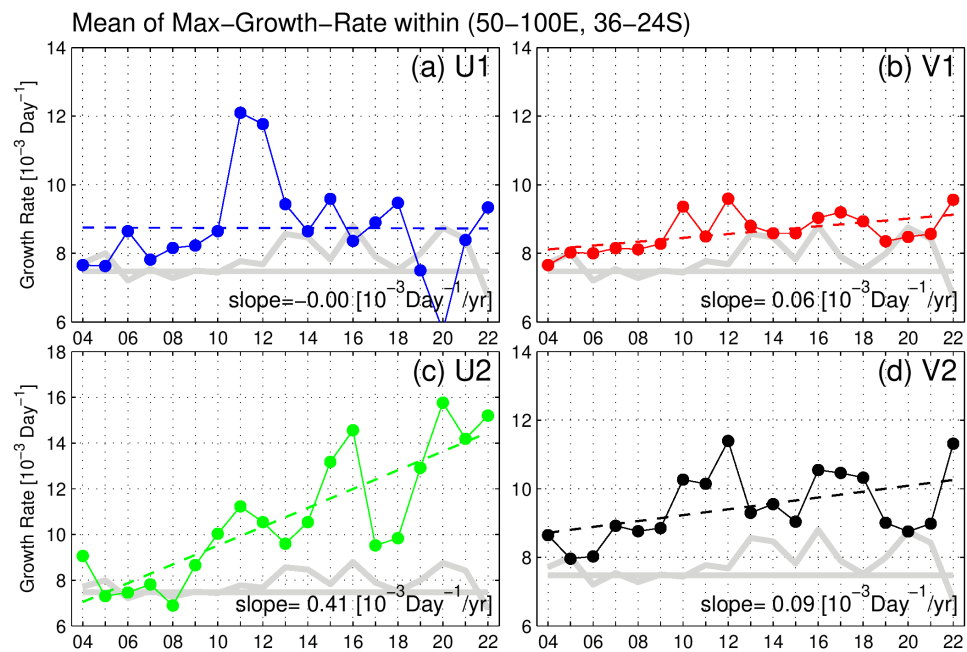
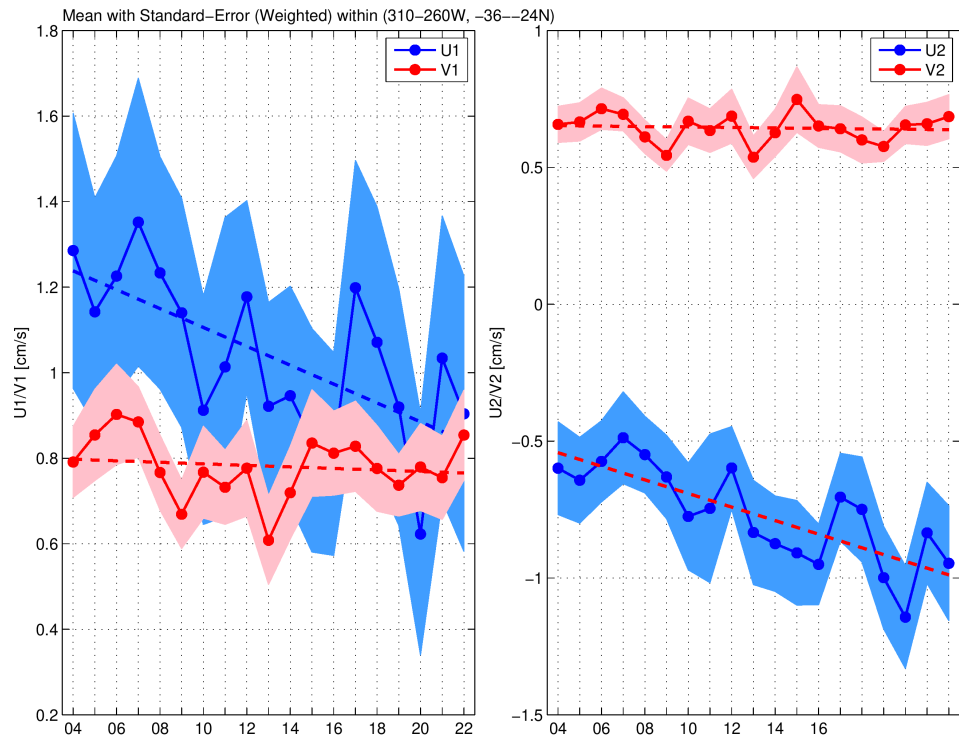
- The increasing negative wind stress curl forcing generates a regional clockwise anomalous circulation
- Its location is such that the anomalous-wind-induced **westward zonal flow** would oppose the climatological **eastward zonal flow** in the NE quadrant of the subtropical gyre

## Takeaway Message

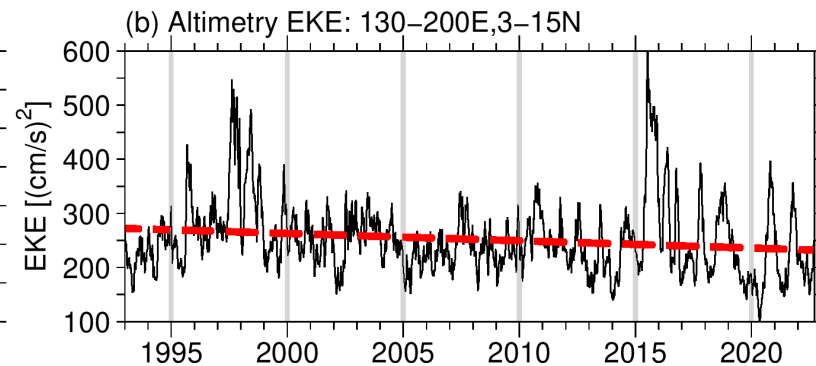
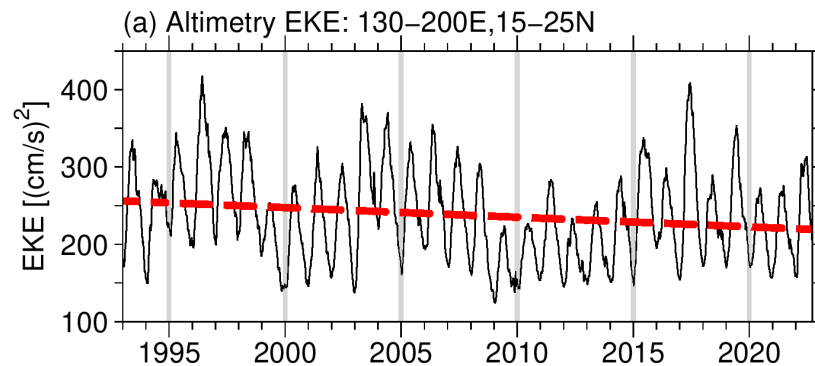
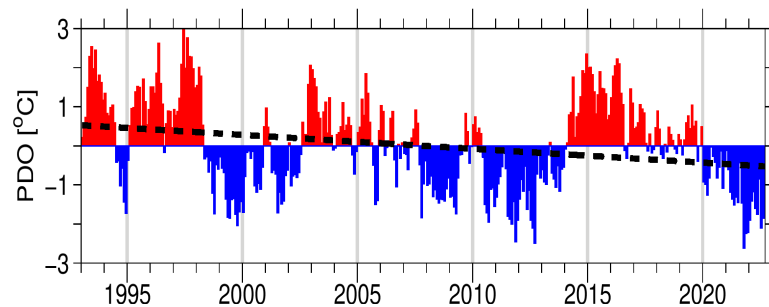
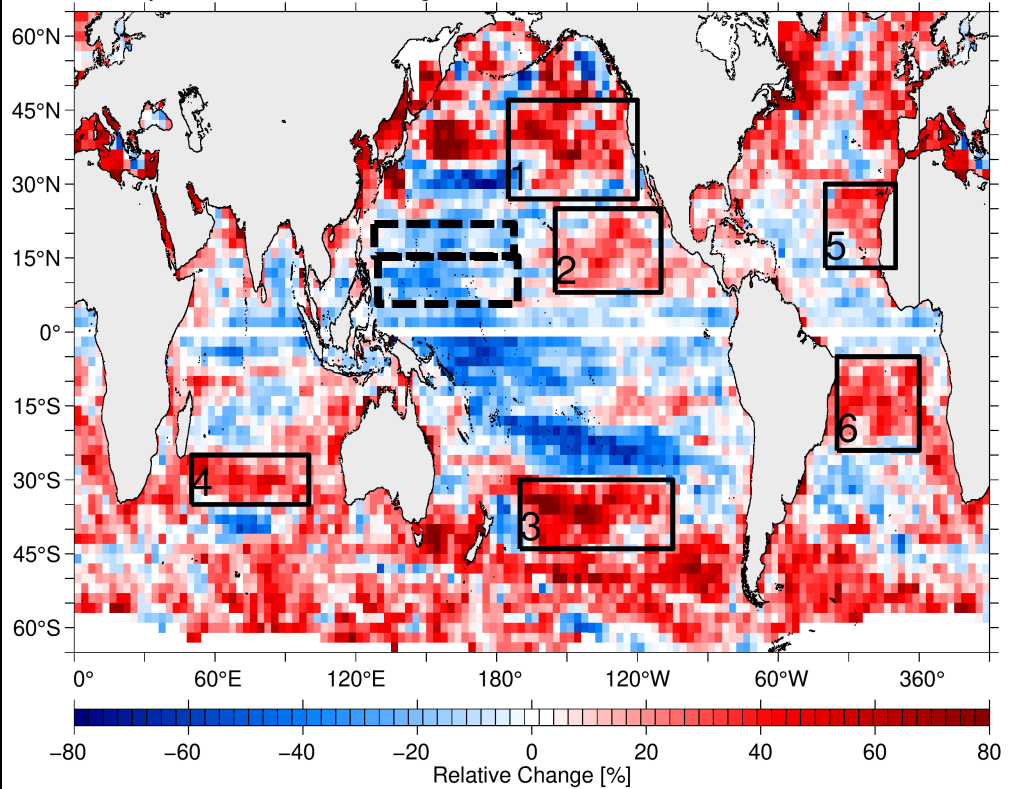
- In many of the weak EKE interior world oceans, there exist overall increasing EKE trends over the past 3 decades
- In the wind-driven subtropical gyre of the NE Pacific, the increasing EKE trend is caused by enhanced baroclinic instability related to broadened  $\beta$ -spiraled upper ocean velocity structure
- This broadening in  $\beta$ -spiraled velocity structure is contributed by 3 concurrent processes:
  1. Enhanced upper ocean stratification resulting from ocean warming,
  2. Increased Ekman pumping relating to the positive-to-negative phase shift of the PDO forcing, &
  3. Spin-up of an anomalous regional Sverdrup circulation that weakens the zonal flow in the NE Pacific subtropical gyre
- Our preliminary analyses indicate that the increasing EKE trends have different governing dynamics in different oceanic regions

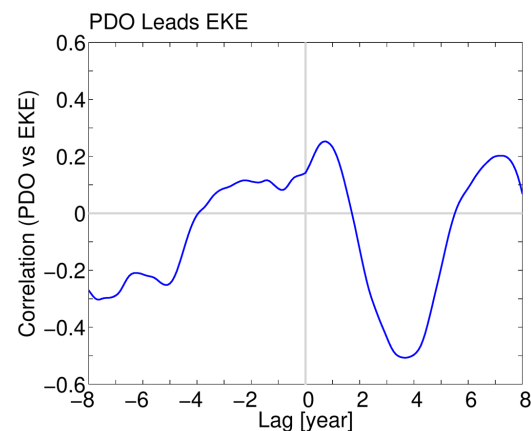
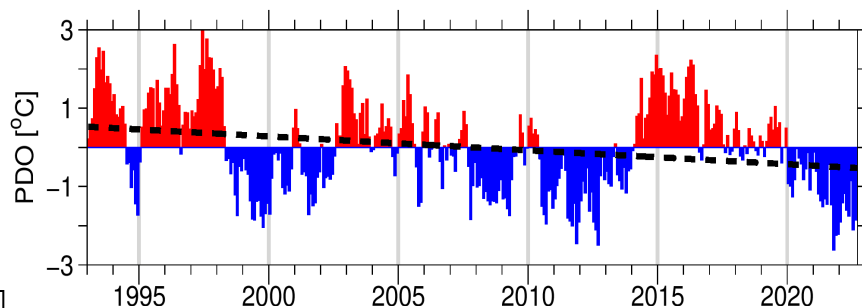
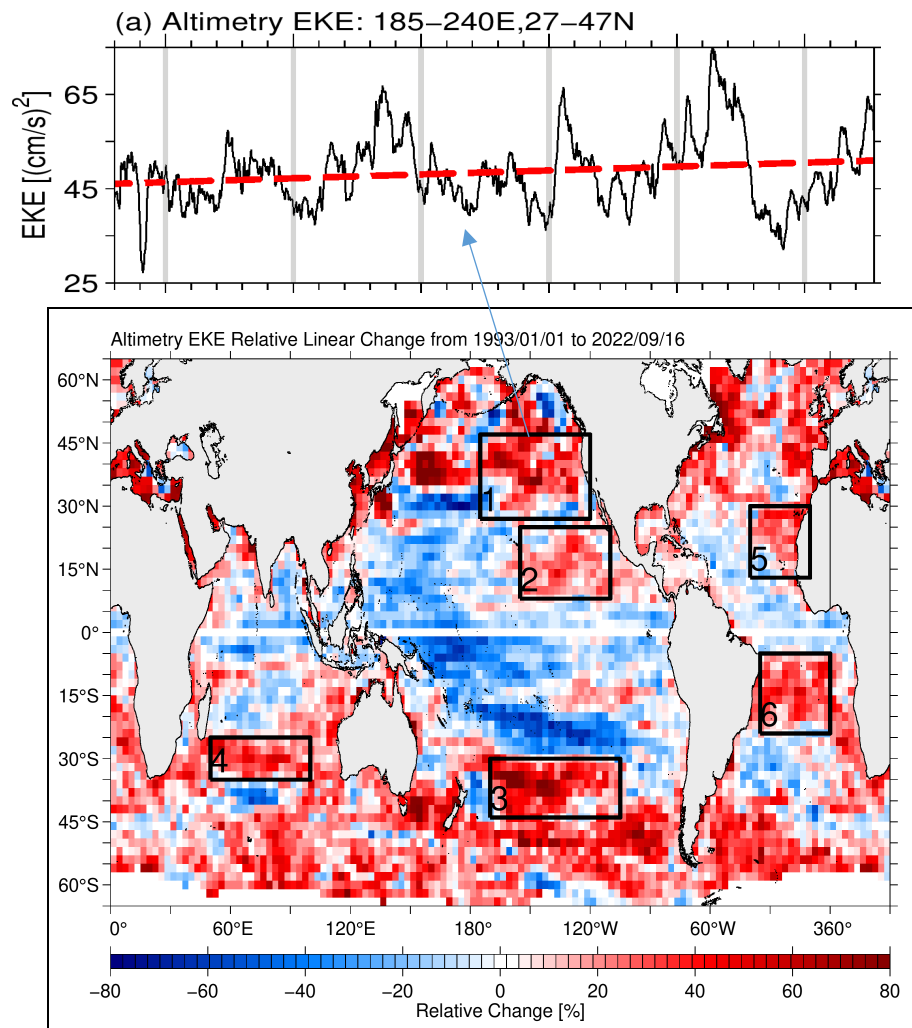
- Slight angle increase, amplitude of  $w$  constant,  $\partial\rho/\partial z$  increases &  $U$  decreases





Altimetry EKE Relative Linear Change from 1993/01/01 to 2022/09/16

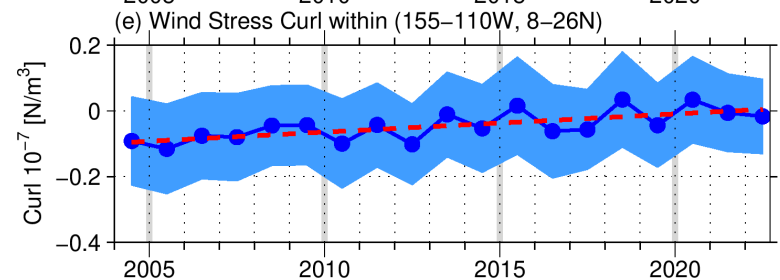
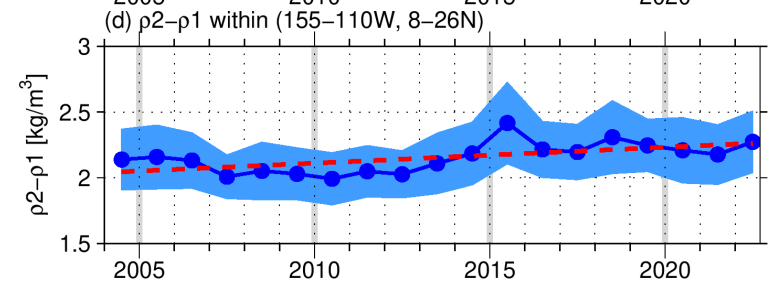
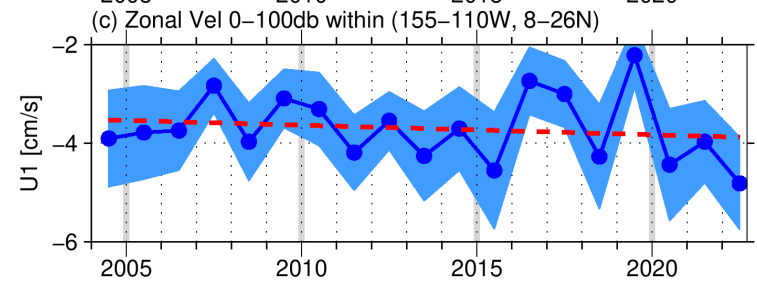
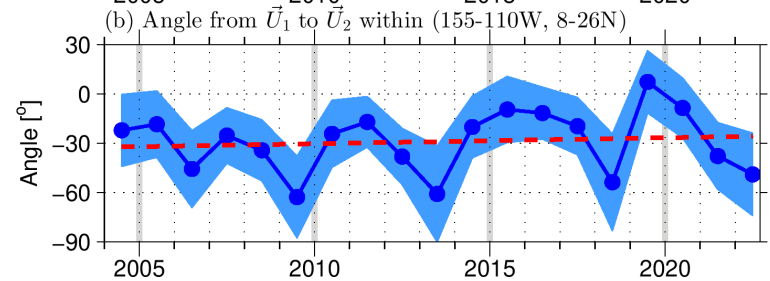
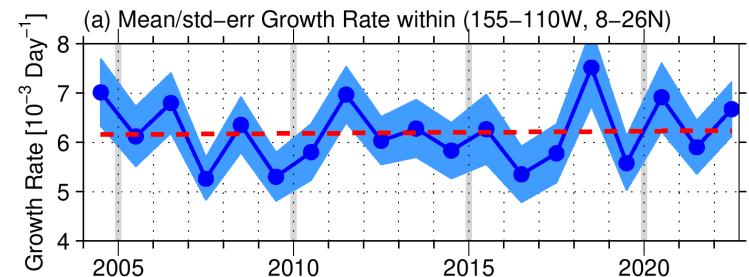
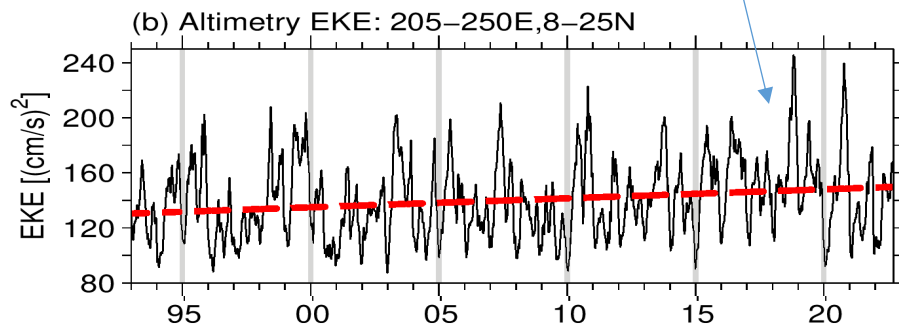
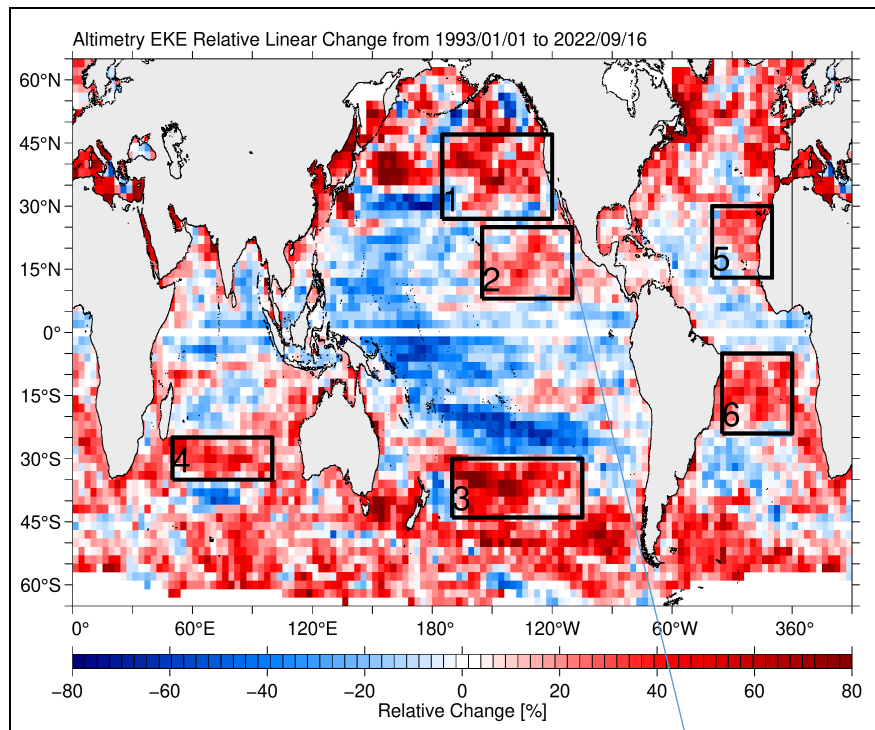




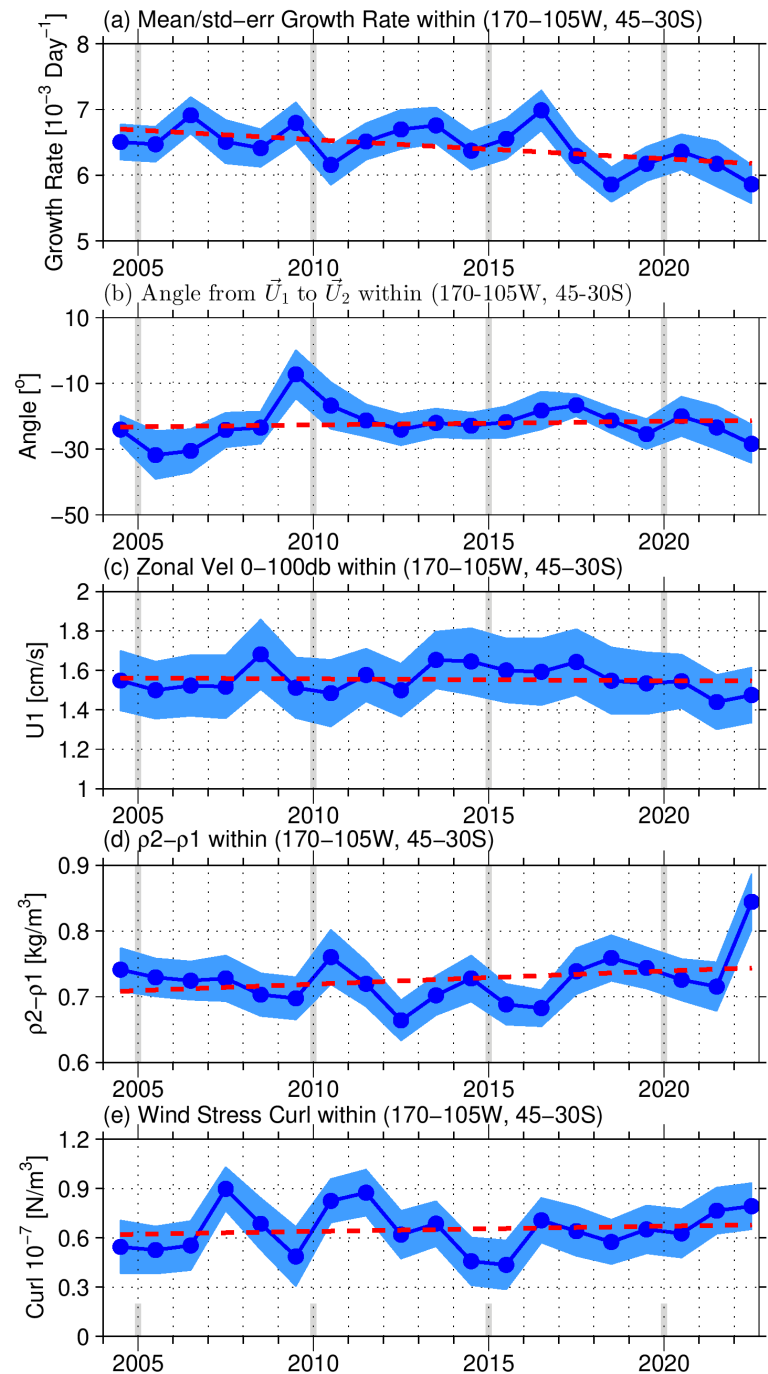
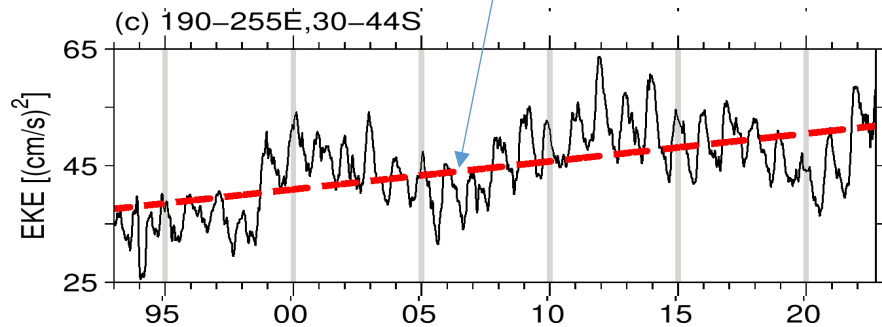
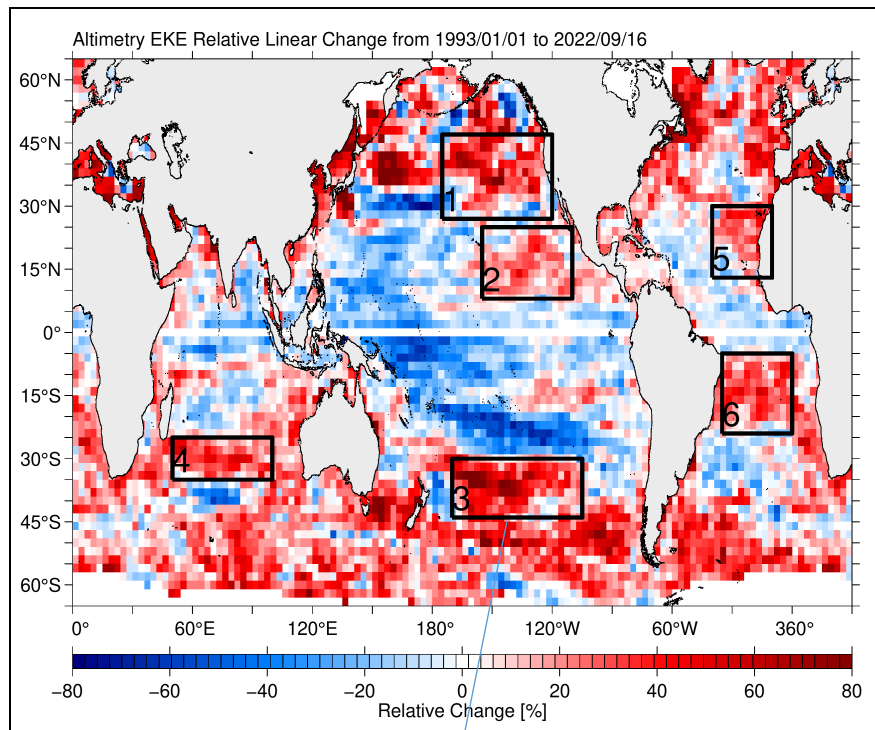
- Although PDO index shows a negative trend, the interannual-decadal PDO variability doesn't match the EKE signals.

- Given the low-frequency EKE signals in the NE Pacific, it's natural to seek connection with the PDO wind forcing

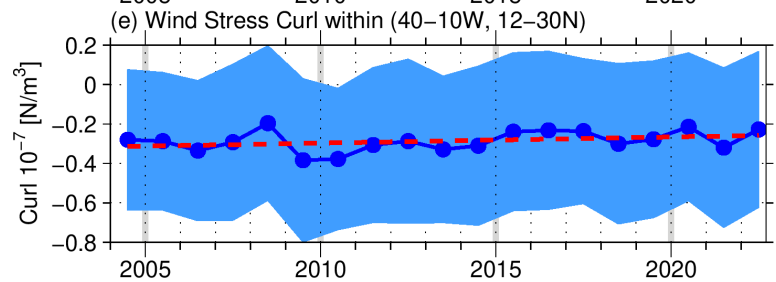
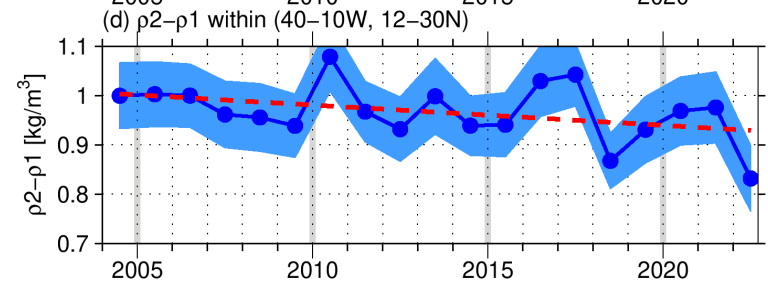
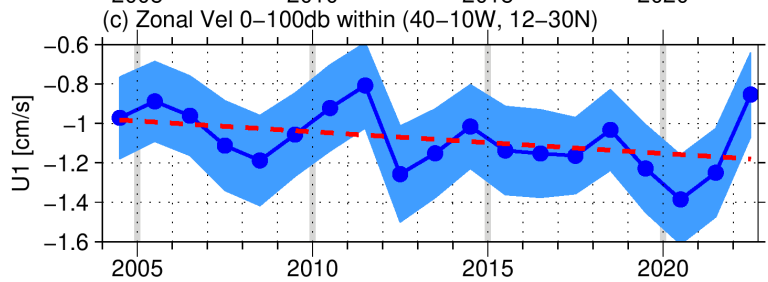
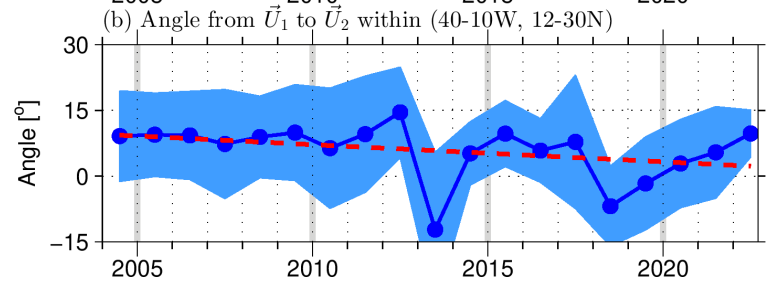
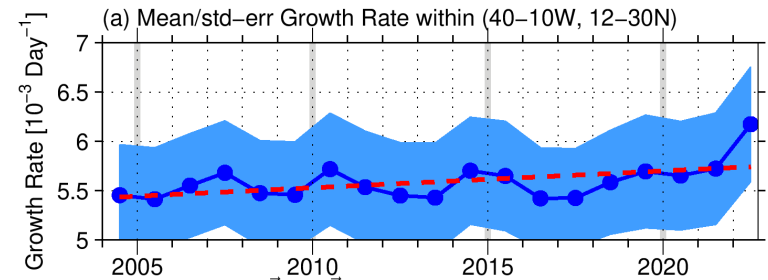
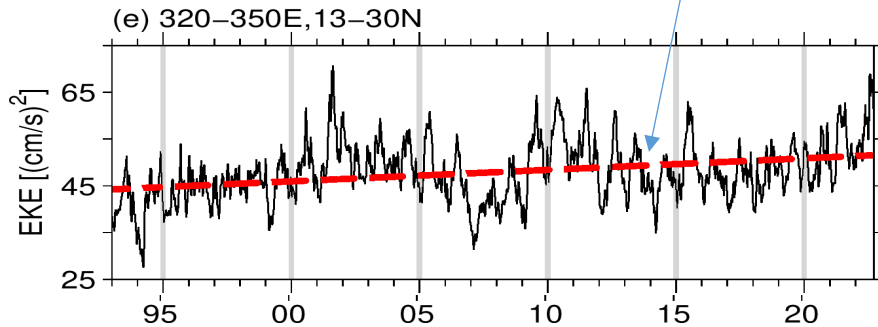
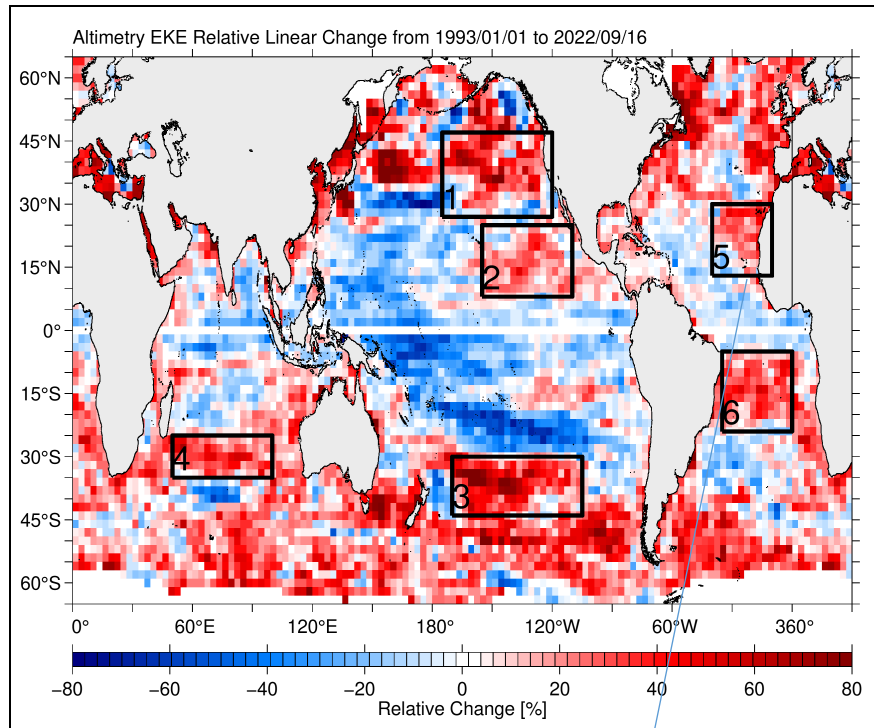
- Little angle change, amplitude of  $w$  decreases and  $\partial\rho/\partial z$  &  $U$  increases, cancelling out effects on  $\partial\theta/\partial z$



- Little angle change,  $w$  &  $U$  largely constant &  $\partial\rho/\partial z$  increase



- Angle decreases, amplitude of  $w$  constant,  $\partial\rho/\partial z$  decreases &  $U$  increases



- Angle increases, amplitude of  $w$  decreases,  $\partial\rho/\partial z$  increases &  $U$  increases

