



Nonlinear Short-Term Upper Ocean Circulation Variability in the Western Tropical Pacific

Bo Qiu, Shuiming Chen & Brian Powell

Department of Oceanography, University of Hawaii

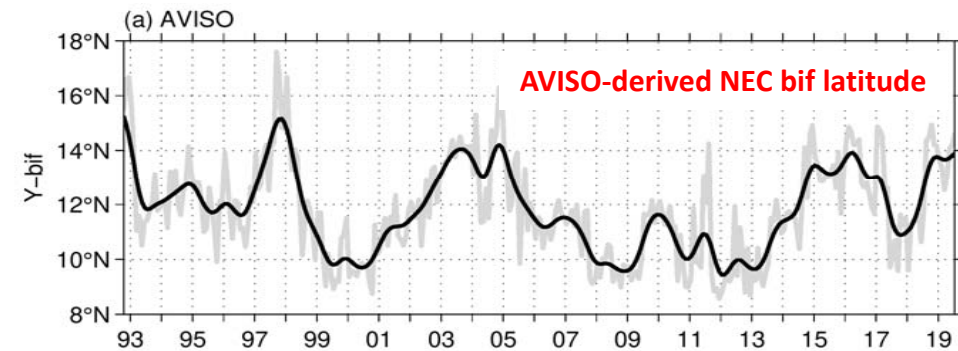
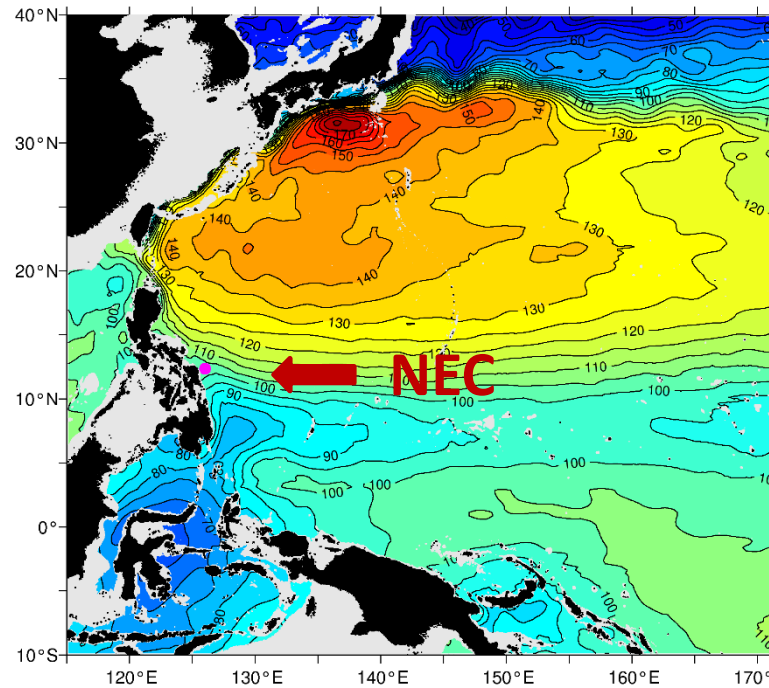
Patrick Colin

Coral Reef Research Foundation, Palau

Dan Rudnick & Martha Schönau

Scripps Institution of Oceanography, UCSD

Low-frequency sea level & circulation variability in tropical Pacific can be understood as wind-forced baroclinic responses



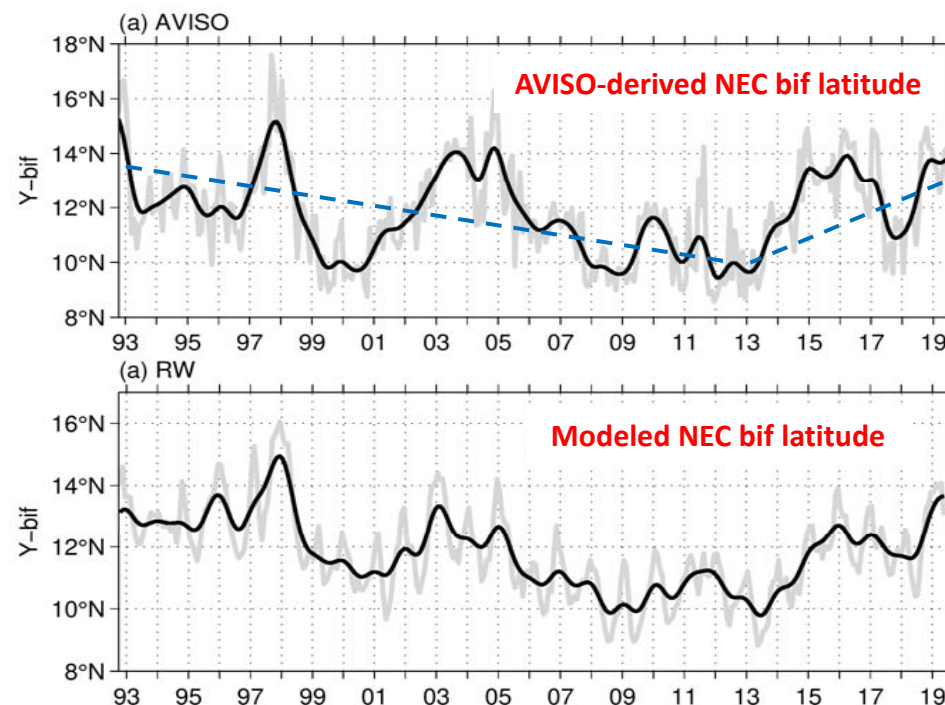
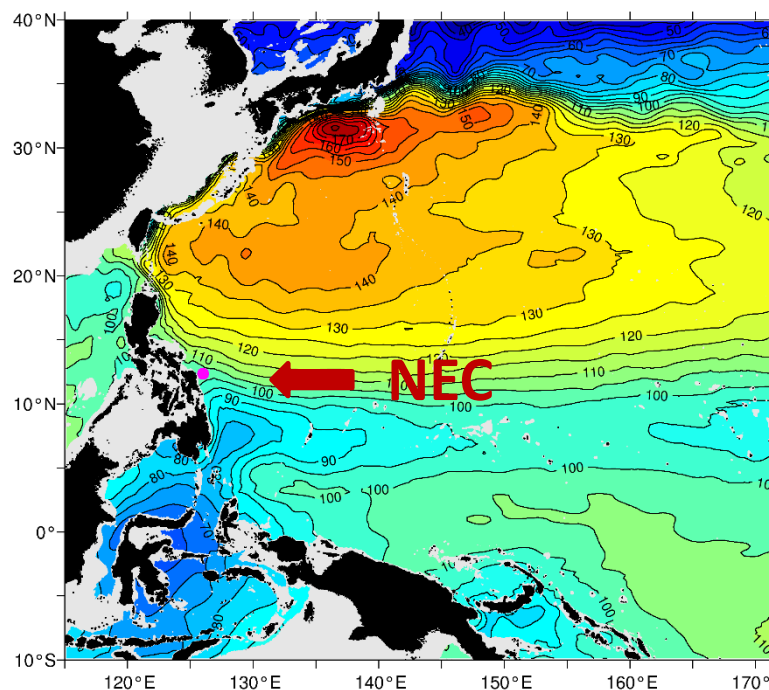
- Utilize the weekly AVISO SSH anomaly data
- Add mean SSH field of Rio et al. (2009): mean NEC bifurcation at ~12°N
- Calculate meridional geostrophic velocity as a function of y along the Philippine coast:

$$v_g(y, t) = \frac{g}{f} [h_e(y, t) - h_w(y, t)],$$

where h_w is SSH in 1°-band east of the coast & h_e in 1°-band further to the east

- NEC bifurcation latitude $Y_b(t)$ is defined at where $v_g = 0$

Low-frequency sea level & circulation variability in tropical Pacific can be understood as wind-forced baroclinic responses

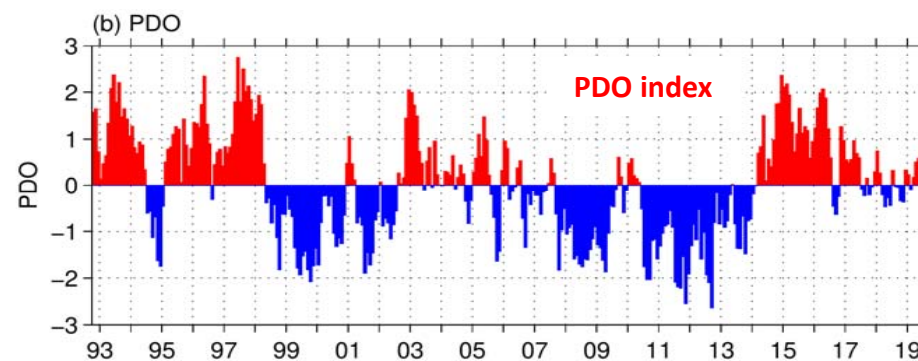
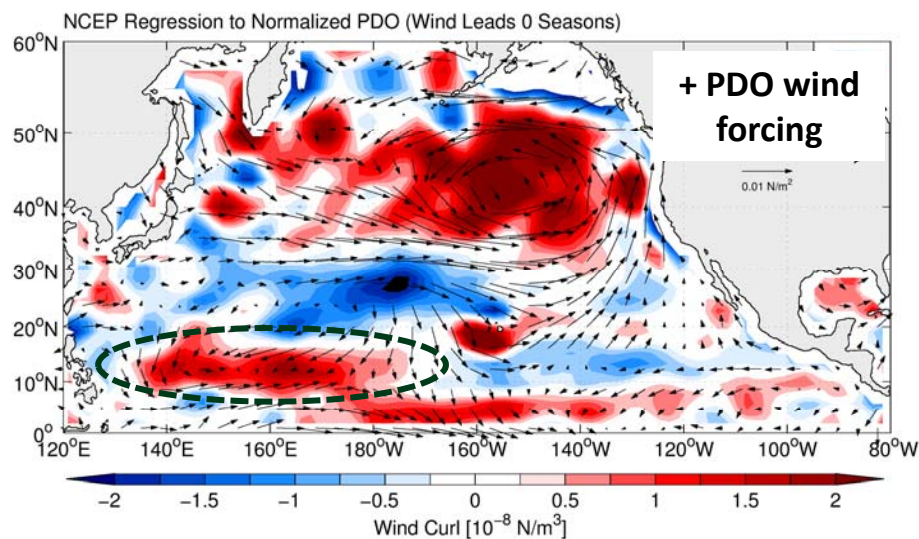
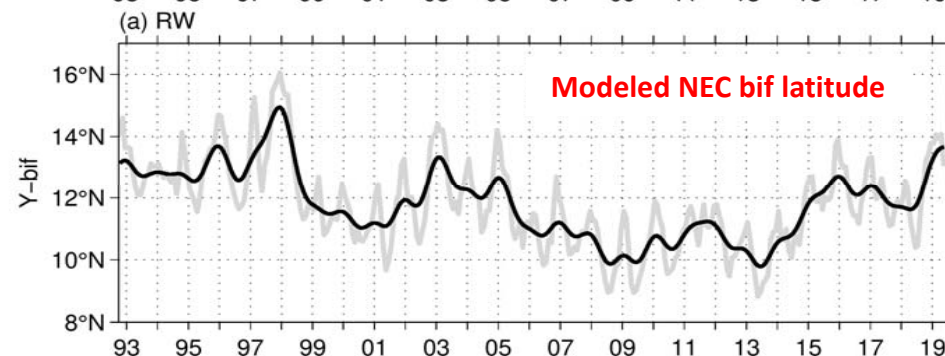
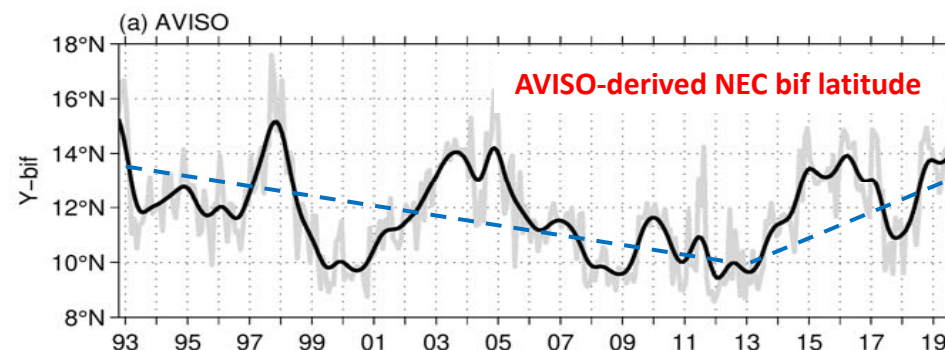
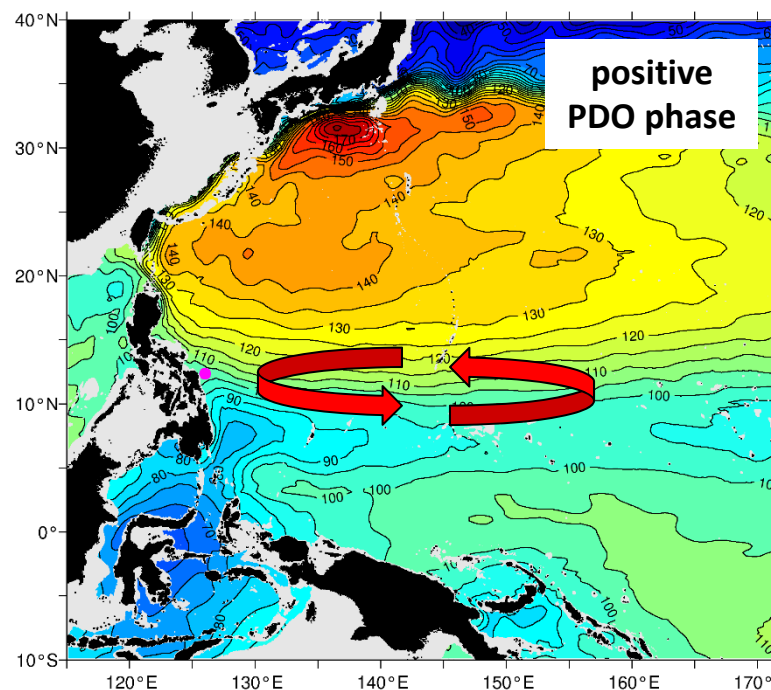


- NEC position change can be represented by sea level change off the Philippine coast
- Low-frequency sea level change is governed by wind-forced baroclinic RW dynamics:

$$\frac{\partial h'}{\partial t} - c_R \frac{\partial h'}{\partial x} = -\frac{g'}{g\rho_o} \text{curl} \left(\frac{\tau}{f} \right) - \epsilon h'$$

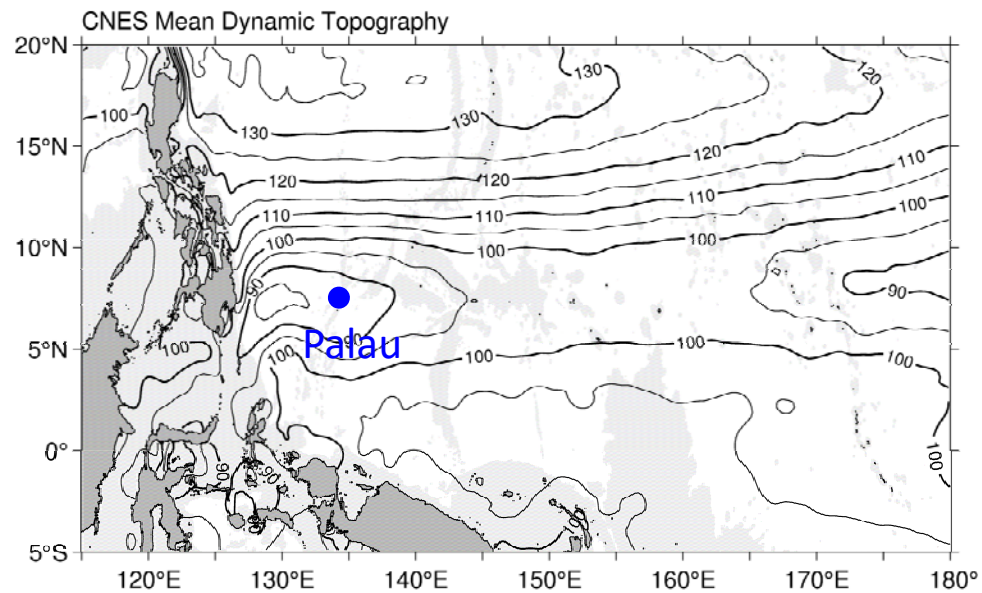
↓
Ekman pumping

Low-frequency sea level & circulation variability in tropical Pacific can be understood as wind-forced baroclinic responses

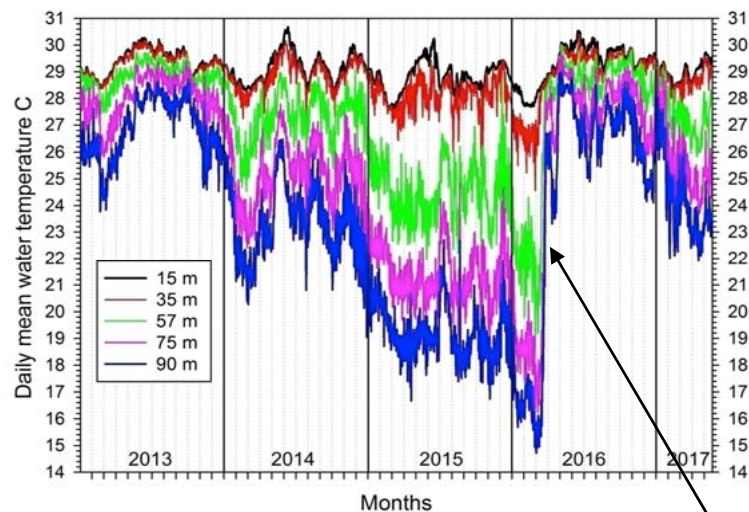
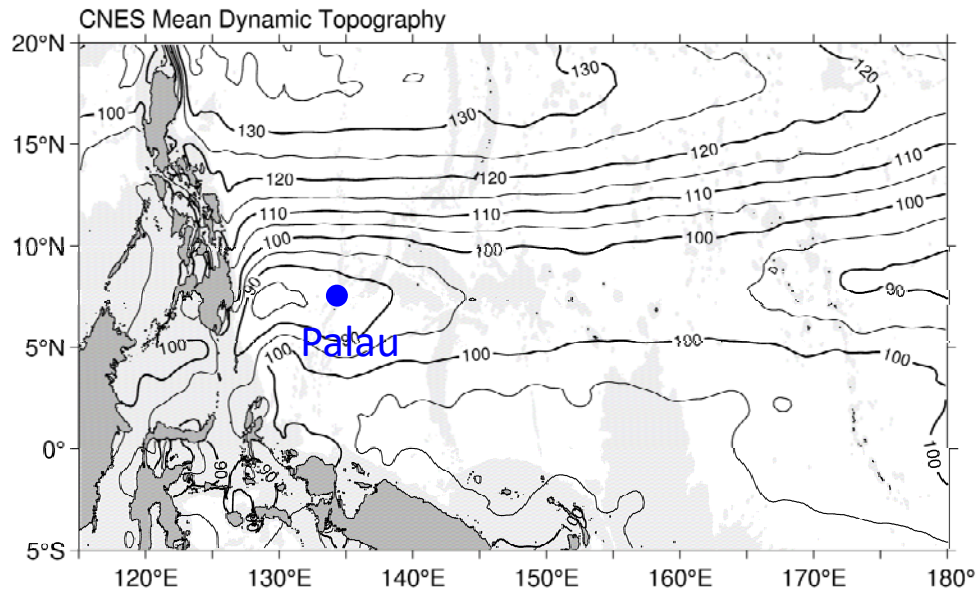


+ PDO forcing → Ekman suction off Philippines
→ low sea level → poleward NEC shift

- Is the **linear RW dynamics** valid for upper ocean variability around **Palau** ?



- Is the **linear RW dynamics** valid for upper ocean variability around **Palau** ?

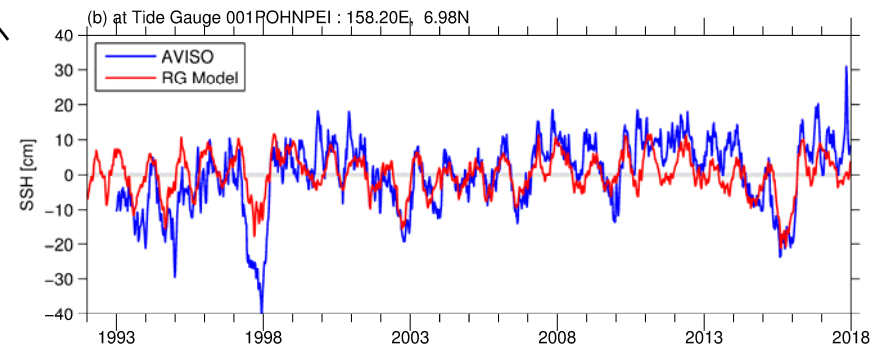
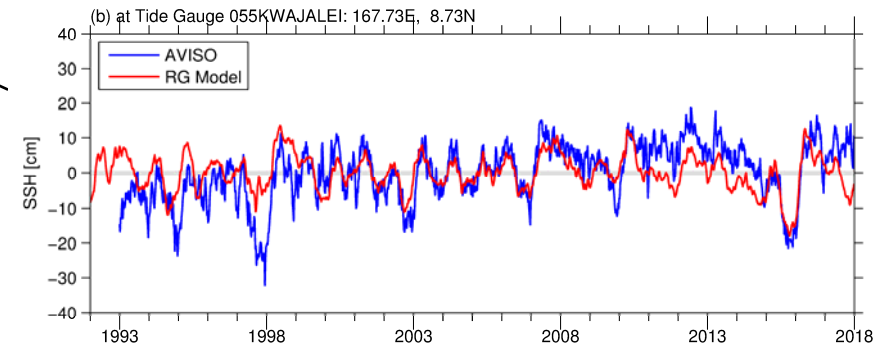
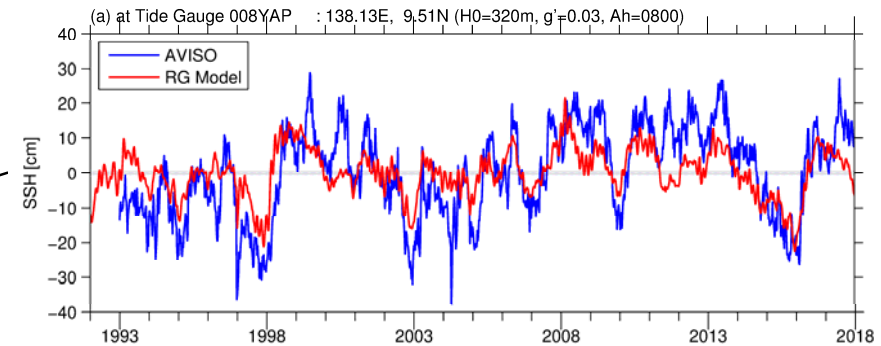
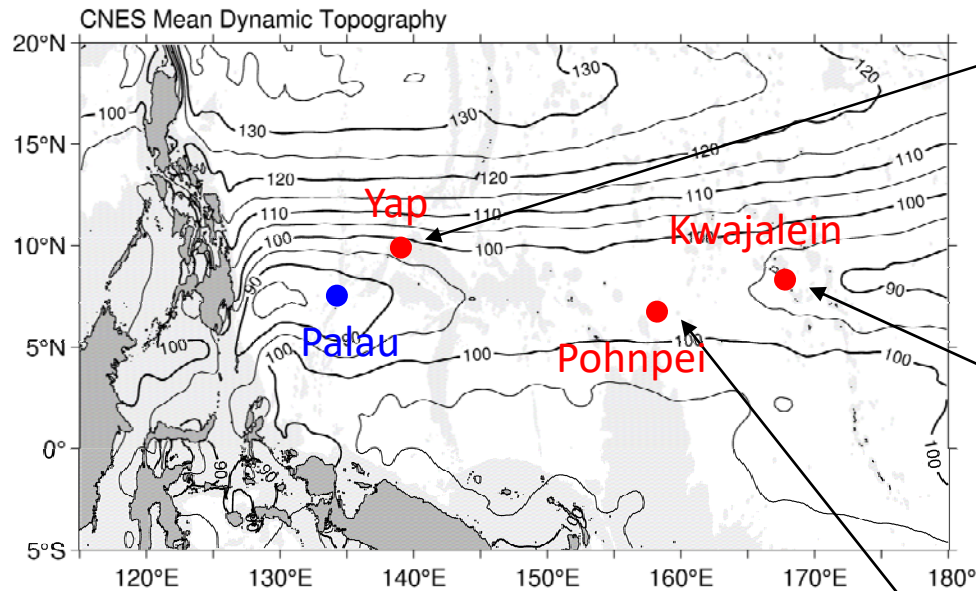


$\Delta T > 14^{\circ}\text{C}$ at Palau in < 2 months



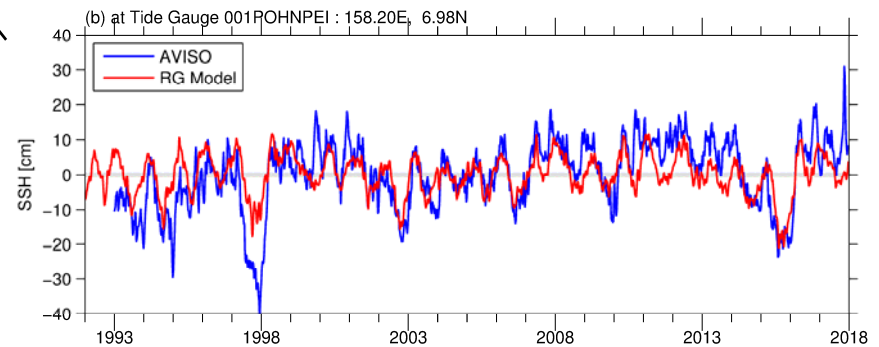
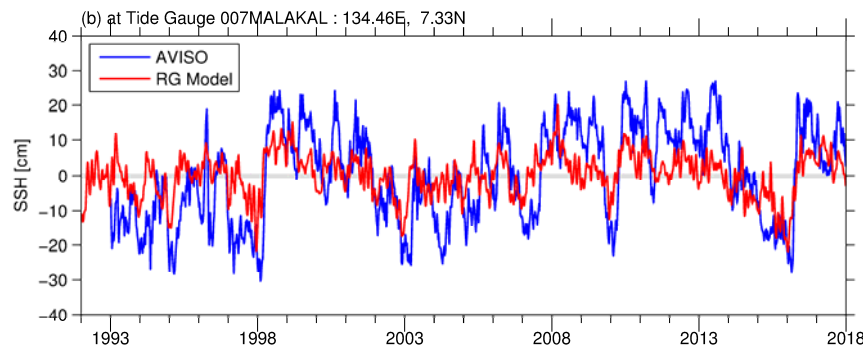
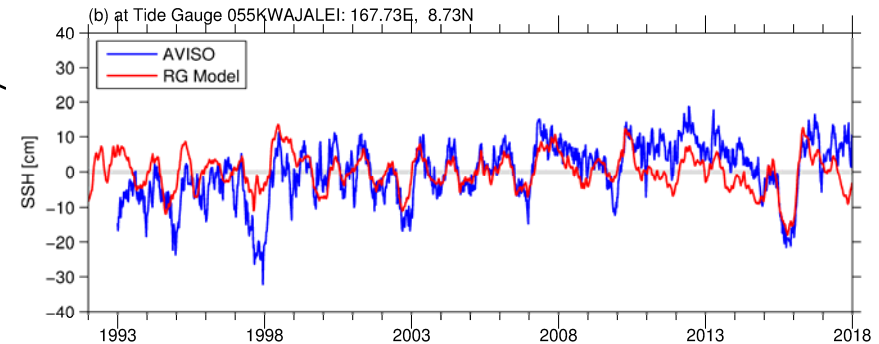
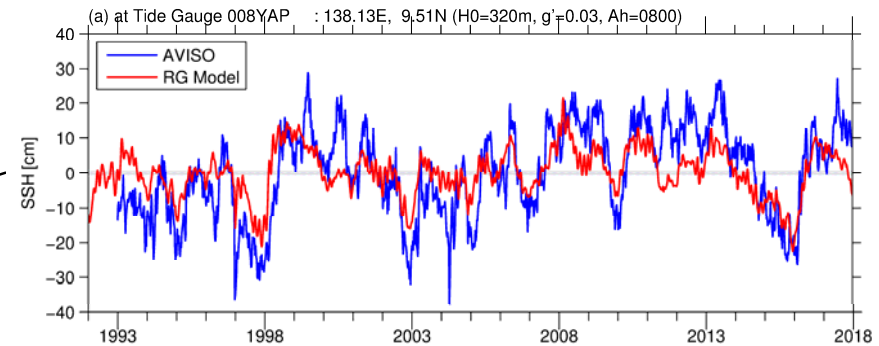
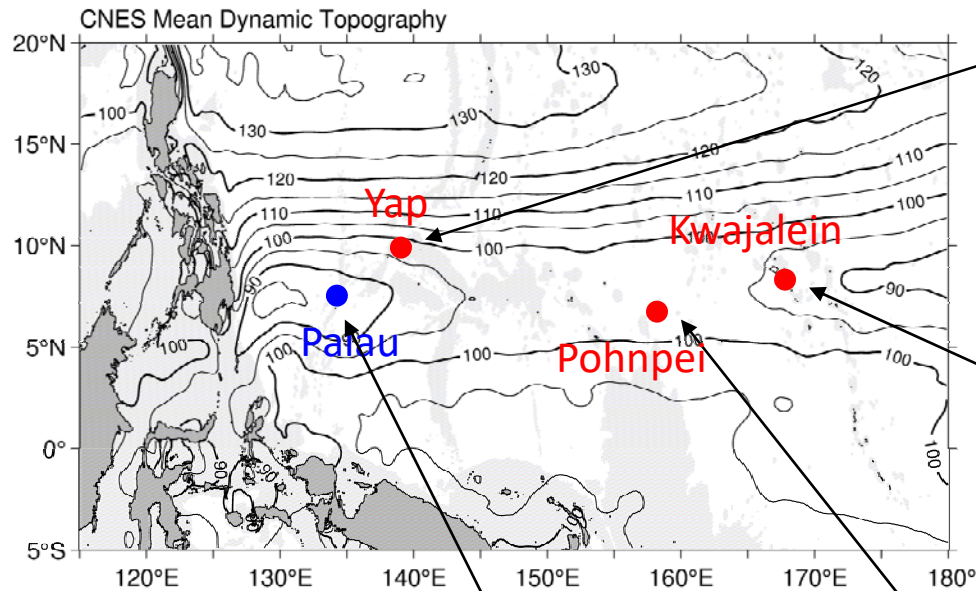
Thermal stress on shallow reef off Palau
(Colin 2018, TOS)

- Is the **linear RW dynamics** valid for upper ocean variability around **Palau** ?



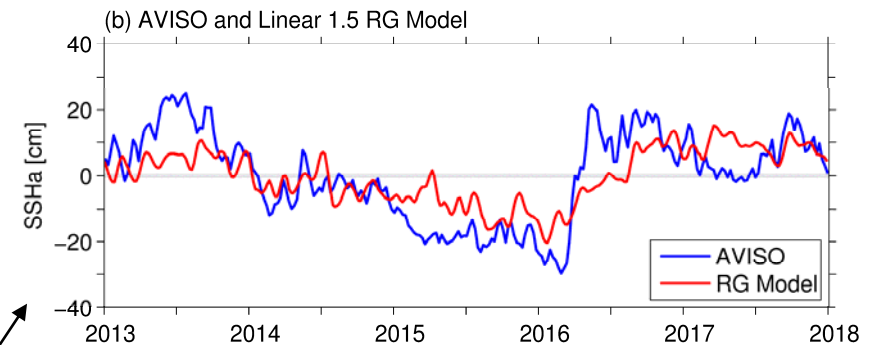
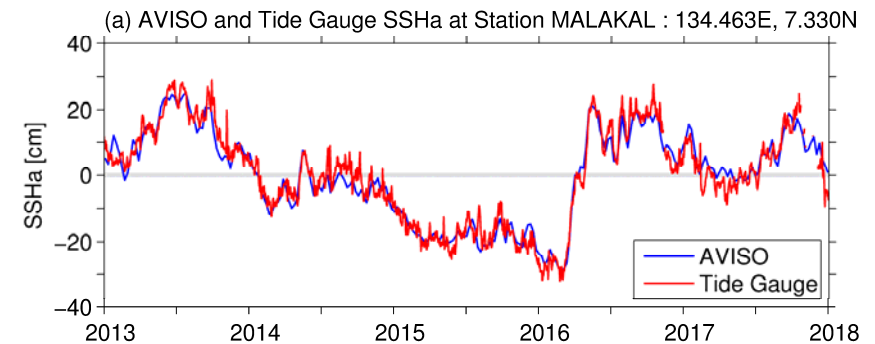
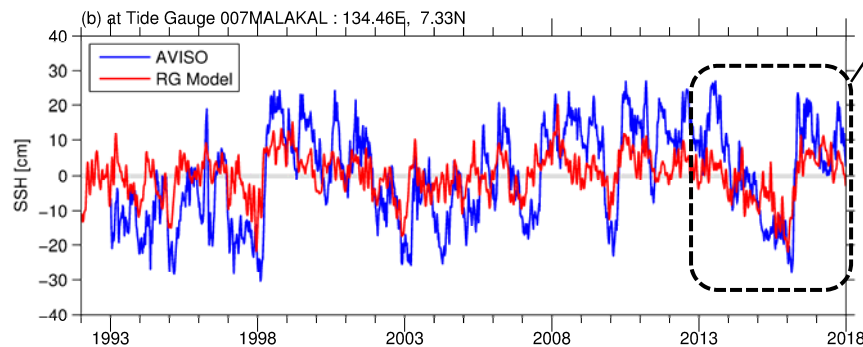
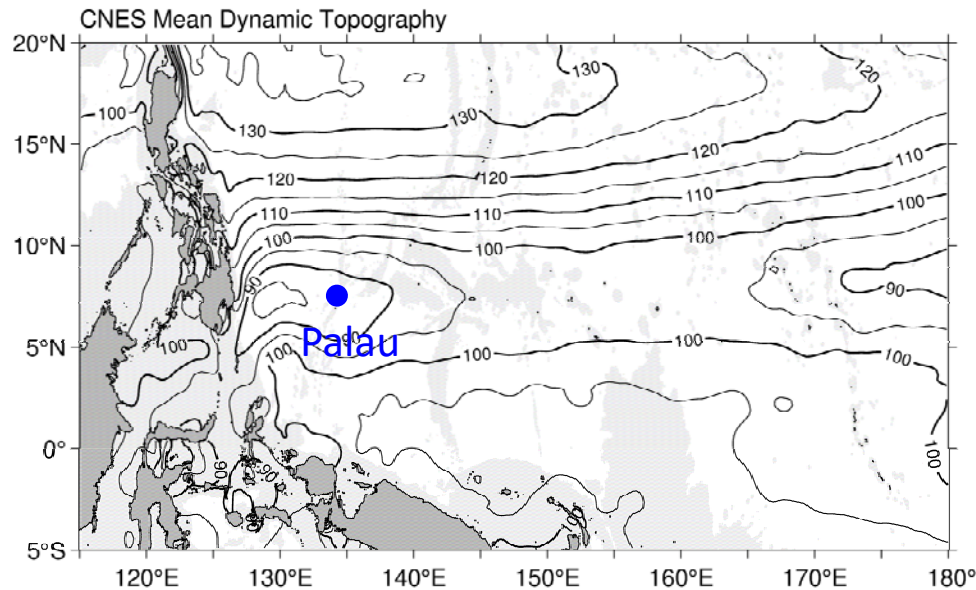
- Linear RW model** does a reasonable job hindcasting the observed SSH variability

- Is the **linear RW dynamics** valid for upper ocean variability around **Palau** ?



- At **Palau**, SSH variability is significantly underestimated by the **linear RW model**

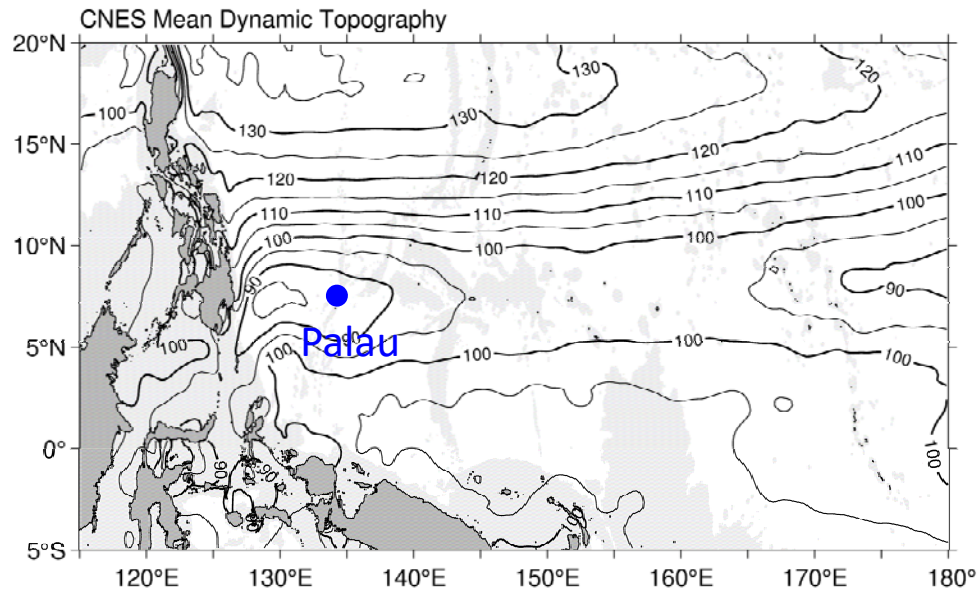
- A closer look at the upper ocean variability at **Palau** during the 2016 El Niño → La Niña transition



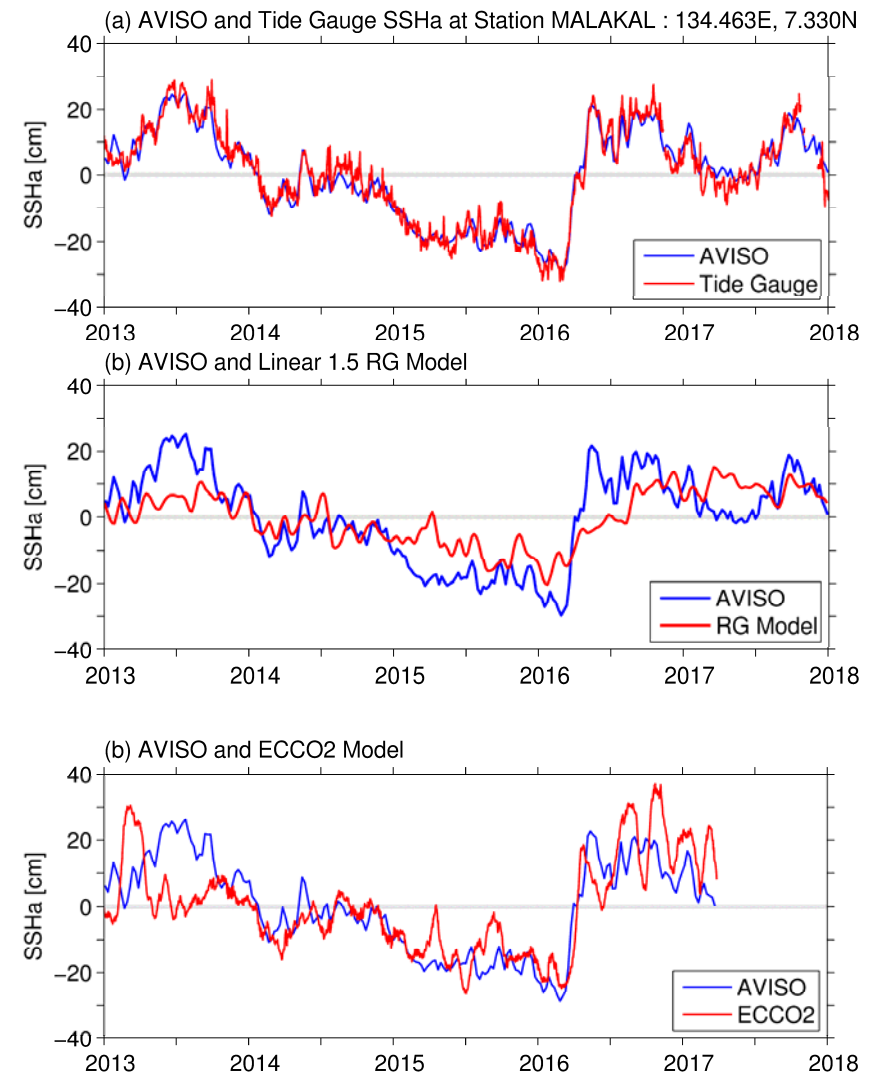
- During the El Niño → La Niña transition in early 2016, the linear model-simulated **SSH jump** is too small & too diffused!

- At **Palau**, SSH variability is significantly underestimated by the **linear RW dynamics**

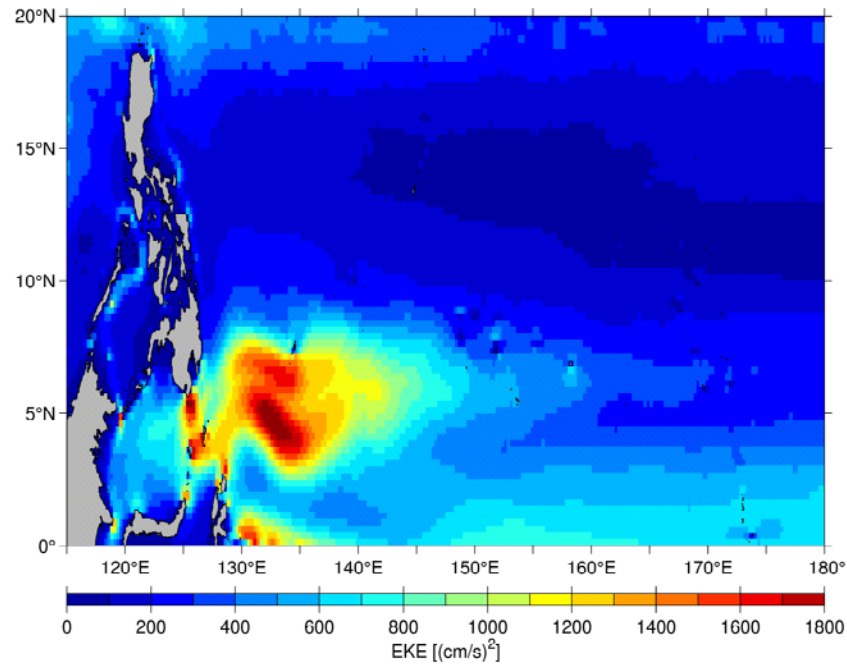
- A closer look at the upper ocean variability at **Palau** during the 2016 El Niño → La Niña transition



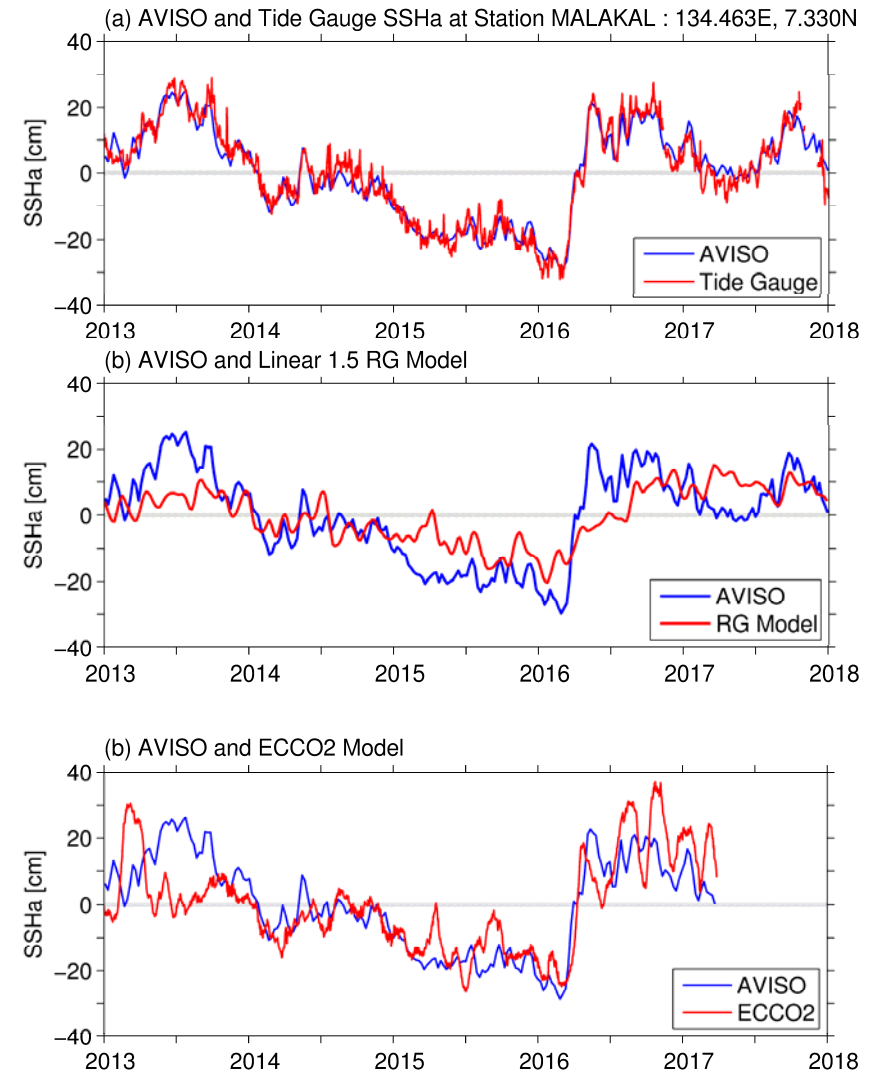
- To explore the intrinsic nonlinear processes, we analyze the **eddy-permitting ECCO2** state estimate that captured well the **short-term** variability around Palau



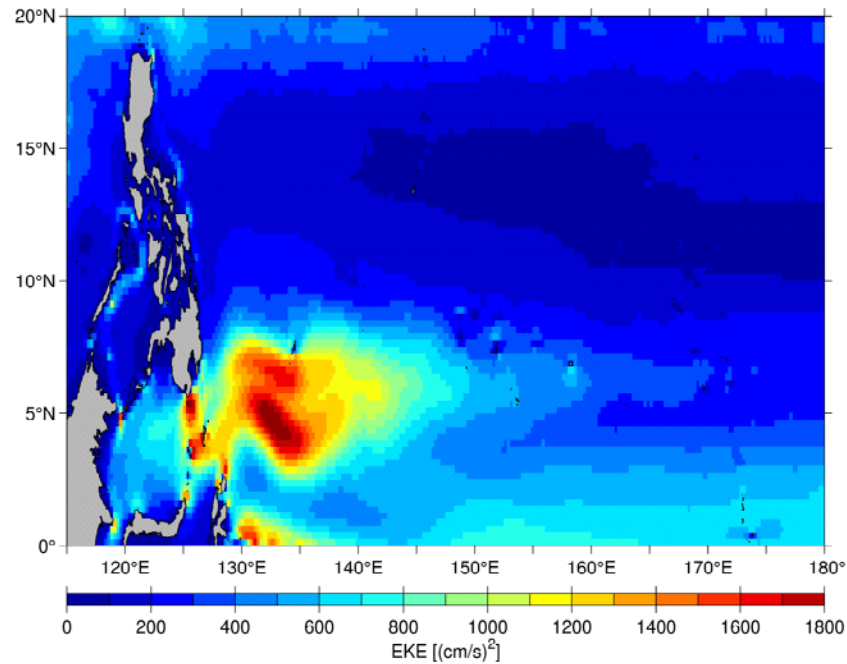
- Quantifying linear vs. nonlinear processes from the $1/4^\circ$ eddy-permitting ECCO2 state estimate



- The area surrounding Palau is where EKE has a **regional maximum** → potential **break-down** of linear vorticity dynamics

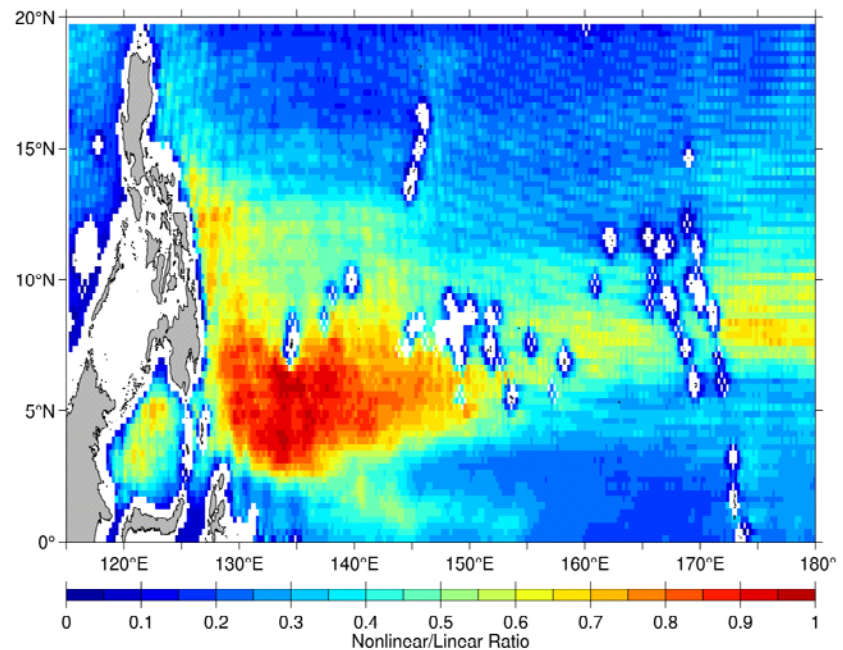


- Quantifying linear vs. nonlinear processes from the 1/4° eddy-permitting **ECCO2 state estimate**



- The area surrounding Palau is where EKE has a **regional maximum** → potential **break-down** of linear vorticity dynamics

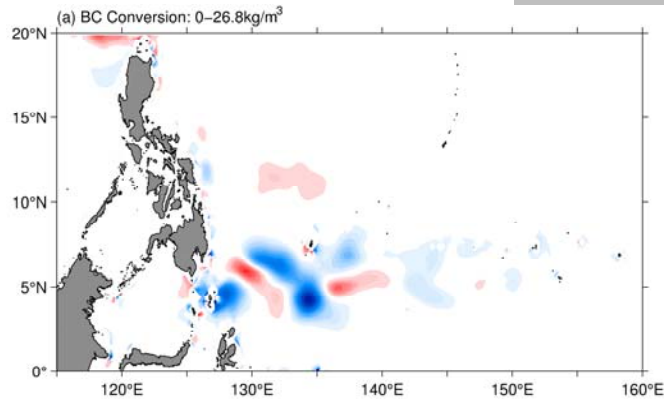
$$\frac{\partial h'}{\partial t} = \underbrace{-\nabla \cdot (\mathbf{u}'\bar{h}) - \nabla \cdot (\bar{\mathbf{u}}h')}_{\text{linear processes}} - \underbrace{\nabla \cdot (\mathbf{u}'h')}_{\text{nonlinear processes}}$$



- Ratio of nonlinear/linear processes contributing to $\partial h'/\partial t$ → the region south of Palau forms a **high-ratio hot spot**

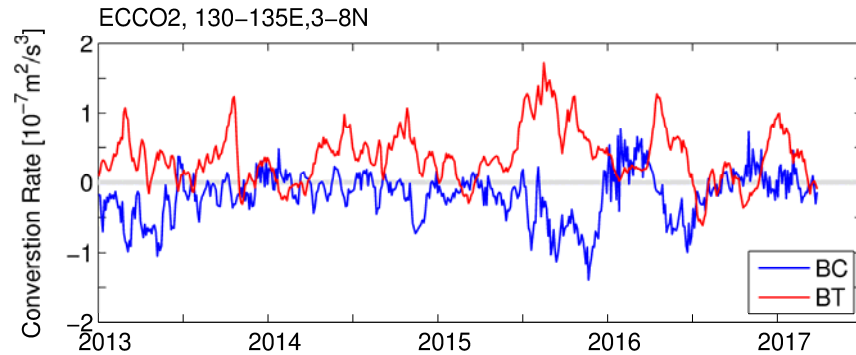
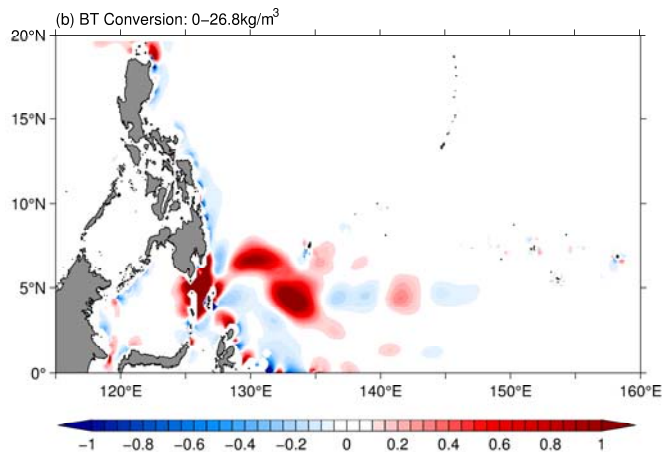
- What is the **energy source** underlying the regional eddy variability around Palau ?

baroclinic conversion rate = $-\frac{g}{\rho_o} w' \rho'$



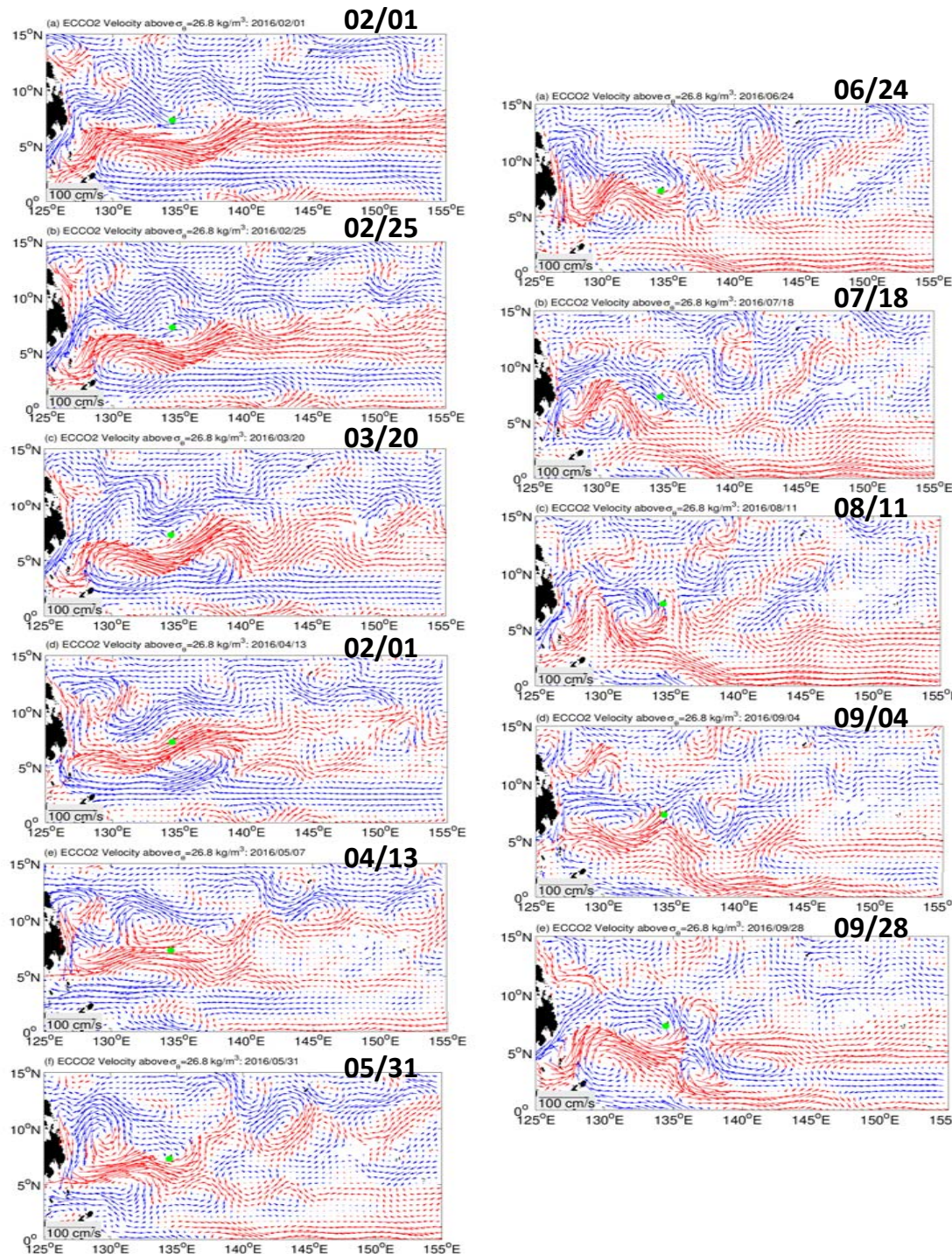
barotropic conversion rate =

$$-(\mathbf{u}'\mathbf{u}') \cdot \nabla \bar{\mathbf{u}} - (\mathbf{u}'\mathbf{v}') \cdot \nabla \bar{\mathbf{v}}$$



- Intensity of NECC's **barotropic instability** is modulated interannually; strengthened NECC during El Nino years favors stronger instability

- In the region south of Palau, eddy energy is mostly derived from mean lateral shear of NECC via **barotropic instability**



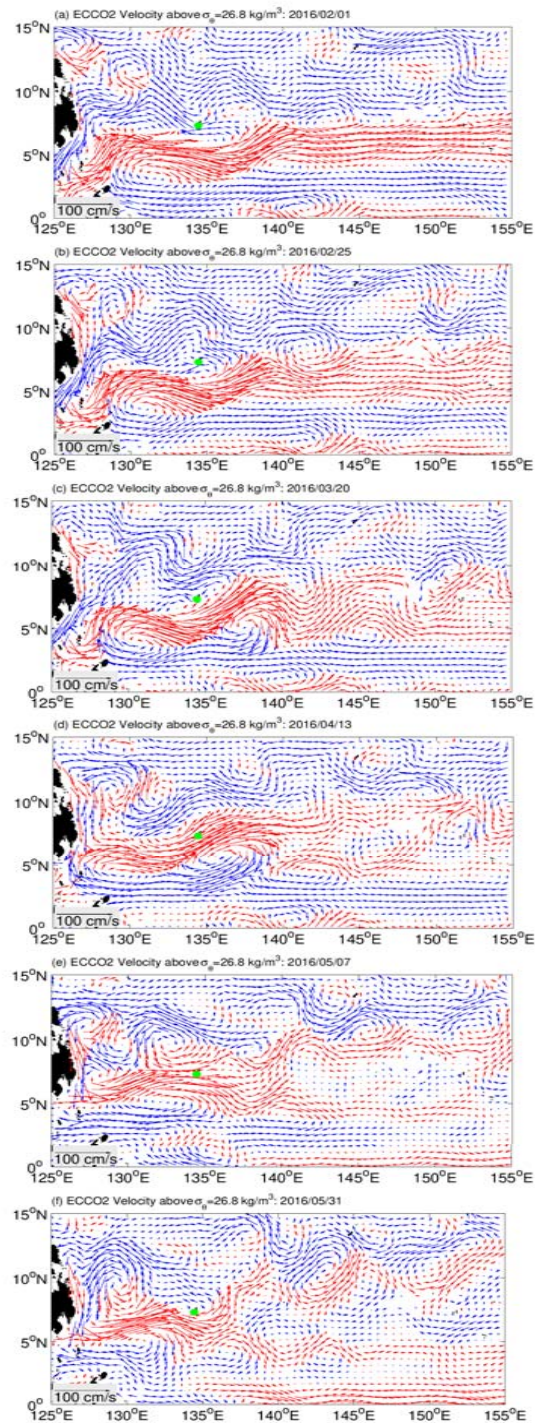
- Upper ocean evolution from the ECCO2 state estimate

- When an El Niño event ends, **trade winds recover** & generate intense off-equatorial **downwelling RWs**

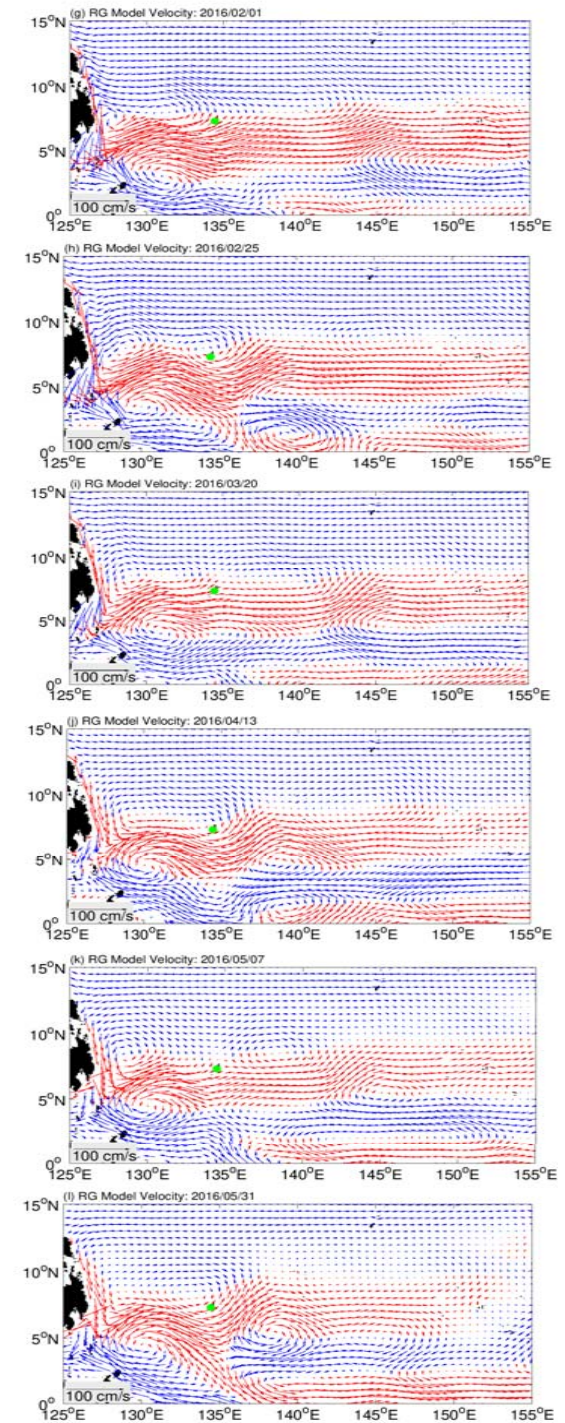
- As these intense RWs propagate westward into the Philippine Sea, they interact with horizontally-sheared NECC & lead to **barotropic instability**

- The instability breaks down the NECC, causing **short-term, large-amplitude** thermal/circulation fluctuations around Palau

- Upper ocean evolution from the ECCO2 state estimate



- Upper ocean evolution from the linear RG model



Summary

- Palau is located at the northern boundary of a “**shadow zone**” where wind-forced linear vorticity dynamics breaks down
- Barotropically-unstable NECC west of 150°E is responsible for the presence of this **shadow zone**
- Effect of **nonlinearity** is enhanced during the El Niño → La Niña transition when intense, wind-forced, downwelling RWs impinge on the seasonally poleward-migrating NECC

