

An analytical method to propagate errors in the Altimeter system

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Introduction

- Altimetry needs accurate geodetic products like: the gravity field and reference frame

=> Analysing the errors is something like understanding the transfer of errors from their origin(s) to the result (e.g. the altitude of the sea surface above a reference)

- Numerical methods are extensively used to determine a solution which is constrained by different models (or approaches)

But it is necessary to evaluate the quality, from external sources, etc. => characterize the errors in time&space

- Analytical methods have the property of providing direct relations (equations) between input parameters (ε) and the solution; e.g. :

$$\text{orbit}(\lambda, \phi, t) = \text{fcnt}(\lambda, \phi, \varepsilon(\text{geophysics, satellite, etc.}, t))$$

They have been used to characterize the Radial Orbit Error(s), using the approach of Kaula, Rosborough, etc. (see *Kaula, Balmino, Colombo, Engelis, Rosborough, Scharroo, Schrama, Tapley, Visser, and Wagner*).

=> radial orbit errors coming from the gravity field essentially, in view of understanding the limits of the altimetry system (see e.g. the so-called Geographically Correlated Errors)

- Here, we develop an analytical approach which is dedicated to the transfer of errors **from the orbit** to geodetic products: the sea surface topography is a first target, before studying the stability of tracking station coordinates.

=> Inversely, we are not able to study the impact, on the orbit, of different error sources or approaches.

Method

- In addition to the 6 classical keplerian elements (aei...), as orbital parameters, we use a set of spherical coordinates ($r'\phi'\lambda'$, strictly dedicated to circular motion) in a privileged frame:

=> the precessing mean orbital plane which is oriented by (i) and $\Omega(t)$ given along the equator

$d \text{ aei} / dt = L(\text{aei}) \delta U / \delta \text{ aie}$ => 6 equations => Kaula solution

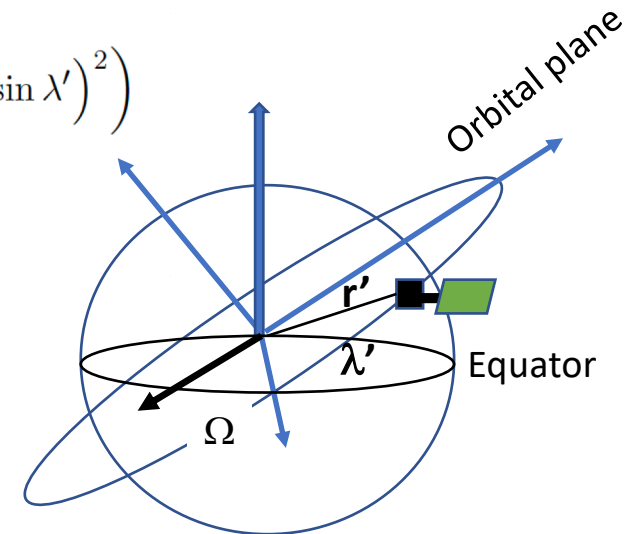
$d^2 r'\phi'\lambda' / dt^2 = G(r'\phi'\lambda'; r, \phi, \lambda) \delta U / \delta r'\phi'\lambda'$ => 3 equations of the second order in t => $r'\phi'\lambda'$ solution

$$T = \frac{1}{2} \left(\dot{r}'^2 + (r' \dot{\phi}' - r' \dot{\Omega} \sin i \cos \lambda')^2 + \left(r' \cos \phi' (\dot{\lambda}' - \dot{\Omega} \cos i) - r' \sin \phi' \dot{\Omega} \sin i \sin \lambda' \right)^2 \right)$$

with : $L = T + U(r', \phi', \lambda'; i, \Omega)$

$$\begin{cases} r' = r'_0 + \varepsilon r'_1 \\ \phi' = 0 + \varepsilon \phi'_1 \\ \lambda' = n(t - t_0) + \varepsilon \lambda'_1 \end{cases}$$

$$\begin{cases} r'_1 = \sum_{k=1}^L \frac{r_1^{ck} \cos \psi_k + r_1^{sk} \sin \psi_k}{\dot{\psi}_k} \\ \phi'_1 = \sum_{k=1}^L \frac{\varphi_1^{ck} \cos \psi_k + \varphi_1^{sk} \sin \psi_k}{\dot{\psi}_k} \\ \lambda'_1 = \sum_{k=1}^L \frac{\lambda_1^{ck} \cos \psi_k + \lambda_1^{sk} \sin \psi_k}{\dot{\psi}_k} \end{cases}$$



Method-II

- the gravity field of the solid earth, and its time variable components

$$U_{n,m} = \frac{GM}{r} \left(\frac{ae}{r}\right)^n P_{n,m}(\sin \phi) (C_{n,m} \cos m\lambda + S_{n,m} \sin \lambda)$$

where: $(C/S)_{n,m} = (C/S)_0 + (t - t_{ref})(C/S)_t + (C/S)_{\text{annual}} + (C/S)_{\text{semi-annual}}$

- the ocean tides (11 waves)

$$U_{f,n,m}^{+-} = \sum_f \sum_{+-} \frac{4\pi G a e \rho}{g} \left(\frac{ae}{r}\right)^n \left(\frac{2n+1}{1+k'_n}\right) P_{n,m}(\sin \phi) (C_{f,n,m}^{+-} \cos m\lambda + S_{f,n,m}^{+-} \sin \lambda)$$

- the solar radiation pressure considered as a field (U_{SRP})

$$\begin{aligned} U_{SRP} &= -\sigma r P_{1,0}(\cos \Psi) \\ \cos \Psi &= \sin \varphi \sin \varphi_{\odot} + \cos \varphi \cos \varphi_{\odot} \cos(\alpha_{\odot} - \lambda) \end{aligned}$$

with: $\sigma = \frac{S}{m} \cdot Cr \cdot C_{sol}$

and r is the Sun-Earth direction $(\varphi_{\odot}, \alpha_{\odot})$, and from which $\partial U_{SRP} / \partial r$ leads to the well-known acceleration.

The analytical solution, here, is used to propagate (in time):

$\delta(C_{nm}, S_{nm}, C_r, \dots)$

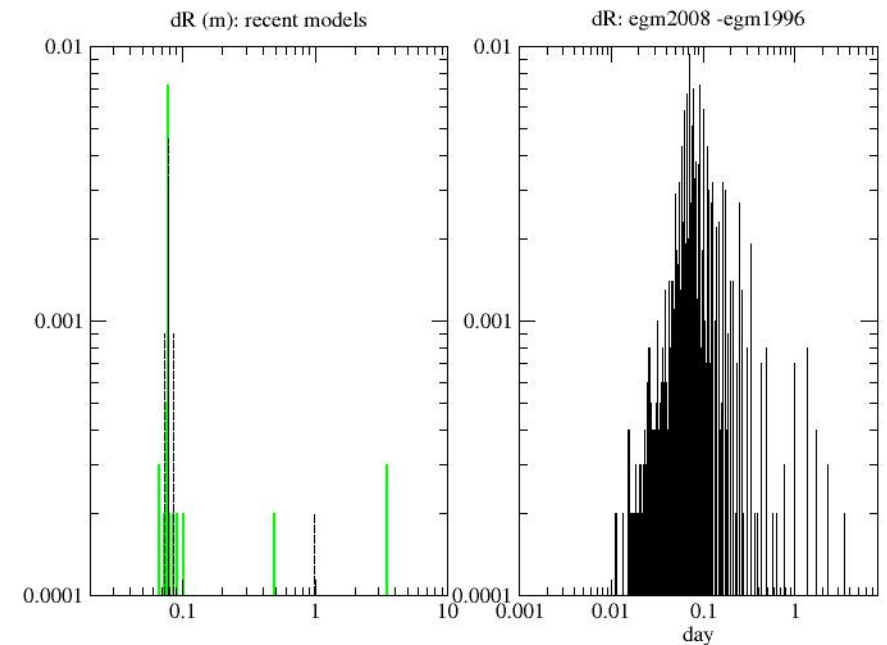
instead of

$(C_{nm}, S_{nm}, C_r, \dots)$

The “solution” directly gives orbit errors $\delta r' \phi' \lambda'$ (t) (equivalent to δRNT_{sat})

Orbit dynamics

- One of the important issue is to characterize the “errors” $\delta(C_{nm}, S_{nm}, C_r, \dots)$ used here, notably:
 - $\delta(C_{nm}, S_{nm})$ for the gravity field
 - $\delta(C_{fnm}, S_{fnm} / \text{ocean wave f})$
 - $\delta(C_r, S/m)$ for the Solar Radiation Pressure
- Gravity field model (constant and variable parts)
 - Variance/Covariance matrix
 - Differences between 2 models
 - Variable part of the geopotential (essentially annual and semi-annual terms)
- Influence of the model on precise orbits of altimetry satellites; see for recent papers: *Couhert, Rudenko, Zelensky*
- X-over analysis is used since years (see: *Rosborough, Engelis, Schrama*)



Orbit dynamics-II

Gravity field model (constant and variable parts)

Variable part of the geopotential, essentially annual and semi-annual terms: grgs-RL02

Differences between 2 models: eigen-GL04s1 and grgs-RL03-v2 (GDR-E products)

Variance/Covariance matrix (calibrated)

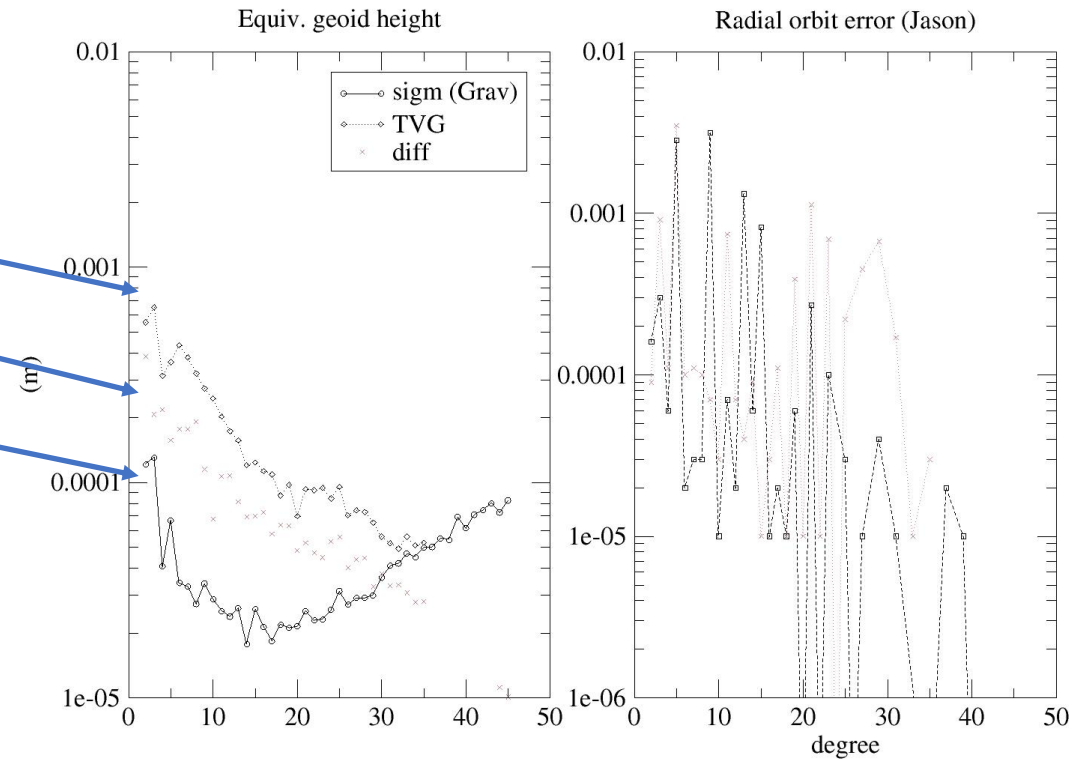
In addition:

degree 1 : δC_{10} , δC_{11} $\Leftrightarrow$ a few mm: 1.9, 1.3/1.5 at annual period

Solar Radiation Pressure

δC_r : a few %

$\delta S/m / (S/m)$: 0.007 with $[\Omega - \text{Sun-longitude}]$ (t)



Orbit dynamics-III

Ocean tide model (e.g. FES and GOT):

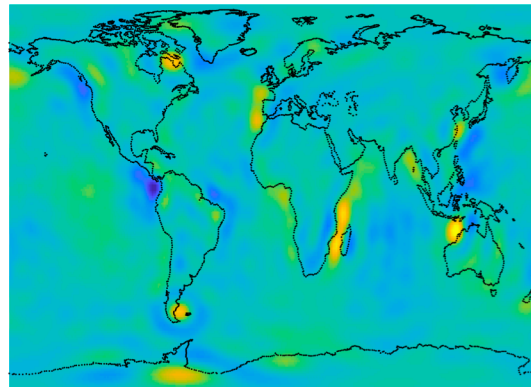
=> only part of discussions about orbit errors (see, however, the work of *Bettadpur & Eanes*)

	M2
Amplitude	~100 cm
Accuracy < 1995	2.87 cm
GOT.99v2	1.47
CSR 4.0	1.64
FES 95	1.65
FES2012	1.30 cm
# FES2012 - 2014	< 1%

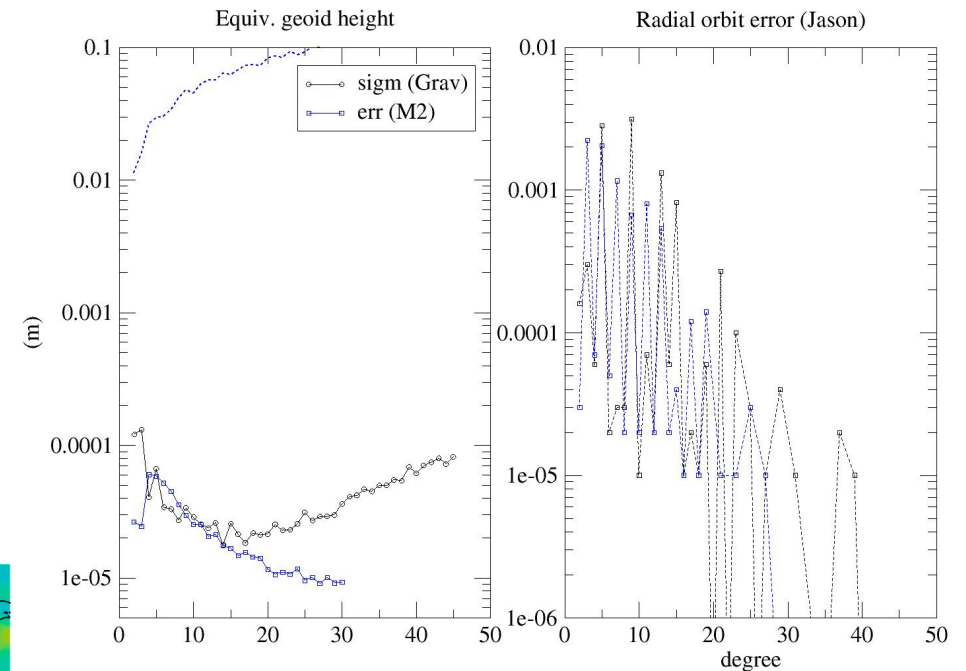
=> tide deep-ocean tide gauge

=> GRACE

Ray, Lyard: 2019, 2014



"errors" in water height => ± 8 mm



11 waves (d/o: 30)

Sa, Ssa, Mf

Q, O, P, S, K : diurnal

N, M, S, K : semi-diurnal

Models and tests

Model 1

eigen-gl04s1 : variance of SH coefficients, d/o 45

Radial orbit error (mm):

Total: mean:	0.0 ± 0.42	mean of RMS:	8.1
X-over: mean:	0.0 ± 0.2	...	1.0
Asc/Desc: mean:	$0.0/0.0 \pm 0.4$	...	8.0

Model 2

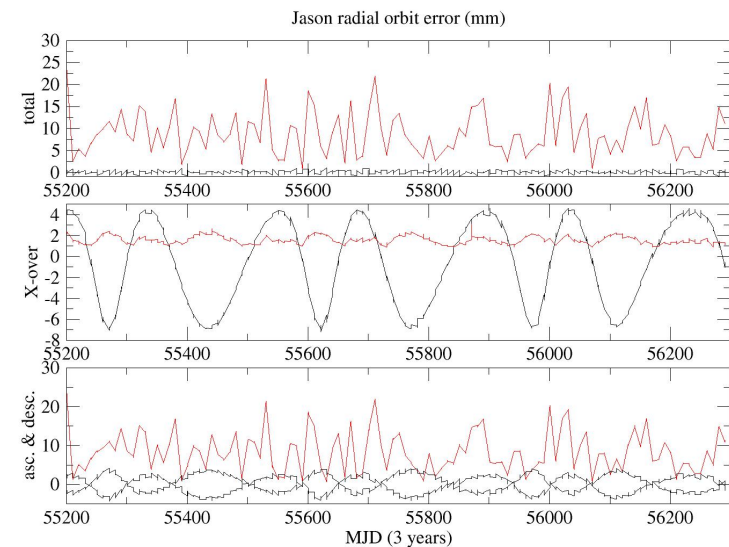
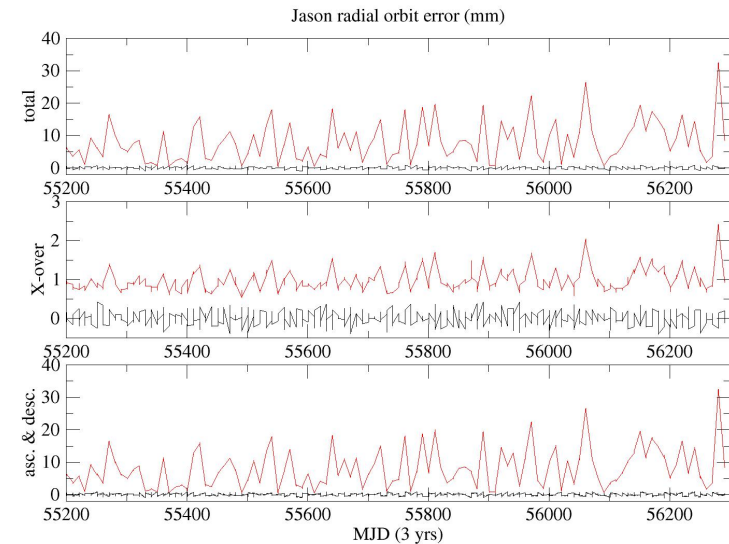
eigen-gl04s1 : variance of SH coefficients, d/o 45

C/S11: 1.3/1.5 C10: 1.9 (annual, in mm)

SRP: $\delta S/m / (S/m) + \delta Cr/Cr = 0.00694$

Radial orbit error (mm):

Total: mean:	0.0 ± 0.43	mean of RMS:	8.7
X-over: mean:	-0.6 ± 3.9	...	1.6
Asc/Desc: mean:	$-0.3/+0.3 \pm 2.0$	...	8.3



Model 1

eigen-gl04s1 : variance of coefficients, d/o 45

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Model 2

eigen-gl04s1 : variance of coefficients, d/o 45

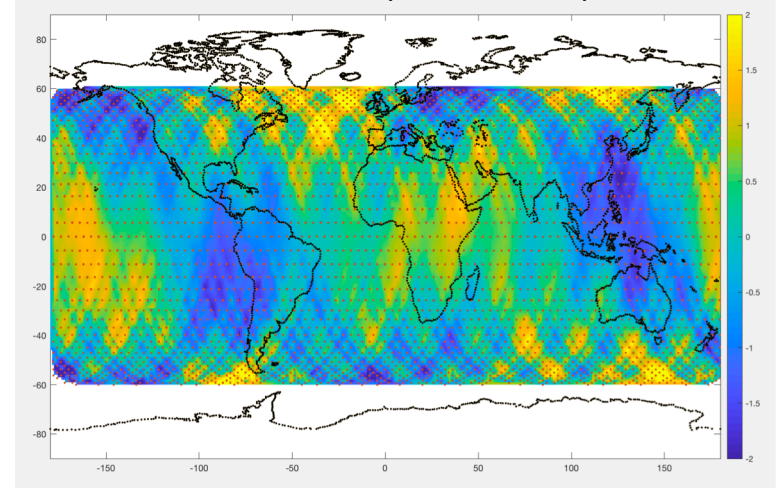
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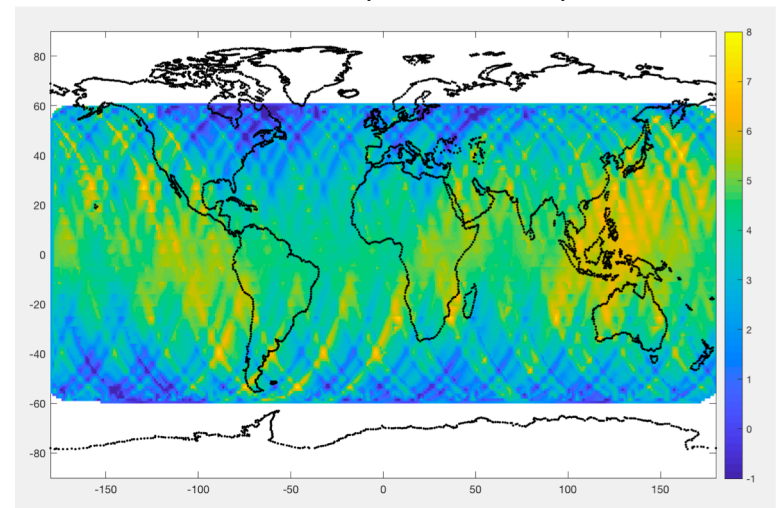
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Model 1 (X-over, mm)



Model 2 (X-over, mm)



Models and tests-II

Model 3

eigen-gl04s1 : variance of SH coefficients, d/o 45
ocean tides: 11 waves (o/ 0, 1, 2 & d/ 30)

Radial orbit error (mm):

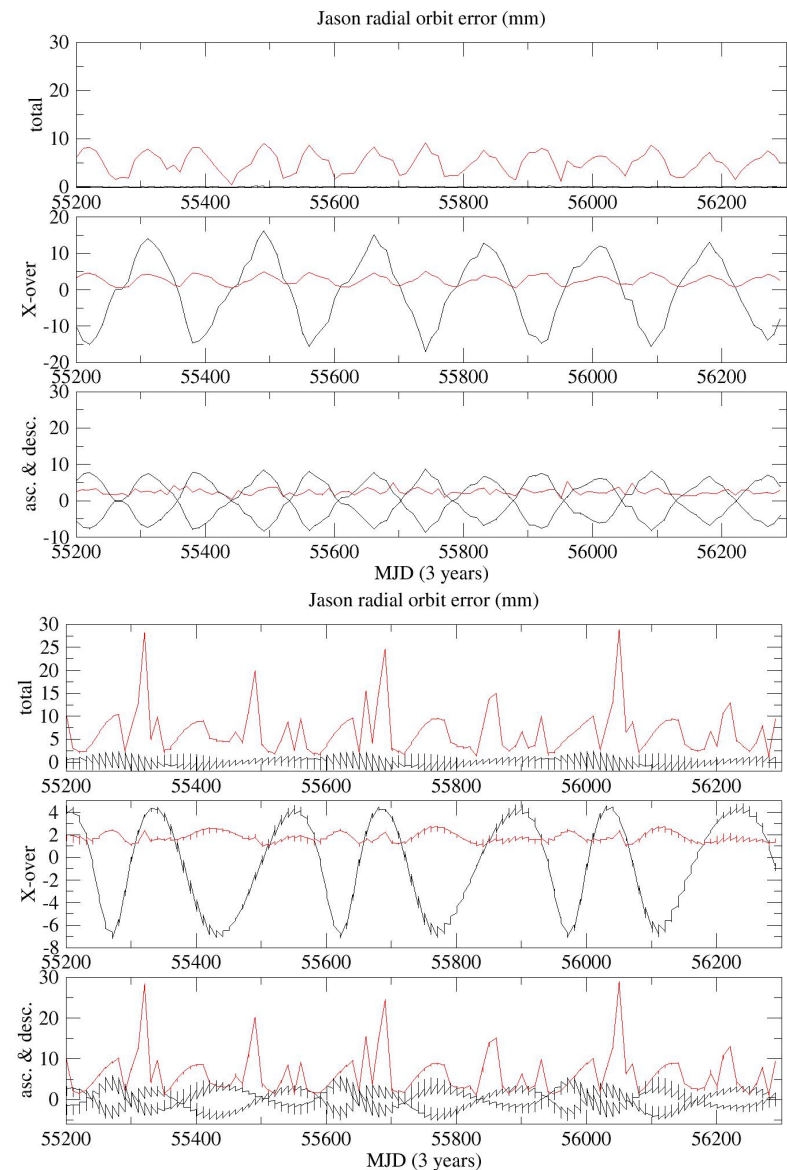
Total: mean:	0.0 ± 0.1	mean of RMS:	5.0
X-over: mean:	-0.7 ± 9.0	...	2.6
Asc/Desc: mean:	$-0.3/+0.3 \pm 4.8$	...	2.4

Model 4

grgs-RL02_MF : time var. coefficients, d/o 45
C/S11: 1.3/1.5 C10: 1.9 (annual, in mm)
SRP: $\delta S/m / (S/m) + \delta Cr/Cr = 0.00694$

Radial orbit error (mm):

Total: mean:	0.0 ± 1.2	mean of RMS:	6.7
X-over: mean:	-0.6 ± 4.0	...	1.7
Asc/Desc: mean:	$-0.3/+0.3 \pm 2.4$	...	6.3



Model 3

eigen-gl04s1 : variance of SH coefficients, d/o 45
ocean tides: 11 waves (o/ 0, 1, 2 & d/ 30)

Radial orbit error (mm):

Total: mean:	0.0 ± 0.1	mean of rms:	5.0
X-over: mean:	-0.7 ± 9.0	:	2.6
Asc/Desc: mean:	$-0.3/+0.3 \pm 4.8$	:	2.4

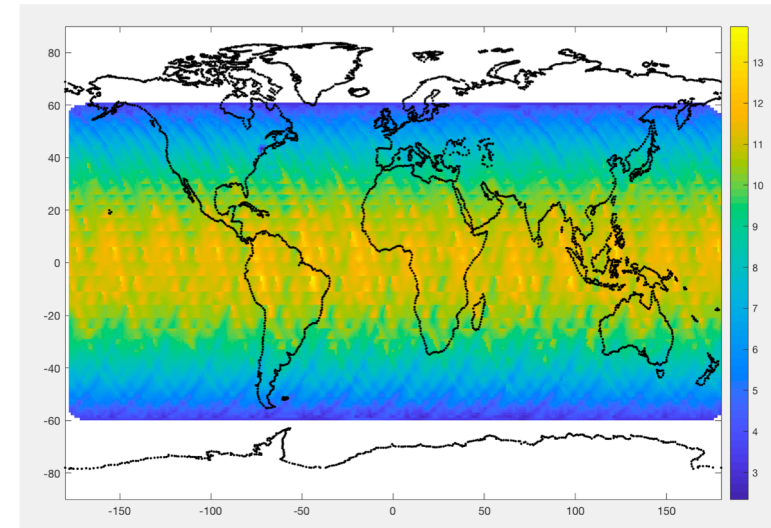
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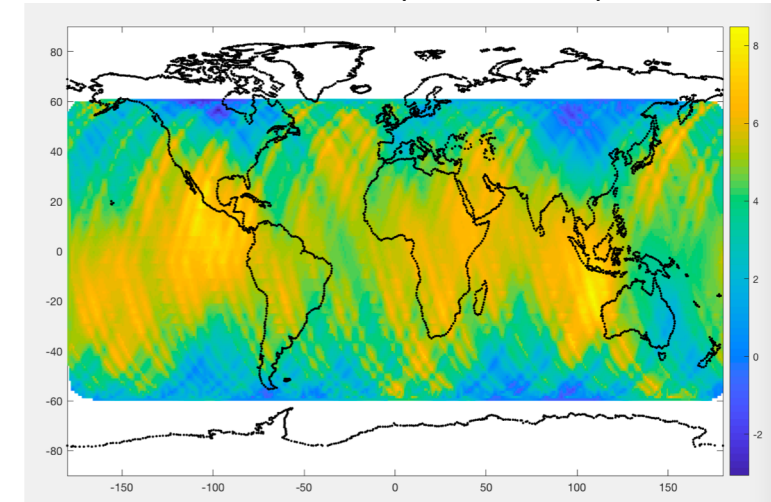
Radial orbit error (mm):

Total: mean:	0.0 ± 1.2	mean of rms:	6.7
X-over: mean:	-0.6 ± 4.0	:	1.7
Asc/Desc: mean:	$-0.3/+0.3 \pm 2.4$	:	6.3

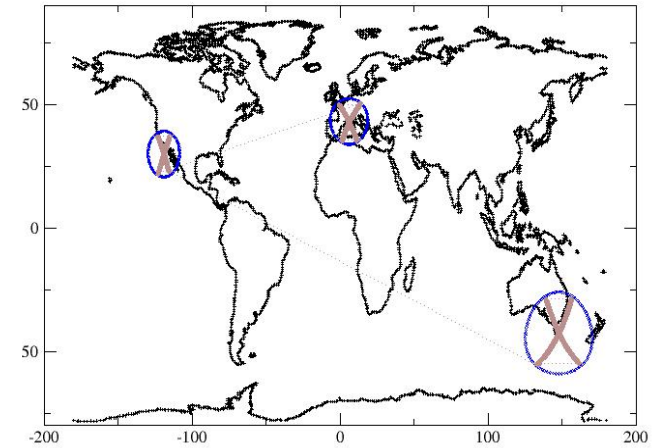
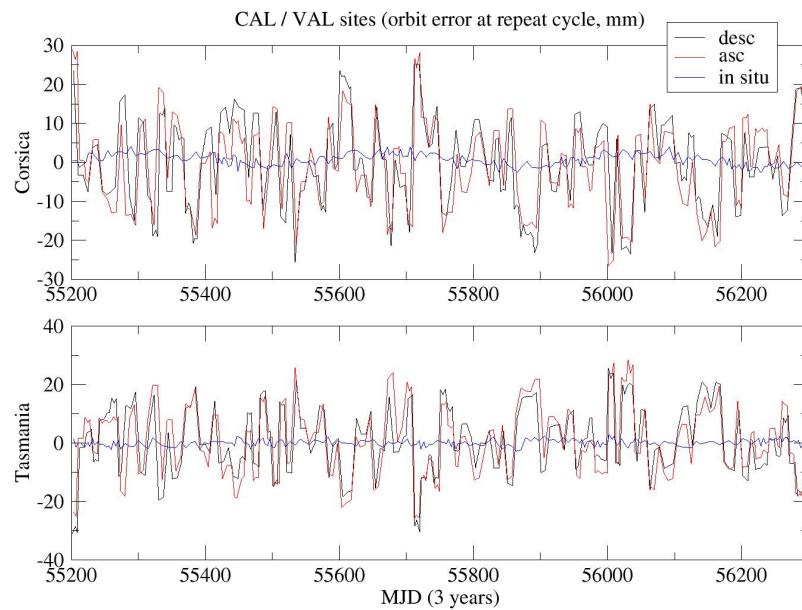
Model 3 (X-over, mm)



Model 4 (X-over, mm)



Radial orbit errors at CAL/VAL sites



Model 2

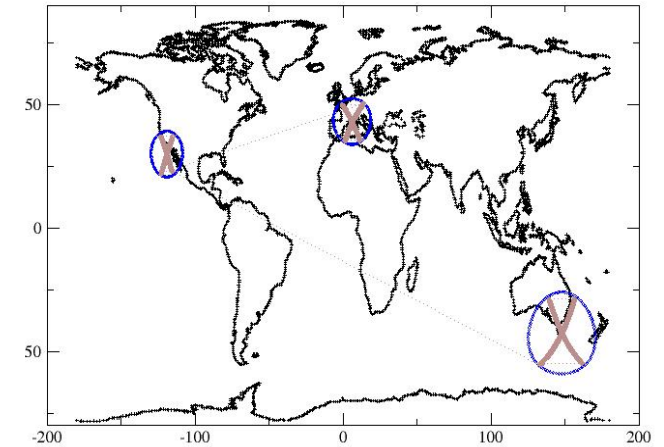
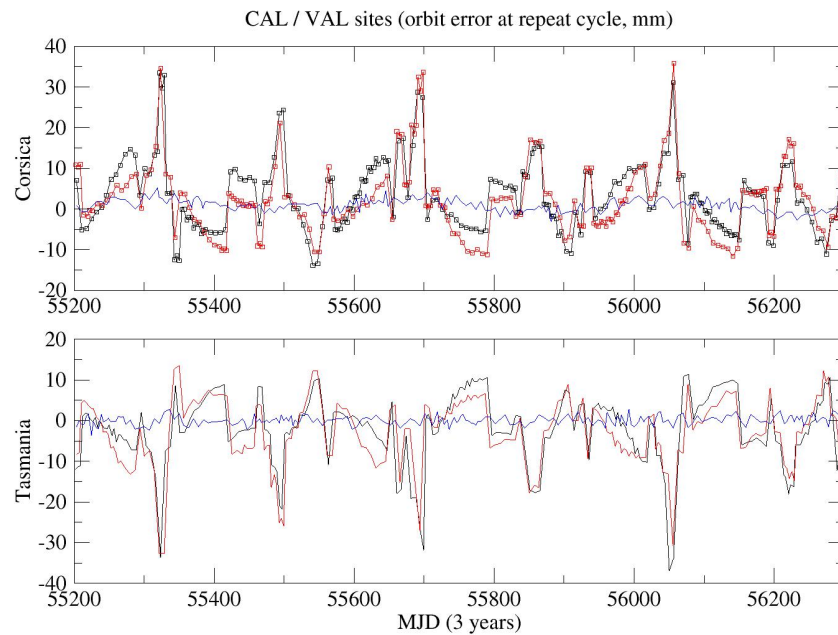
eigen-gl04s1 : variance of SH coefficients, d/o 45

C/S11: 1.3/1.5 C10: 1.9 (annual, in mm)

SRP: $\delta S/m / (S/m) + \delta Cr/Cr = 0.00694$

Corsica (repeat arc, mm)			Tasmania (id.)	
	mean	RMS	mean	RMS
Desc. arc:	-0.70	11.7	0.46	11.9
Asc. arc:	-0.10	11.8	0.57	12.3

CAL/VAL sites-II



Model 4

grgs-RL02_MF : time var. coefficients, d/o 45

C/S11: 1.3/1.5 C10: 1.9 (annual, in mm)

SRP: $\delta S/m / (S/m) + \delta Cr/Cr = 0.00694$

Corsica (repeat arc, mm)

Tasmania (id.)

	mean	RMS	mean	RMS
Desc. arc:	3.10	8.8	-2.30	8.8
Asc. arc:	2.50	8.9	-2.84	8.1

Summary (theory)

The present analytical approach is used to study the orbit errors due to several and potential origins:

- Gravity field: constant and time variable parts
- Ocean tide model: 11 waves
- Solar Radiation Pressure

After validating the first order analytical solution (spherical coordinates in the mean orbital plane) at 10^{-3} , errors in the “geodynamical” parameters are introduced in the solution equations in order to propagate the signals.

Errors in the “geodynamical” parameters are defined following different techniques:

- Variance of SH coefficients (with white noise) <= geopotential
- Part (10%) of the time variable geopotential
- Differences between existing models
- Global analysis of coefficient accuracy <= tide model
- Empirical results <= solar radiation pressure

Conclusion (results)

The present analytical approach is applied in 4 different situations:

- constant errors (variance of the geopotential)
- a slowly changing SPR (draconitic period)
- a part of the annual and semi-annual changes (hydrology) of the TVG field
- < 1% error in tide waves (global ocean tide model)

Global radial orbit error is determined at the level of 5 to 8 mm RMS (depending on the “model”)

At CAL/VAL sites, the orbit errors (repeat arcs) are of 11 mm RMS with constant part of 1 to 3 mm

The regional part of Sea Surface Height errors coming from the orbit errors is of 1 to 8 mm

Over a 3 years period, the radial drift rate is reaching 0.3 mm/yr (depending on the “model”)

A global test will be performed in order to cumulate all the different signals presented here: TVG, ocean tides and SRP

The propagation of orbit errors will be applied to study the stability of tracking station coordinates.

