Combining coastal altimetry and in situ observations to improve Meridional Overturning Circulation estimates: focus on the Southwestern Atlantic

#### Matthieu Le Hénaff<sup>1,2</sup>

Marion Kersale<sup>1,2</sup>, Christopher S. Meinen<sup>2</sup> Renellys C. Perez<sup>2</sup>, Maria Paz Chidichimo<sup>3</sup> Marcello Passaro<sup>4</sup>, Christian Schwatke<sup>4</sup> Florence Birol<sup>5</sup>, Yannice Faugère<sup>6</sup>



ARTMENT OF CO

OF MIAMI ROSENSTIEL SCHOOL of MARINE & ATMOSPHERIC SCIENCE

UNIVERSITY

UM-CIMAS, Miami, FL
 NOAA-AOML, Miami, FL
 SHN, Buenos Aires, Argentina
 DGF-TUM, Germany
 LEGOS, Toulouse, France
 CLS, Toulouse, France

# Outline:

- 1. Motivation
- 2. Methodology and data
- 3. Initial results and discussion
- 4. Concluding remarks

# Motivation



- The Atlantic **Meridional Overturning Circulation** (MOC) is a crucial component of the **climate system**
- The North Atlantic is a region of deep water formation
- The **South Atlantic** is a **nexus** for **water masses formed remotely**: Indian and Pacific Oceans, Antarctic bottom waters, North Atlantic Deepwater

The South Atlantic is the only basin where heat is fransported equatorward (and not poleward)

Schematic of the main MOC in the Atlantic (Perez et al., 2015). Blue arrows: cold, dense water in the lower limb of the MOC cell. Red arrows: warm, light waters in the upper limb of the cell.

- North Atlantic (26.5 deg.N): RAPID-MOCHA array since 2004
- South Atlantic MOC Basin-wide Array (SAMBA, 34.5 deg.S) since 2009



Location of moorings of the 34.5 deg.S array, with U.S., Brazilian, French, and South African moorings.

SAMBA array mostly composed of Pressure Inversed Echo Sounders
 PIES/CPIES moored instruments: acoustic measurements (travel time) + look up tables from *in situ* data:

### $\Rightarrow$ vertical thermohaline structure

- Between 2 moorings: integrated geostrophic current anomaly
- Dense array at both western and eastern boundaries
- **Total MOC transport** (geostrophic + Ekman) between most eastern and western moorings: **14.7 Sv**, standard deviation **8.3 Sv** (Meinen et al., 2018)

# **Motivation**



- Limitations of the SAMBA array:
  - Most inshore moorings at ~1000 m depth
  - No continuous measurements inshore
- Currently, **transport inshore** of the moorings estimated using **model simulations**
- Inshore transport: ~3 Sv, i.e. ~20% of the total MOC estimated without direct observations

Surface velocities (blue arrows) from shipboard-ADCP in 2009. Locations of the PIES/CPIES instruments (yellow + magenta)

• We proposed to use **coastal altimetry** to **fill the gap** between the most inshore moorings and the coasts:

Project "Combining coastal altimetry and in situ observations to improve Meridional Overturning Circulation estimates in the South Atlantic", supported by NASA Physical Oceanography Program

#### Methodology and data



Jason tracks (magenta), with, in red, tracks used in our project. In yellow and cyan: PIES and CPIES, in green: tall moorings

- We are using the **reference TOPEX/Poseidon-Jason** series **tracks**:
  - Track #254 in the West along South America
  - Track #209 in the East along South Africa
- Focus on **Jason-2** data first (2008-2016, longest overlap with SAMBA array), to be extended to T/P, Jason-1 and Jason-3

### • Along-track Sea Level Anomaly, then Geostrophic Current Anomaly

- *In situ* observations: barotropic and 1<sup>st</sup> baroclinic modes => integrate the Geostrophic Current Anomaly => **Geostrophic Transport Anomaly**
- Mean geostrophic transport: MDT or model outputs (if MDT not reliable)
- Ageostrophic component: **Ekman transport** from atmospheric product

# Methodology and data

# • Coastal altimetry data from various sources:

- Centre of Topography of the Oceans and the Hydrosphere (CTOH, F. Birol)
- **PISTACH/PEACHI** products from CLS (Y. Faugère)
- ALES retracker data (M. Passaro and C. Schwatke)
- **GPD+** wet tropospheric correction (J. Fernandes)
- Advantages of using coastal altimetry datasets:
  - Closer to the coast
  - **High-frequency** data (20 Hz): allows higher spatial resolution
  - Improved retracking methods (ALES, PISTACH/PEACHI)
  - Improved geophysical corrections: MSS (CTOH), Tidal signal (CTOH), wet troposphere (GPD+), DAC (PISTACH/PEACHI)
- We combine these datasets to:
  - Take advantage of the strengths of each dataset
  - Estimate **uncertainties** from the various available coastal altimetry data and their combination



- CTOH: correlation 0.51; Root-Mean Square Error (RMSE) 9.62 cm
- PEACHI: correlation 0.53, RMSE 10.79 cm
- $\Rightarrow$  better correlation for PEACHI, lower error with CTOH data
- Root Mean Square Difference between CTOH and PEACHI = 3.79 cm
  ⇒ suggests most errors are shared between datasets
- Both datasets capture 2010 low values and 2015 high values

#### Initial results: focus on the Southwestern Atlantic **Comparison** of **geostrophic current anomalies** along the track: Geostro<u>phic current anomaly, CTOH</u> r<u>ef</u>. data South of 34.5 deg.S: <sup>100</sup><sub>80</sub> - Variability associated with **Brazil Current** <sup>20</sup> - Signature **more intense in** -20 **PEACHI** data • North of 34.5 deg.S: - CTOH data extend closer to the coast Geostrophic current anomaly. PEACHI data - Positive anomaly band around 34 deg.S in **PEACHI** 🖞 data; reason not clear - Root Mean Square **Difference** between both datasets: 11.9 cm/s, -60 -80 <sup>100</sup> comparable to RMSE in other coastal regions (Le J M M J S N J M M

Geostrophic Current Anomalies (cm/s) from CTOH (top) Hénaff et al., 2011) and PEACHI (bottom). The dashed line is at 34.5 deg.S.

# **Concluding remarks and next steps**

# Initial results are promising:

- Capacity of the altimetric data to capture observed variability
- Differences in geostrophic currents between datasets in the order of typical observation errors in geostrophic currents

# Next steps:

- Extend to high-frequency data and ALES data (work in progress)
- Implement **specific corrections** (esp. harmonic analysis for the tidal correction, GPD+ wet tropospheric correction)
- Analysis of *in-situ* data to estimate the vertical structure and how to project surface geostrophic current on the vertical to estimate transport. We have already gathered the *in situ* data of the region.
- Perform **comparable analysis** in the **Eastern boundary** (when approach finalized in the Western boundary), to allow refined **estimates of the MOC**
- More results at next OSTST Meeting!



